

BRIGHAM YOUNG UNIVERSITY

GEOLOGY STUDIES

VOL. 24 PART 2

DEC. 1971



Brigham Young University Geology Studies

Volume 24, Part 2

CONTENTS

Studies for Students: A Question Set for Sands and Sandstones	Sedimentation Seminar
A New Species of <i>Arroyocrinus</i> (<i>Inadunata</i>) from the Park City Formation (Upper Permian) of Utah	H. L. Strimple and J. F. Miller
Additional Specimens of the Hypsilophodontid Dinosaur <i>Dryosaurus altus</i> from the Upper Jurassic of Western North America	Jeffrey D. Shepherd, Peter M. Galton, and James A. Jensen
Paleoenvironments of the Moenave Formation, St. George, Utah	John Daniel Davis
Foraminiferal Abundance Related to Bentonitic Ash Beds in the Tununk Member of the Mancos Shale (Cretaceous) in Southeastern Utah	Rebecca Lillywhite Bagshaw
Paleoecology of the Lower Carmel Formation of the San Rafael Swell, Emery County, Utah	Lawrence H. Bagshaw
Structure, Stratigraphy, and Tectonic History of the Indianola Quadrangle, Central Utah	David M. Runyon
Compound Faceted Spurs and Recurrent Movement in the Wasatch Fault Zone, North Central Utah	Thomas C. Anderson
A Subsurface Correlation of Permian-Triassic Strata in Lisbon Valley, Utah	Ralph T. Bohn
Mesozoic-Cenozoic Structural Development of the Kern Mountains, Eastern Nevada-Western Utah	Robert C. Ahlborn
Publications and Maps of the Geology Department	



Cover: Sahara dune sand, X130. Photo courtesy Amoco Production Research, Tulsa, Oklahoma, contributed by Paul E. Potter, Department of Geology, University of Cincinnati, Cincinnati, Ohio 45221.

A publication of the
Department of Geology
Brigham Young University
Provo, Utah 84602

Editors

W. Kenneth Hamblin
Cynthia M. Gardner

Brigham Young University Geology Studies is published semiannually by the department. *Geology Studies* consists of graduate-student and staff research in the department and occasional papers from other contributors. *Studies for Students* supplements the regular issues and is intended as a series of short papers of general interest which may serve as guides to the geology of Utah for beginning students and laymen.

ISSN 0068-1016

Distributed December 1977

Price \$5.00

(Subject to change without notice)

12-77 525 28856

CONTENTS

STUDIES FOR STUDENTS: A QUESTION SET FOR SANDS AND SANDSTONES	1	PALEOENVIRONMENTS OF THE MOENAVE FORMATION, ST. GEORGE, UTAH	17
I. Questions primarily answered in the laboratory	1	Introduction	17
II. Questions primarily answered by the megascopic examination of outcrops and cores and by the study of wire line logs	4	Previous work	17
III. Format for a basin—Questions primarily answered by study of the entire basin or perhaps even a larger area such as a continental margin	6	Location	18
Figures		Methods of study	18
1. Definition sketch of a typical sandstone	2	Acknowledgments	18
2. Minerals filling sandstone interstices	3	Geologic setting	19
3. Flow chart using wire line logs	5	General sedimentary features	19
4. Flow chart for basin analysis	6	Geometry	21
Tables		Channel-filling mechanisms	21
1. Classification of sand and sandstone bodies	4	Composition	23
2. Format for basin study: properties and methods	7	Sedimentary structures	23
A NEW SPECIES OF <i>ARROYOCRINUS</i> (INADUNATA) FROM THE PARK CITY FORMATION (UPPER PERMIAN) OF UTAH	9	Paleocurrent data	24
Introduction	9	Fossils	26
Systematic paleontology	9	Interpretation and reconstruction of the Moenave environment of deposition	27
Description	9	Interpretation I	27
Discussion	9	Interpretation II	28
Measurements of holotype in millimeters	10	Reconstruction of the Moenave tidal flat environment	29
Occurrence	10	Summary	30
Holotype	10	References cited	30
Acknowledgments	10	Figures	
References	10	1. Index map	18
Figures		2. Moenave Formation stratigraphy	19
1. Photographs	9	3. North channel in airport roadcut	20
2. Camera lucida drawings	10	4. South channel in airport roadcut	20
ADDITIONAL SPECIMENS OF THE HYPSILOPHODONTID DINOSAUR <i>DYROS AURUS ALTUS</i> FROM THE UPPER JURASSIC OF WESTERN NORTH AMERICA	11	5. Schematic drawing of tabular appearance of massive mudstone layers which enclose the channel deposits	20
Introduction	11	6. Bedding plane inclinations in channels	21
Data on specimens	11	7. Bundles of ribbon sand bodies at Interstate 15 road- cut	22
Acknowledgments	11	8. Schematic drawing of ribbon sand bodies, Interstate 15 roadcut	22
Description and comparisons	11	9. Intraformational conglomerate lens	23
Vertebral column	11	10. Flaser bedding	23
Pectoral girdle	12	11. Wavy bedding	23
Fore limb	12	12. Planar cross-bedding	24
Pelvic girdle	12	13. Small-scale trough type cross-bedding	24
Hind limb	13	14. Paleocurrent trends in channel undulations	24
Discussion	14	15. Vector diagrams of paleocurrent trends	25
References cited	15	16. Location map of paleocurrent measurements along the erosional cliffs north of St. George Blvd.	26
Figures		17. Location map of paleocurrent measurements for posi- tions G and H	26
1. Vertebrae	12	18. Paleocurrent summary diagram of positions A-J	27
2. Pectoral girdle and forelimb	13	19. Vertical burrows	28
3. Pelvic girdle	13	20. Horizontal burrows	28
4. Femora and tibiae	14	21. Burrows preserved as cavities	28
5. Pes	15	22. Reconstructed model diagram of the Moenave tidal flat in the St. George area	29
Table		23. Regional paleocurrent trends	30
1. Summary of current flow data	25	FORAMINIFERAL ABUNDANCE RELATED TO BENTONITIC ASH BEDS IN THE TUNUNK	

