Vol. 5

No. 7

July, 1958

PALEOECOLOGY OF THE MANNING CANYON SHALE IN CENTRAL UTAH

by
Richard W. Moyle

Brigham Young University

Department of Geology

Provo, Utah

PALEOECOLOGY OF THE MANNING CANYON SHALE

IN

CENTRAL UTAH

A thesis

submitted to

the Faculty of the Department of Geology

Brigham Young University

In partial fulfillment
of the requirements for the degree

Master of Science

by

Richard W. Moyle

July 1958

CONTENTS

LIST OF ILLUSTRATIONS		• •		•		· ii
ACKNOWLEDGMENTS						iii
ABSTRACT	٠.					iv
INTRODUCTION						
Purpose and scope	• •		•	•		1
Locations	• •	•	•	•	•	1
Previous work	• •		•	•	•	1
Field work		٠	. 19		•	4
Laboratory work		٠	•		i .	5
2	•	•	•	•	٠	5
MISSISSIPPIAN SYSTEM						
Great Blue Limestone	•	٠	•	•	٠	7
	•	٠		•	*	7
MISSISSIPPIAN-PENNSYLVANIAN SYSTEMS						
Manning Canyon Shale	•	•	• •	٠	1300	8
Soldier Canyon type section		•	• •	a		8
Ophir Pass	•	•		•	3 4 €	9
Manning Canyon type locality	•					19
West Canyon	•	٥	• •	٠	*	21
Lake Mountain	•	٥		3 # 35	٠	23
Provo Canyon	•	*	• •	۰	•	26
Traverse Mountains	•	٠	• •		٠	27
Five Mile Pass	٠	ø	*	•	•	28
Stansbury Mountains	٥	•	(*)	•	3 .6 3	28
Stansbury Mountains	•	٥	• •	•	•	29
Onaqui Mountains	•	٠	• •	277.0		29
CONCLUSIONS						
	7.4	٠	• •	•	*	31
Correlation and age	•	•		٠	٠	31
Environment of sedimentation	•	•	• •		*	35
Cyclic sedimentation	٠			•		36
Source area	•					39
Paleoecology	٠	٠	• •	•	•	41
APPENDIX		•		٠	٠	48
SELECTED REFERENCES						12 121

LIST OF ILLUSTRATIONS

Plate 1	Vertical air photo showing location of Soldier Canyon measured section 10
Plate 2	Photographs of Manning Canyon Shale upper contact, medial limestone, and lower contact in Soldier Canyon, Oquirrh Mountains, Utah . 13
Plate 3	Type section of Manning Canyon Shale, Soldier Canyon, Oquirrh Mountains, Utah 15
Plate 4	Stratigraphic sections of Manning Canyon Shale in central Utah
Plate 5	Lithologic oscillation chart of Manning Canyon Shale in central Utah
Plate 6	Photographs of thin sections
Plate 7	Photographs of thin sections
Figure 1	Index map
Figure 2	Bed thickness chart 6
Figure 3	Cyclic patterns in Manning Canyon Shale, Oquirrh Mountains, Utah 40
Figure 4	Isopachous map of Manning Canyon Shale and correlative formations in central Utah and eastern Nevada
Figure 5	Fossil-sediment facies patterns in Manning Canyon Shale, Soldier Canyon, Oquirrh Mountains, Utah
Table 1	Color, particle size, and soluble material of the shales in Soldier Canyon, Oquirrh Mountains, Utah

ACKNOWLEDGMENTS

Dr. Harold J. Bissell, chairman of the thesis committee, and Dr. J. Keith Rigby gave assistance in the field and laboratory, and made valuable suggestions for the organization and completion of the manuscript. Doctors Kenneth C. Bullock, Lehi F. Hintze, Jess R. Bushman, and Mr. Willis H. Brimhall assisted in rock descriptions and made helpful comments on the thesis.