MEMBER OF THE MANCOS SHALE (CRETACEOUS) IN SOUTHEASTERN UTAH	33		
Introduction	33	Autecology	59
Acknowledgments	33	<i>Trigonia</i> and <i>Vaugonia</i>	
Previous work	33	<i>Camptonectes</i>	
Field and laboratory methods	34	<i>Gryphaea</i> (?)	
Fauna	34	<i>Modiolus</i> (?) and <i>Nucula</i> (?)	
Paleoecology	35	Environment	59
Conclusions	39	General	59
Systematic paleontology	39	Salinity	59
References cited	48	Temperature	59
Figures		Interpretation of sedimentary environments	59
1. Stratigraphic section	33	Tidal flat	59
2. Index map	34	Open marine	60
3. View of Tununk stratigraphy	35	Barrier lagoon	60
4. View of trench through ash 2	35	Sedimentary model	60
5. Graph of foraminiferal abundance in ash 1	36	Conclusions	61
6. Graph of foraminiferal abundance in ash 2	37	Appendix	61
7. Graph of foraminiferal abundance in ash 3	38	References cited	62
8. Graph of foraminiferal abundance in ash 4	39	Figures	
9. Graph of foraminiferal abundance in ash 5	40	1. Index map	51
10. Foraminiferal fauna	42	2. Stratigraphic section	52
11. Foraminiferal fauna	44	3. Generalized east-west Carmel time cross section	53
12. Foraminiferal fauna	46	4. Paleogeographic setting of the study area	53
13. Foraminiferal fauna	47	5. Field views	54
Tables		6. Photomicrographs	55
1. Foraminiferal occurrence in ash 1	36	7. Stratigraphic section	56
2. Foraminiferal occurrence in ash 2	37	8. Fossils and an unusual sedimentary structure	58
3. Foraminiferal occurrence in ash 3	38	9. Sedimentary model	60
4. Foraminiferal occurrence in ash 4	39		
5. Foraminiferal occurrence in ash 5	41		
PALEOECOLOGY OF THE LOWER CARMEL FORMATION OF THE SAN RAFAEL SWELL, EMERY COUNTY, UTAH	51	STRUCTURE, STRATIGRAPHY, AND TECTONIC HISTORY OF THE INDIANOLA QUADRANGLE, CENTRAL UTAH	63
Introduction	51	Introduction	63
Acknowledgments	51	Acknowledgments	63
Location	51	Previous work	63
Methods of study	51	Geologic evolution and regional setting	64
Previous work	52	Present-day setting	65
Geologic setting	52	Stratigraphy	65
Lithologies	53	Jurassic	66
Sandstone	53	Arapien Shale	66
Light yellow brown sandstone	53	Twist Gulch Formation	67
Medium brown sandstone	53	Cretaceous	67
Siltstone	53	Indianola Group	67
Red siltstone	53	South Flat Formation	68
Gypsiferous and dolomitic siltstone	55	Price River Formation	69
Shale	55	Cretaceous-Tertiary	70
Green shale	55	North Horn Formation	70
Grey shale	55	Tertiary	71
Red shale	55	Flagstaff Formation	71
Carbonate rocks	55	Green River Formation	72
Calcareneite	56	Unnamed volcanic rocks	73
Dololutite	56	Quaternary	74
Dolarenite	56	Unnamed	74
Sedimentary structures	56	Sedimentary tectonics	75
Ripple marks	56	Structure	76
Flute casts	57	Faults	76
Mudcracks	57	Joints	78
Miscellaneous	57	Folds	78
Paleontology	57	Tectonic history	78
Faunal assemblages	57	Economic geology	79

North Horn Formation	81	References cited	99
Flagstaff Formation	81		
References cited	81		
Figures			
1. Index map	65	1. Index map	84
2. Stratigraphic column	65	2. Conceptual model of compound faceted spur development	87
3. Hjork Creek diapir (Twist Gulch)	66	3. Spanish Fork Peak	90
4. North San Pitch collapsed diapir (high altitude vertical)	66	4. Selected facets on Spanish Fork Peak	90
5. Hjork Creek vertical outcrop	68	5. Area near Hobble Creek Canyon	90
6. Price River outcrop, overturned South Flat	69	6. Section of Spanish Fork Peak quadrangle topographic map	91
7. Formline contour map	69	7. Same map as figure 6 with facets and pediments identified	91
8. North Horn channels	71	8. Profile of rangefront	92
9. North Horn "giant" algal balls	71	9. Loafer Mountain area with key pediment correlations	92
10. North Horn algal ball	71		94
11. Pollen photomicrographs	73	10. Spanish Fork Peak area with key pediment correlations	94
12. Tertiary volcanic stream deposits	74	11. Hobble Creek area with key pediment correlations	94
13. North San Pitch collapsed diapir (low oblique)	75		94
14. Salt Valley cuesta	75	12. Maple Flat area with key pediment correlations	95
15. Green River cuesta	75	13. Mount Timpanogos area with key pediment correlations	96
16. Seven-step geologic evolution	76	14. Alpine area with key pediment correlations	96
17. Stream pattern analysis. Stippled areas denote areas of controlled drainage	77	15. Stranded streams at Corner Creek area	97
18. Rebound diagram	79	16. Generalized profile of range showing main correlations and uplift patterns	97
19. Geologic map with cross sections A-A', B-B', and C-C'	in pocket	17. Sketches of the Spanish Fork Peak area showing probable historical development	98
 COMPOUND FACETED SPURS AND RECURRENT MOVEMENT IN THE WASATCH FAULT ZONE, NORTH CENTRAL UTAH	83	 A SUBSURFACE CORRELATION OF PERMIAN-TRIASSIC STRATA IN LISBON VALLEY, UTAH	103
Introduction	83	Introduction	103
Acknowledgments	83	Acknowledgments	103
Previous studies	83	Previous work	103
Early work	83	Geologic setting	103
General geology	84	Stratigraphy	103
Wasatch fault geometry	85	Structure	107
Geophysics	86	Cutler stratigraphy	107
Origin of pediments	86	Sandstones	107
Conceptual model	86	Calcareous arkose	107
Pediment formation	86	Arkose wacke	107
Recurrent uplift	88	Mudstones	108
Pediment preservation	88	Sandy mudstone	108
Downcutting	88	Mudstone	108
Slope retreat	89	Sandy limestone	109
Structural control	89	Chinle stratigraphy	109
Sequential development	89	Sandstone	109
Methods	89	Calcareous arkose	109
Aerial photography	90	Arkose wacke	109
Pediment mapping	90	Mudstone	109
Profile projection	91	Sandy mudstone	109
Correlation	91	Mudstone	110
Results	93	Conglomerate	110
Payson Canyon to Spanish Fork Canyon	95	Mudstone conglomerate	110
Spanish Fork Canyon to Springville	95	Stratigraphic correlation	111
Springville to Provo Canyon	95	Appendix	113
Provo Canyon to American Fork Canyon	95	Measured section	113
American Fork Canyon to the Traverse Mountains	96	Core 3	114
Conclusions	97	Core 9	115
Pediment formation and slope retreat	97	References cited	115
Uplift and quiescence	97	Figures	
Absolute dating	99	1. Index map	104
Tilt	99	2. Fence diagram	105
Regional geodynamics	99	3. Map of Lisbon Valley	106
Suggested further studies	99		

4. Stratigraphic column	106	General statement	123
5. Cutler Formation calcareous arkose	107	Metamorphic rocks	125
6. Cutler Formation arkosic wacke	108	General statement	125
7. Mudstone	108	Metamorphic rock description	125
8. Cutler Formation sandy limestone	109	Marble and dolomarble	125
9. Diagram of cores	110	Phyllitic rocks	126
10. Chinle Formation calcareous arkose	111	Calc-silicate rocks	126
11. Chinle Formation arkosic wacke	111	Schist	126
12. Chinle Formation conglomerate	112	Quartzite	126
13. Cutler and Chinle formations	112	Metigneous rocks	126
Tables		Mineral assemblages	126
1. Logs and drilling data	104	Mafic assemblages	126
2. Core composition	106	Pelitic assemblages	126
3. Mineral assemblages	126	Calcareous and calc-silicate assemblages	127
MESOZOIC-CENOZOIC STRUCTURAL DEVELOPMENT OF THE KERN MOUNTAINS, EASTERN NEVADA-WESTERN UTAH	117	Contact vs. regional metamorphism	127
Introduction	117	Age of metamorphic rocks	127
Previous work	117	Igneous rocks	128
Acknowledgments	117	General statémant	128
Structures of the sedimentary cover	117	Volcanic rocks	128
General statement	117	Rhyodacite	128
Low-angle faults	118	Dacitic rocks	129
Regional denudation	119	Tuff	129
Local denudation	119	Summary and conclusions	129
High-angle faults	120	References cited	130
Folds	120	Figures	
Metamorphic structures	121	1. Index map	118
General statement	121	2. Structural outline	119
Skinner Canyon study	121	3. Geologic map	in pocket
Macroscopic structures	121	4. Sterographic projection of fold axes	121
Mesoscopic fabric	121	5. Sterographic projection of axial planes	121
Microscopic fabric	122	6. Sterographic projection of metamorphic foliation	121
Structure of the plutonic complex	122	7. Sterographic projection of granitic foliation	123
General statement	122	8. Sterographic projection of joints	123
Foliation	122	9. Sterogtaphic projection of aplitic dikes	123
Joints	123	10. Generalized stratigraphic column of regional stratigraphic and Kern Mountains stratigraphic column ..	124
Dikes	123	11. Generalized geologic map of Kern Mountains	125
Stratigraphy	123	12. Metamorphic rock correlation	129
Table		1. Kern Mountains geologic history	120

Additional Specimens of the Hypsilophodontid Dinosaur *Dryosaurus altus* from the Upper Jurassic of Western North America

JEFFREY D. SHEPHERD², PETER M. GALTON¹ AND JAMES A. JENSEN³

¹ Department of Biology, University of Bridgeport, Bridgeport, Connecticut 06602.

² Department of Biology, Westchester Community College, Valhalla, New York 10595.

³ Earth Science Museum, Brigham Young University, Provo, Utah 84602.

ABSTRACT.—Specimens of the hypsilophodontid dinosaur (Reptilia: Ornithischia: Ornithopoda) *Dryosaurus altus* (Marsh) are described from the Morrison Formation (Upper Jurassic) of Colorado, Utah, and Wyoming. As shown by these specimens, the postcranial anatomy of *Dryosaurus altus* (Marsh) is very similar to that of *Dryosaurus lettow-vorbecki* Pompeckj from the Tendaguru Formation (Upper Jurassic) of Tanzania, East Africa.

INTRODUCTION

The purpose of this paper is to describe four specimens (AMNH 834, CM 1949, CM 21786, DNM 1016)* of an ornithopod dinosaur (Reptilia: Ornithischia) from the Morrison Formation (Upper Jurassic) of Colorado, Utah, and Wyoming. Comparisons with YPM 1876 (holotype of *Laurasaurus altus* Marsh, 1878, the type species of genus *Dryosaurus* Marsh, 1894) show that these four specimens are referable to the hypsilophodontid *Dryosaurus altus* (Marsh). A summary of previous work on hypsilophodontid dinosaurs from the Upper Jurassic of North America is given by Galton and Jensen (1973:137-38).

Data on Specimens

AMNH 834.—From Bone Cabin Quarry, Medicine Bow anticline, Wyoming; specimen consists of one cervical and several fragmentary dorsal and caudal vertebrae; right scapula, coracoid, humerus, ilium, ischium and pubis, both femora, incomplete left tibia, pes and some phalanges (figs. 1A-D, 2A-D, 3F, 4A-C, 5A-B). Length of right femur 222 mm and original length of animal about 2.0 m.

CM 1949.—Collected by W. H. Utterback in 1905 from the Elk Mountains near Brown's Ranch, Johnson County, Wyoming; specimen consists of several dorsal and caudal vertebrae, right ilium, femur and tibia (figs. 3I, 4D-E, 5N-O). Length of tibia 496 mm and original length of animal about 3.8 m.

CM 21786.—Collected by J. L. Kay and A. C. Lloyd in 1955 from Lily Park, Moffat County, Colorado; specimen consists of 41 vertebrae, incomplete left ilium, ends of all long bones (left tibia with intact calcite core, length 356 mm, and original length of animal about 2.7 m) with both pes almost complete (figs. 1E-V, 2E-H, 3G-H, 4F-O, 5C-M).

DNM 1016.—Collected by J. Adams from the exhibit cliff at Dinosaur National Monument near Jensen, Uintah County, Utah (White 1964, fig. 1); left ilium (fig. 3A-E), length of original animal about 2.4 m.

ACKNOWLEDGMENTS

We thank the following people for all their assistance while studying specimens under their care: Dr. D. S. Berman, Carnegie Museum; Dr. E. S. Gaffney, American Mu-

seum of Natural History; Dr. J. H. Ostrom, Peabody Museum of Yale University; and Dr. T. E. White and Mr. J. Adams of Dinosaur National Monument. NSF grant DEB 76-09769 to P. M. Galton partly supported this work and provided \$250 toward the publication of this paper.