Messrs. Gilbert Greene and Lloyd Burkle, graduate geology students of Brigham Young University, assisted in measuring some sections in the field and identified some fossils. David Moyle assisted in measuring sections and collecting samples in the field.

To his wife Belva, the writer expresses gratitude for assistance in the preparation of the manuscript, as well as help and encouragement toward the completion of the work.

ABSTRACT

Five sections of the Upper Mississippian and Lower Pennsylvanian Manning Canyon Shale in central Utah were measured, and five additional partially exposed sections nearby were visited for comparison.

The sections range from 1000 to 2600 feet thick and average between 1500 and 1600 feet. Shale predominates throughout. Clay shales characterize the lower 600 feet of the formation and silty shales and shaly limestones typify the upper 900 to 1000 feet. Lake Mountain section is the exception in that quartzites, arkoses and subgraywackes make up about one-third of the formation. Fresh, brackish, and shallow warm marine water conditions are suggested in central Utah throughout Manning Canyon time as evidenced by typical marine fossil assemblages, pelecypod and gastropod faunas, and plant remains in the sediments.

The sponges Talpaspongia clavata, Wewokella contorta, Wewokella solida, and Wewokella n. sp. were found in Mississippian rocks for the first time. They are present in West Canyon, Soldier Canyon and Ophir Pass.

Correlation and study of lithology demonstrate the cyclic nature of the Manning Canyon seas and reflect three regressions and two transgressions punctuated with many minor climatic and depositional changes. The formation thickens and size of the clastics increases to the east. Conclusive evidence of the Mississippian-Pennsylvanian time boundary was not found; however, it is tentatively placed above a fossiliferous limestone zone about 575 feet above the base of the formation and below a plant zone within the interbedded shales and quartzites 700 feet above the base of the formation.

INTRODUCTION

Purpose and Scope

The Manning Canyon Shale at the type locality in Manning Canyon of the Oquirrh Mountains in central Utah, and in contiguous area, bridges the Mississippian-Pennsylvanian time boundary and in addition reflects considerable variation in depositional environments.

Various geologists have mapped, measured and discussed the formation in central Utah, but a detailed study of the Soldier Canyon type section in the Oquirrh Mountains has not been reported. Nor has anyone heretofore compared the relationship of fauna and flora to the sediments, the fluctuation in salinity and depth of water, or the fluctuation between marine, brackish and non-marine water within the Manning Canyon Shale of central Utah.

The writer attempts to present data concerning the systemic boundary, and to reconstruct the sedimentary environments by means of field and laboratory evaluation of the fossils and the sediments of the formation. Also, it is a definite objective to ascertain guide fossils and correlate lithologic units from area to area.

Only a selected few complete sections were studied in detail. Partly exposed sections were visited to see if correlation could be obtained.

Locations

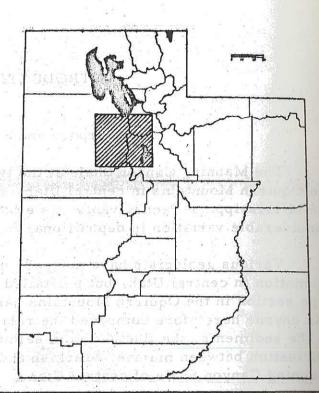
The Manning Canyon Shale crops out in many places in central Utah. Five complete sections were visited and studied along with five incompletely exposed sections (Fig. 1). Nine of the ten studied sections lie within the easternmost ranges of the Basin and Range Physiographic Province, and the remaining one lies in the westernmost part of the Wasatch Range in the Central Rocky Mountain Province.

The detailed complete sections include:

 Soldier Canyon; (Type Section) on the north side of Soldier Canyon, about one-half mile above its mouth, Sec. 33, T. 5 S., R. 4 W., Tooele County, Utah, (BYU Loc. Nos. 11018, 11019, 11020, 11021).