DESCRIPTION AND COMPARISONS

No attempt is made to describe the specimens in detail because there are several good descriptions of the bones of cursorial ornithopods in the literature. Instead comparisons have been drawn primarily with the following taxa: the fabrosaurid *Fabrosaurus australis* (Thulborn 1972) from the Upper Triassic of Lesotho, southern Africa; the hypsilophodontids *Hypsilophodon foxii* (Galton 1974) from the Lower Cretaceous of England; *Dysalotosaurus lettow-vorbecki* (Janensch 1955) from the Tendaguru Formation (Upper Jurassic) of Tanzania, East Africa, and *Othnielia* (Galton 1977, based on *Nanosaurus rex* Marsh, 1887; see Galton & Jensen 1973) from the Morrison Formation of western North America; and the iguanodontid *Camptosaurus dispar* (Gilmore 1909; the several Morrison species described are probably all referable to *C. dispar*) from western North America. Unless stated to the contrary, data for comparisons with these taxa are taken from the papers indicated above.

Vertebral Column

Using the vertebral column of *Hypsilophodon foxii* (Galton 1974, figs. 19-22) as a guide, the vertebrae of *Dryosaurus* represented by centra in CM 21786 (fig. 1E-V) are identified as shown. The two cervical centra (fig. 1E-H) are tentatively identified as cervicals four and seven, with the seventh showing a stouter medioventral keel than does the fourth (fig. 1F, H). A well-preserved cervical of *Dryosaurus* (fig. 1A-D) is tentatively identified as the ninth, and presumably it is the last of the cervical keels and, on the basis of *Hypsilophodon*, are identified as dorsals one and three. The remaining seven centra are tentatively identified as shown (fig. 1J-M). The dorsal centra of CM 21786 increase in width passing posteriorly to culminate in a heavy, squat centrum that is wider than long (fig. 1N, O), and clearly this is the last dorsal. The sacrum of *Dryosaurus* probably consisted of six vertebrae as in *Hypsilophodon*. The last four are preserved in CM 21786 (fig. 1P; first four vertebrae, unnatural hyperflexion between sacrals 3 and 4 not compensated for). In the caudal series (fig. 1P-V) the first chevron was borne between the centra of the first two caudals as indicated by the presence of ventral articular facets. Starting with about the twentieth caudal, the transverse process is represented by a small ridge on the side of the centrum (fig. 1S-V).

*Institutional names cited in this paper have been abbreviated as follows: AMNH, American Museum of Natural History, New York; CM, Carnegie Museum of Natural History, Pittsburgh; DNM, Dinosaur National Monument Quarry, Jensen, Utah; YPM, Peabody Museum, Yale University, New Haven.

Pectoral Girdle

The scapula and coracoid (fig. 2A) of *Dryosaurus* (AMNH 834) are not well preserved, but are typically hypsilophodontid in form. The ratio of the maximum length of humerus to that of scapula is 0.7 and, judging from the measurements given by Gilmore (1925), the corresponding ratio for another specimen of *Dryosaurus* (CM 3392) is about 0.8. In *Hypsilophodon*, *Othnielia*, *Dysalotosaurus*, *Thescelosaurus neglectus* (Gilmore 1915) and *Parksosaurus warreni* (Parks 1926, as *Thescelosaurus warreni*), this ratio is 1.0. The scapula of *Dysalotosaurus*, as figured by Janensch (1955), is slightly shorter than the humerus, but with the anteroventral process for the clavicle restored as in other hypsilophodontids, the scapula and humerus of *Dysalotosaurus* are of equal length.

Fore Limb

The deltopectoral crest of the humerus (fig. 2D, E) is low but stout as in *Othnielia*, *Dysalotosaurus*, *Campitosaurus*, and in young individuals of *Hypsilophodon*, rather than high and slender as in mature individuals of *Hypsilophodon*. The anterior aspect of the distal outer condyle of *Dryosaurus* (CM 21786) is flat with a groove on its median surface (fig. 2E, H). The form of this condyle is the same as in *Dysalotosaurus* (Janensch 1955, fig. 34d) and is in contrast to the well-rounded condyles of the humeri of *Othnielia*, *Hypsilophodon* and *Campitosaurus*. The distal end of the humerus of AMNH 834 is badly damaged, but the form of the condyles of YPM 1876, the holotype of *Dryosaurus altus*, are identical to those of CM 21786. Posteriorly in CM 21786 there is a well-defined intercondylar groove (fig. 2F, H) with a foramen at the end of the indented nonarticular surface (fig. 2F, H). Only the ends of the radius and ulna of CM 21786 are preserved and, although flattened, are similar to those of other hypsilophodontids.

Pelvic Girdle

The ilia of AMNH 834 (fig. 3F), CM 1949 (fig. 3I), DNM 1016 (fig. 3A-E), and the posterior portion of CM 21786 (fig. 3G, H) are hypsilophodontid in form. They differ from the ilia of *Hypsilophodon* (Galton 1969, 1974) in being proportionally low in lateral view (fig. 3B, F, G, I) with a much broader brevis shelf (fig. 3C, D). However, the brevis shelf of DNM 1016 (fig. 3A, D) and its preserved proximal portion in AMNH 834 (fig. 3F) are as in the specimens referred to *Dryosaurus* by Gilmore (1925, fig. 5) and to *Dysalotosaurus* by Janensch (1955, pl. 13, fig. 19). The brevis shelf is wide, thin, and at an acute angle to the main body of the ilium (fig. 3C) to give the V-shaped appearance noted by Gilmore (1925). Although damaged, the brevis shelf of CM 1949 was obviously thin and wider than the main body of the ilium. The form of the posterior part of the main body of the ilium is identical in the ilia of DNM 1016 (fig. 3A-E), AMNH 834 (fig. 3F), CM 1949 (fig. 3I), CM 21786 (fig. 3G), YPM 1884, and those of *Dysalotosaurus*. The posterior portion of the ilium becomes very broad and contrasts with the thinness of the brevis shelf. In this respect these specimens differ from *Hypsilophodon* and *Othnielia* in which the posterior end is much deeper, and the termination is gently rounded. The dorsal edges of AMNH 834 and DNM 1016 have a distinct lateral bevel (fig. 3B, F, G). Posteriorly this beveled surface becomes deeper to include most of the height of the ilium. This feature is also shown quite clearly in the ilia of *Dryosaurus* (YPM 1884 and Gilmore 1925, fig. 6) and *Dysalotosaurus* (Janensch 1955, fig. 37a). The posterior surface of the ilium of CM 21786 (fig. 3G, H) has a matt finish, and in life this part was covered in cartilage.

The ischium of *Dryosaurus* (fig. 3F) has a proximally placed obturator process, and posterior to it the two ischia are in contact for the rest of their length. The distal half of

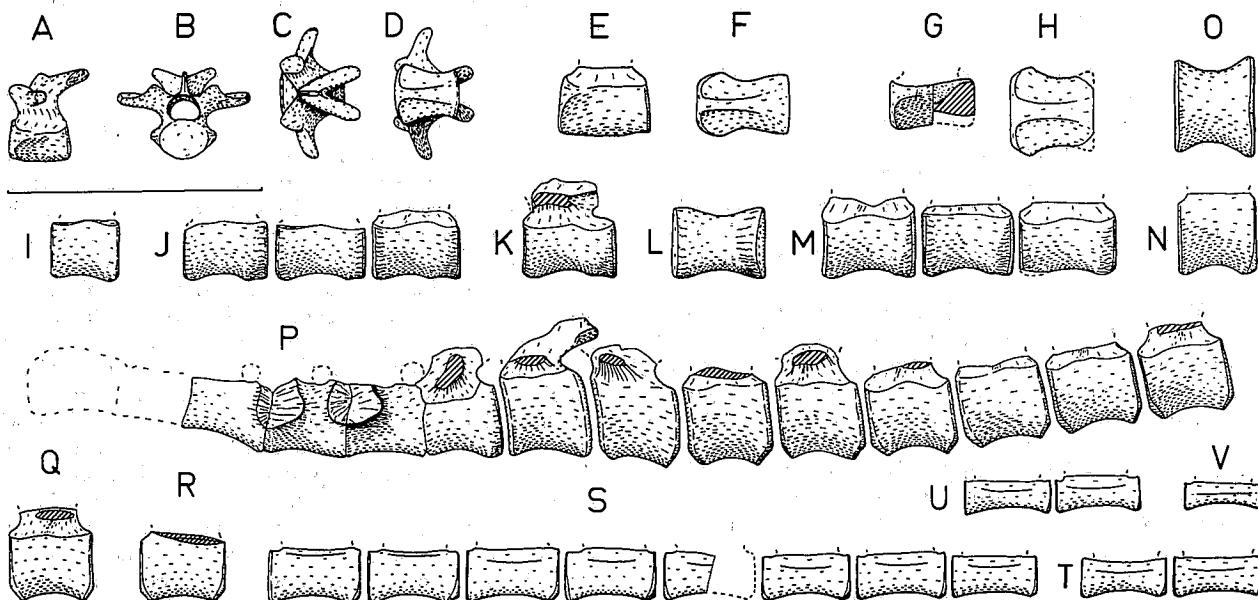


FIGURE 1.—*Dryosaurus altus*, referred specimens, vertebrae X ½. A-D.—Ninth cervical vertebra, AMNH 834, in A, lateral; B, anterior; C, dorsal; D, ventral views; E-V.—vertebral centra of CM 21786: E, cervical 4 in lateral view; F, as E in ventral view; G, cervical 7 in lateral view with posterior half sectioned along midline; H, cervical 7 in ventral view; I-O, dorsal vertebrae: I, dorsal 1 in lateral view; J, dorsals 3 to 5 in lateral view; K, dorsal 7 in lateral view; L, as K in ventral view; M, dorsals 10 to 12 in lateral view; N, last dorsal in lateral view; O, as N in ventral view; P, three sacral vertebrae conjoined with caudal vertebrae; Q-V, caudal vertebrae in lateral view, tentatively identified as follows: Q, 9; R, 15; S, 20 to 27; T, 31 and 32; U, 36 and 37; V, 42. Scale line represents 10 cm.

the ischium is bar-shaped (fig. 3F, sections) and gently curved in lateral view (fig. 3F). The ischia of *Dryosaurus* (Marsh 1896, pl. 55, fig. 4), *Dysalotosaurus* (Janensch 1955, pl. 13, figs. 21, 22) and, apart from the distal half being more curved with the end more expanded, the ischia of *Campitosaurus* (Gilmore 1909, 1912) are also very similar. In *Othnielia*, *Parksosaurus* (Parks 1926), and *Thescelosaurus* (Gilmore 1915) the obturator process is also proximally placed, but the distal half of the ischium is dorsoventrally flattened and blade-like. This is also the case in the ischia of *Hypsilophodon* in which the obturator process is at midlength (Galton 1969, 1974).

The preserved part of the right pubis (fig. 3F) of *Dryosaurus* is similar to the corresponding region of the pubis of YPM 1876 (Marsh 1896, pl. 55, fig. 4), *Othnielia*, *Dysalotosaurus*, and *Hypsilophodon*. In *Dryosaurus* (AMNH 834, YPM 1876) and *Othnielia* the prepubic process is long and in marked contrast to the short and sharply pointed prepubis, as figured by Marsh (1894, 1896) for *Laosaurus cossorus*, and to the very short prepubic process of *Fabrosaurus*.

Hind Limb

The femora of hypsilophodontids are the most diagnostic postcranial bones for taxonomic identification. The femora of CM 21786 (fig. 4J-O), AMNH 834 (fig. 4A-C), and CM 1949 (fig. 4D, 5O) are almost identical to that of YPM 1876 (Galton 1975, fig. 2G-L), the holotype of *Dryosaurus altus*. In particular all the femora show the following combination of characters: a deep cleft between the greater and lesser trochanters (fig. 4A, B, D, J), the fourth trochanter is on the

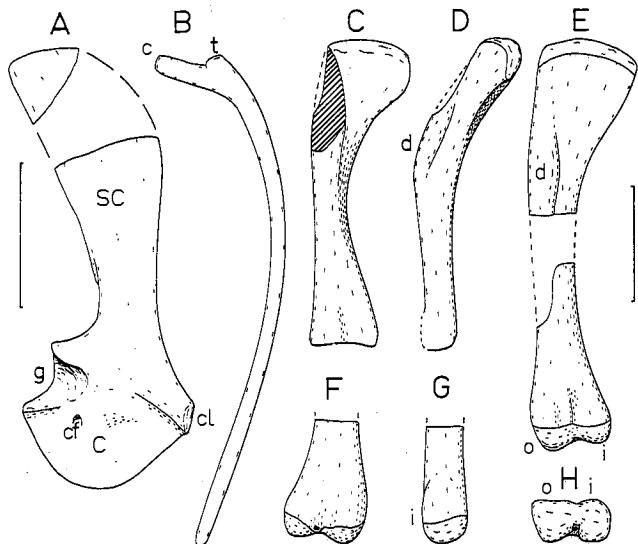


FIGURE 2.—*Dryosaurus altus*, pectoral girdle and forelimb, referred specimens. A, right scapula and coracoid in lateral view, AMNH 834 X0.38, dashed line based on impression in plaster jacket into which specimen fits. B, dorsal rib, AMNH 834, X0.38; C-D, right humerus, AMNH 834, X0.38 in C, anterior and D, medial views; E-H, right humerus CM 21786, X0.3 in E, anterior view; F, posterior view of distal end; G, medial view of distal end; H, distal end. c, capitulum; C, coracoid; cf, coracoid foramen; cl, facet for clavicle; d, deltopectoral crest; g, glenoid cavity; i, inner condyle; o, outer condyle; SC, scapula. Scale lines represent 5 cm.