Index Map

- 1 Soldier Canyon
- 2 Ophir Pass
- 3 Manning Canyon
- 4 West Canyon
- 5 Lake Mountain
- 6 Provo Canyon
- 7 Traverse Mountains
- 8 Five Mile Pass
- 9 Stansbury Mountains
- 10 Onaqui Mountains



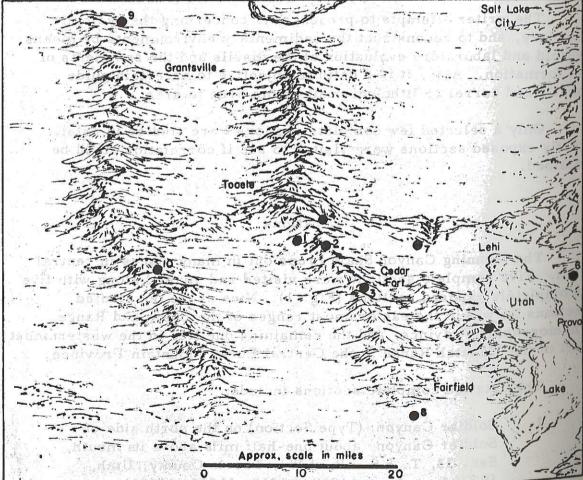


Figure 1

- 2. Ophir Pass; on the divide between the head of the South Fork of Soldier Canyon and the head of Ophir Canyon, Sec. 11, T. 6 S., R. 4 W., Tooele County, Utah, (BYU Loc. No. 11022).
- 3. Manning Canyon; (Type Locality) in Manning Canyon at the head and about one-half mile below the head, in Sec. 4, T. 6 S., R. 3 W., Utah County, Utah, (BYU Loc. Nos. 11023, 11024).
- West Canyon; on the north side about one to one and one-fourth miles west up canyon from the mouth of Iron Canyon, Sec. 3, T. 5 S., R. 3 W., Utah County, Utah, (BYU Loc. Nos. 11005, 11025, 11026).
- 5. Lake Mountain; on west side of Seep Canyon north and west of spring, east side of range, Sec. 36, T. 6 S., R. 1 W., Utah County, Utah, (BYU Loc. No. 11027).

The incomplete sections observed:

- 6. Provo Canyon; at the head of Little Rock Canyon just east of the road, Secs. 16 and 21, T. 6 S., R. 3 E., north side of creek at Canyon Glen Camp Grounds, Secs. 4 and 5, T. 6 S., R. 3 E., and above cliffs north of the Provo City Power Plant in the mouth of Provo Canyon one-fourth S. E., Sec. 1, T. 6 S., R. 2 E., Utah County, Utah.
- 7. Traverse Range; measured north-south across the clay pits in center of Sec. 9, T. 5 S., R. 1 W., Utah County, Utah.
- 8. Five Mile Pass; on south slope of east-west trending spur, S. E. one-fourth, Sec. 2, T. 8 S., R. 3 W., Utah County, Utah.
- 9. Stansbury Mountain; north end of Stansbury Mountains east side in clay pit, center of Sec. 8, T. 2 S., R. 6 W., Tooele County, Utah.
- Onaqui Mountains; north end of Onaqui Mountains north and south of the road about one-fourth to one-half mile east of Johnson Pass, Sec. 31, T. 5 S., R. 6 W.,
 Tooele County, Utah, (BYU Loc. No. 11028).

Previous Work

Gilluly (1932, p. 31) named the Manning Canyon Shale where he observed outcrops of the formation in Manning Canyon in the southeastern part of the Oquirrh Mountains. He measured the type section of the formation in Soldier Canyon to the north and west of the Manning Canyon type locality.

Nolan (1935, pp. 32-33) described a section of the Manning Canyon Shale in the Gold Hill district of Nevada and correlated the section with that of the Oquirrh Mountains. Bissell (1935, p. 163) suggested from a fossil study in the Wasatch Mountains that the systemic boundary occurred in a ten foot limy-sandstone within the formation.