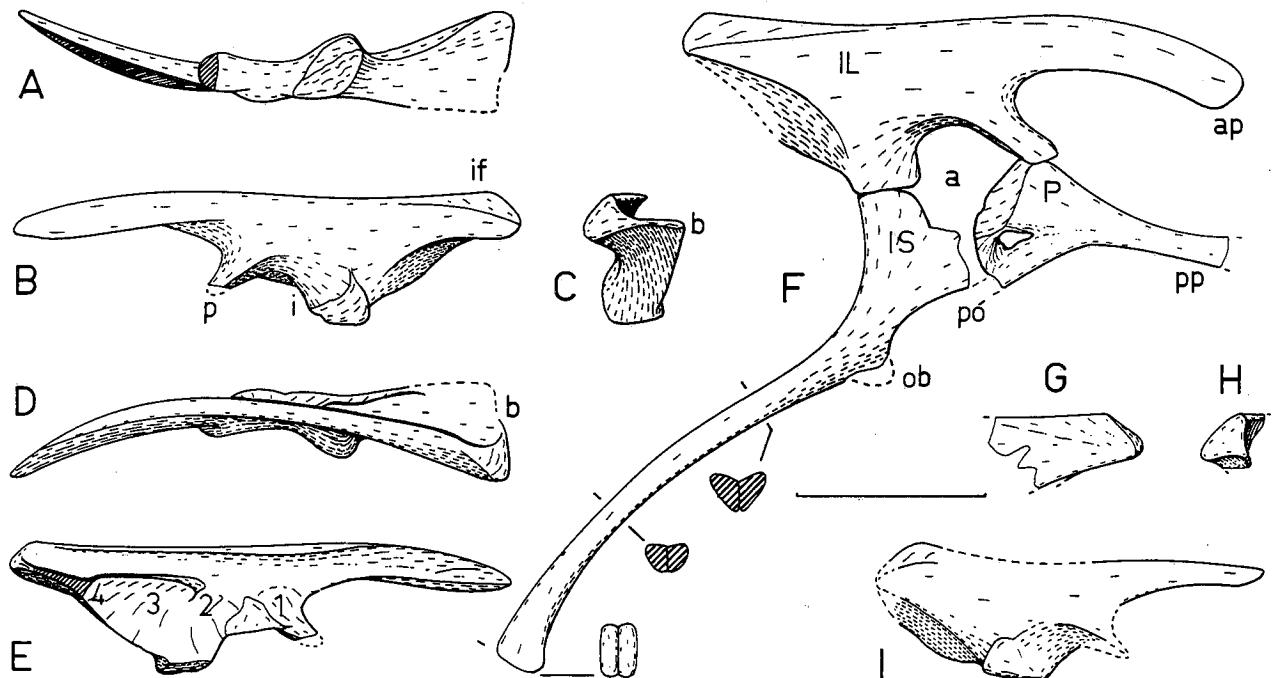


FIGURE 3.—*Dryosaurus altus*, referred specimens, pelvic girdle. A-E—Ilium DNM 1016, X0.25 in A, ventral; B, lateral; C, posterior; D, dorsal; E, medial views; F, pelvic girdle of AMNH 834 in lateral view, X0.5; G, lateral and H, posterior views of posterior portion of ilium, CM 21786, X0.25; I, ilium, CM 1949 in lateral view X0.125. Scale line represents 20 cm for I, 10 cm for A-E, G-H, and 5 cm for F-I. a, acetabulum; ap, anterior process; b, brevis shelf; i, ischiadic head; if, area for M. iliotibialis; IL, ilium; IS, ischium; P, pubis; po, postpubic rod; pp, prepubic rod or anterior process; 1, 2, 3, 4, attachment areas for sacral ribs 1 to 4 (area for 5 broken).

proximal half of the femur (fig. 4A, B, D), the deep depression for the *M. caudi-femoralis longus* (Galton 1969) is set well on the shaft (fig. 4A), the distal end is nearly square (fig. 4C, O) with a well-developed anterior intercondylar groove (fig. 4C, O, 5O), and posteriorly a proportionally small lateral condyle (fig. 4C, O). The femora of *Dysalotosaurus* (Janensch 1955, pl. 14, figs. 1, 2) are very similar. The femora of *Camptosaurus* have a much more anteroposteriorly expanded lesser trochanter and a more distally placed fourth trochanter. Distally the femora of *Othnielia*, *Hypsilophodon*, *Parksosaurus* (Parks 1926) and *Thescelosaurus* (Gilmore 1915) lack an anterior intercondylar groove.

The slender tibia of CM 1949 is longer than the femur (fig. 4D, E) as in all other hypsilophodontids. The form of the tibia (fig. 4E-I, 5B, N) of *Dryosaurus* is almost identical to that of YPM 1876, *Dysalotosaurus*, *Othnielia*, and *Hypsilophodon*. In *Camptosaurus* and *Thescelosaurus* the tibia is much

stockier and shorter than the femur. The fibula of CM 1949 is incomplete (fig. 4E), and it is of normal hypsilophodontid form. The fibulae of CM 21786 are represented by crushed ends that are not well preserved.

Only three metatarsals (fig. 5A) and some phalanges are represented in AMNH 834, but the pes of CM 21786 (fig. 5C-J) is the most complete pes of *Dryosaurus* discovered to date. CM 21786 possesses two incomplete right distal tarsals (fused 2-3, 4), the preserved parts of which are of the standard hypsilophodontid type (Galton 1974, figs. 57F, G). Although the shafts of the metatarsals are rather crushed, all the articular surfaces are well preserved (fig. 5H-J). The general form of the pes is shown (fig. 5C, D), and it is obviously slender. In *Hypsilophodon* (Galton 1974, fig. 58), *Parksosaurus* (Parks 1926, figs. 15, 16), and *Othnielia* (Galton and Jensen 1973, fig. 6D) the first metatarsal is relatively large. In *Dryosaurus* the pes of YPM 1884, AMNH 834, and CM 21786 (fig. 5C) all have associated phalanges, but there is no trace of a good-sized metatarsal I or of any phalanges referable to this digit. However, in the case of CM 21786, metatarsal I is probably represented by a diminutive element (fig. 5C, D, K-M). This element has a convex proximal articular surface, so it is not a phalanx; and it shows a well-defined distal articular surface with two separate condyles (fig. 5M). It cannot be a fifth metatarsal because where known, as in *Hypsilophodon* (Galton 1974, fig. 58B) and in *Othnielia* (Galton and Jensen 1973, fig. 6A), it is a slender, tapering rod with a rounded distal end without condyles. This small element of CM 21786 (fig. 5C, D, K-M) is very similar in form and size to an element preserved with the pes of *Dysalotosaurus* (Janensch 1955, fig. 40) and identified as metatarsal I. The pes of *Camptosaurus* and *Thescelosaurus* (Gilmore 1915) are much shorter and stockier than are those referred to *Dryosaurus*.

DISCUSSION

On the basis of the comparisons made above, it is apparent that the postcranial anatomy of *Dryosaurus altus* (Marsh) is most similar to that of *Dysalotosaurus lettow-vorbecki* Pompeckj, 1920, from the Upper Jurassic of Tanzania. As noted earlier (Galton 1973), a comparison of the skull of CM 3392 from the Morrison Formation of Utah with the holotype of *Dryosaurus altus* confirms the provisional identification by Gilmore (1925). However, the skull of CM 3392 does not have a supraorbital bar as described by Gilmore (1925, fig. 3). Instead, the supraorbital tapers gently to a point lateral to the postorbital with no suture between these bones. Apart from the longer supraorbital, the skull of CM 3392 (Galton 1977) is almost identical to that of *Dysalotosaurus* (Janensch 1955, fig. 1). In all other genera of hypsilophodontids no two skulls are alike.

In addition to general hypsilophodontid characters, *Dryosaurus* and *Dysalotosaurus* show the same combination of postcranial features. The humeri (fig. 2C-H) are very similar with a low delto-pectoral crest (also shown by *Othnielia*) and a uniquely flat lateral condyle, but in *Dryosaurus* the scapula is longer than the humerus whereas in *Dysalotosaurus* these bones are subequal in length. The ilia are very similar; the ilium of AMNH 834 is almost identical to that of *Dysalotosaurus* as figured by Janensch (1955). In particular both show a very low main body to the ilium, an obliquely truncated posterior end, and a broad brevis shelf (fig. 3A-C). In the ischia (fig. 3F) the obturator process is proximal in position (also shown in *Othnielia*, *Parksosaurus*), and the distal halves are bar-shaped and curved in lateral view. In the hind limb

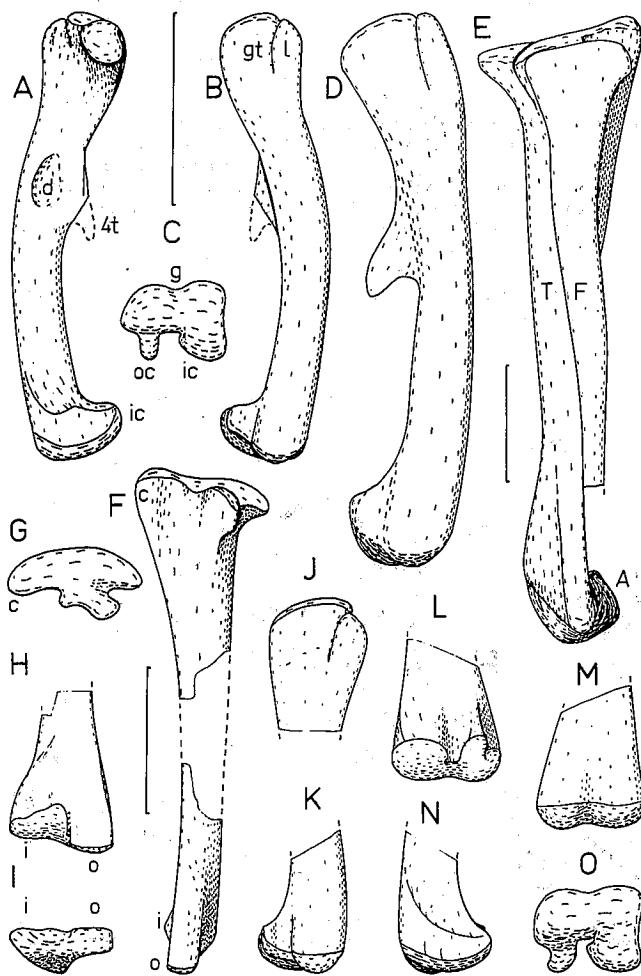


FIGURE 4.—*Dryosaurus altus*, referred specimens, femora and tibiae. A-C.—Right femur, CM 1949, in lateral view, X0.2. E.—right tibia, fibula, CM 1949, lateral view, X0.2. F-I.—Left tibia, CM 21786, X0.25 in F, medial view; G, proximal end; H, anterior view of distal end; I, distal end. J-O.—Right femur, CM 21786, X0.25; J, proximal end in lateral view; L, distal end in posterior view; M, distal end in anterior view; N, distal end in medial view; O, distal end. A, astragalus; c, cnemial crest; d, deep depression for insertion of *M. caudi-femoralis longus*; F, fibula; g, greater trochanter; i, inner malleolus; ic, inner condyle; o, outer malleolus; oc, outer condyle; 4t, fourth trochanter. Scale lines represent 10 cm.

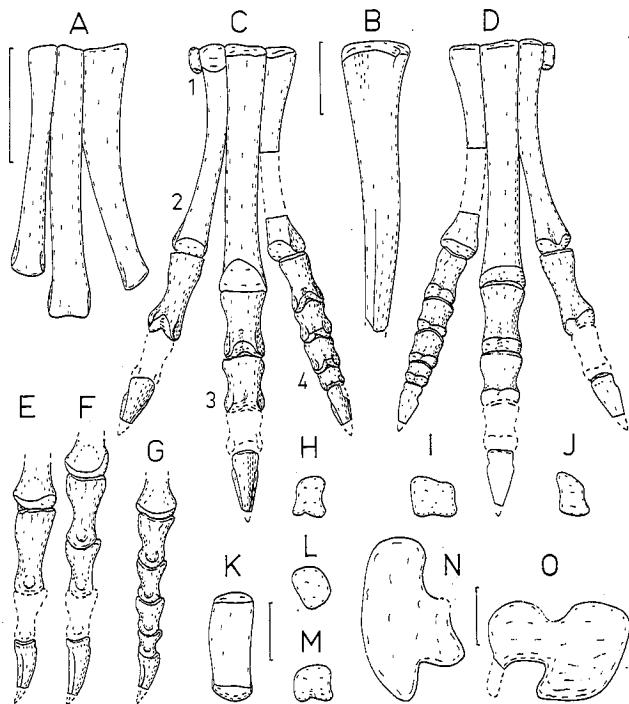


FIGURE 5.—*Dryosaurus altus*, referred specimens, pes. A, metatarsals 2-4, AMNH 834, dorsal view, X0.3. B, left tibia, AMNH 834, X0.3; C-M, left pes of CM 21786, X0.2; C, dorsal; D, ventral views; E-G, medial view of digits: E, 2; F, 3; G, 4; H-J, distal view of metatarsals: H, 2; I, 3; J, 4; K-M, metatarsal 1 in K, dorsal; L, proximal; M, distal views; N, proximal end of left tibia, CM 1949, X0.15; O, distal end of right femur, CM, 1949, X0.15. 1-4, digits 1 to 4. Scale lines represent 5 cm except K-M where 1 cm.

the femora of *Dryosaurus* (figs. 4A-D, J-O, 5O; Galton 1975, fig. 2G-L) and *Dysalotosaurus* are nearly identical and possess the following combination of characters: lesser trochanter is rodlike and separated from the greater trochanter by a deep cleft, the deep depression level with the fourth trochanter is set well anteriorly on the shaft, and the distal end is square with a well-developed anterior intercondylar groove. In the pes metatarsal I is rudimentary (figs. 5C, K) and is represented by a small nubbin of bone in both genera.

Dryosaurus altus and *Dysalotosaurus lettow-vorbecki* are very similar and, as discussed elsewhere (Galton 1977, in press), these species are congeneric and are referred to the genus *Dryosaurus* Marsh 1894 as *Dryosaurus altus* (Marsh 1878) and

Dryosaurus lettow-vorbecki (Pompeckj 1920). The occurrence of *Dryosaurus* in the Upper Jurassic of North America and East Africa provides good evidence of the presence of a land route between Laurasia and Gondwanaland in the early Upper Jurassic (Galton 1977, in press).

REFERENCES CITED

- Galton, P. M., 1969, The pelvic musculature of the dinosaur *Hypsilophodon* (Reptilia: Ornithischia): Postilla, no. 181, 64 p.
 —, 1973, Redescription of the skull and mandible of *Parkosaurus* from the Late Cretaceous with comments on the family *Hypsilophodontidae* (Ornithischia): Life Sci. Contr., Royal Ont. Mus., no. 89, 21 p.
 —, 1974, The ornithischian dinosaur *Hypsilophodon* from the Wealden of the Isle of Wight: Bull. Br. Mus. (Nat. Hist.) Geol., v. 25, p. 1-152.
 —, 1975, English hypsilophodontid dinosaurs (Reptilia: Ornithischia): Palaeontology, v. 18, no. 4, p. 741-52.
 —, 1977a, The Upper Jurassic ornithopod dinosaur *Dryosaurus*—evidence for a Laurasia-Gondwanaland connection: Jour. Paleont., v. 51 (2) III, p. 11-12.
 —, 1977b, The ornithopod dinosaur *Dryosaurus* and a Laurasia-Gondwanaland connection in the Upper Jurassic: Nature, v. 268, p. 230-32.
 —, in press, Upper Jurassic ornithopod dinosaur *Dryosaurus* and a Laurasia-Gondwanaland connection: Milwaukee Public Mus. Spec. Pap. Biol. Geol.
 Galton, P. M. and Jensen, J. A., 1973, Skeleton of a hypsilophodontid dinosaur *Nanosaurus* (?) rex from the Upper Jurassic of Utah: Brigham Young Univ. Geol. St., v. 20, no. 4, p. 137-57.
 Gilmore, C. W., 1909, Osteology of the Jurassic reptile *Campitosaurus* with a revision of the species of the genus, and description of two new specimens: Proc. U.S. Nat. Mus., v. 36, p. 197-333.
 —, 1912, The mounted skeletons of *Campitosaurus* in the United States National Museum: Proc. U.S. Nat. Mus., v. 41, p. 687-96.
 —, 1915, Osteology of *Thescelosaurus*, an ornithopodous dinosaur from the Lance Formation of Wyoming: Proc. U.S. Nat. Mus., v. 49, p. 591-616.
 —, 1925, Osteology of ornithopodous dinosaurs from the Dinosaur National Monument, Utah: Mem. Carnegie Mus., v. 10, no. 4, p. 385-409.
 Janensch, W., 1955, Der Ornithopode *Dysalotosaurus* der Tendaguru-schichten: Palaeontographica Suppl., v. 7(3), p. 105-76.
 Marsh, O. C. 1877, Notice of new dinosaurian reptiles from the Jurassic Formation: Amer. Jour. Sci. (3) v. 14, p. 514-16.
 —, 1878, Principal characters of American Jurassic dinosaurs: Amer. Jour. Sci. (3) v. 16, p. 411-16.
 —, 1894, The typical Ornithopoda of the American Jurassic: Amer. Jour. Sci. (3) v. 48, p. 85-90.
 —, 1896, The dinosaurs of North America: 16th Ann. Rpt. U.S. Geol. Surv., 1894-95, pt. 1, p. 133-244, pls. 2-85.
 Parks, W. A., 1926, *Thescelosaurus warreni*, a new species of ornithopodous dinosaur from the Edmonton Formation of Alberta: Univ. Toronto Geol. Ser., no. 21, 42 p.
 Pompeckj, J. F., 1920, Das angebliche Vorkommen und Wandern des Parietal foramens bei dinosauriern: Sber. Ges. Naturf. Freunde Berlin, v. 1920, p. 109-29.
 Thulborn, R. A., 1972, The postcranial skeleton of the Triassic ornithischian dinosaur *Fabrosaurus australis*: Palaeontology, v. 15, no. 1, p. 29-60.
 White, T. E., 1964, The dinosaur quarry: Intermountain Assoc. of Petrol. Geol. 13th Ann. Fld. Conf., v. 1964, p. 22-28.