Robertson (1940, p. 70) studied the selenium percentage and content in various shale samples from the Manning Canyon Shale in Provo Canyon and Manning Canyon, Utah. Baker (1947) measured a thick section of the formation near the mouth of Provo Canyon, Utah. Bullock (1949) mapped and measured a section of the formation in the north end of the Pelican Hills on the east side of Lake Mountain and also another section on the west side of Lake Mountain.

Hebertson (1950, p. 31) attempted a correlation of the facies within the Manning Canyon Shale by means of physical and chemical sedimentary characteristics. He investigated the formation in: (1) Manning Canyon in the Oquirrh Mountains; (2) Lake Mountain, northeastern end; and (3) Provo Canyon in the Wasatch Mountains.

Calderwood (1951, pp 63-65) mapped and measured a section in the Cedar Valley Hills on the west side of Lake Mountain. Ornelas (1953) attempted to classify and evaluate the clay possibilities of the Manning Canyon Shale throughout Utah County and the adjacent areas.

McFarland (1955, pp. 11-12) mapped and measured a complete section in West Canyon in the Oquirrh Mountains. Sadlick (1955, p. 3) studied sections of the formation in: (1) Uinta Mountains, north and south flanks; (2) Wasatch Mountains, at Ogden, Salt Lake, and Provo, Utah, areas; and (3) Pisgah Hills, southwest of Logan, Utah. Young (1955) mapped an incomplete section of the formation in the southern Lakeside Mountains, Utah.

Croft (1956) mapped and measured a section in the northern Onaqui Mountains, but due to faulting in the area an accurate stratigraphic measurement could not be made. Hyatt (1956) using Ornelas' 1953 thesis as a nucleus made an additional classification and evaluation of the clay

deposits within the formation. Pitcher (1957, p. 8) mapped an incomplete section in the south and east end of the Traverse Mountains, Utah.

Field Work

Dr. Harold J. Bissell suggested the problem in the spring of 1952. The field work commenced in April of the same spring, but an interruption in August, 1952, delayed the work until the spring of 1957 when again it was commenced, being completed in June, 1958. The work consisted of detailed measuring with steel tape and sampling each rock unit in the selected stratigraphic sections.

inge if is exposed in Iron Canyon the south to just north of he town

Laboratory work included preparation and study of thin sections of the limestone and quartzites, insoluble residues of the shales, and the preparation of fossils for identification.

The writer prepared thin sections using standard procedures, and studied them with binocular and petrographic microscopes. Insoluble residues were prepared by dissolving 10 gram samples of the shales broken into pea size, in one part hydrochloric acid to one part water. These were washed, dried, weighed, and the residues determined in per cent of insoluble material.

Fossil preparation consisted of placing silicified fossiliferous limestones in one part phosphoric acid and five parts water for a controlled period of observed digestion, then carefully washing and drying the fossils. Plastic spray preserved the more fragile fossils, especially the bryozoans. Subjecting fossiliferous limestone and shale samples to heat in an electric furnace until red hot, then immersing them in cold water, caused some of the fossils to loosen from the matrix due to slaking of lime at the contact of fossil and matrix.

The color description of fresh and weathered rock surfaces conforms to that of the National Research Council rock color chart, with the exception that the color olive was excluded and different combinations of the colors green, yellow, and brown were used.

Wentworth's scale was used for grain size analysis.

A combination of Kelley (1956), and McKee and Weir (1953) terminology for bed thickness and weathering structure was used as given in the following chart:

Name	Thickness	Weathering Structu
Massive	Greater than 6 ft.	Massive
Thick	3 to 6 ft.	Blocky
Medium	l to 3 ft.	Slabby
Thin	l in. to l ft.	Flaggy
Laminae	2 mm. to 1 in.	Platy (or Shaly)
Thin Laminae	Less than 2 mm.	Papery

Fig. 2 Bed Thickness Chart

ory work included preparation and study of thin sortions is made, and quartzites, insoluble residues of the shales, and in of fessils for identification.

gr prepared thin sections using standard procedures, and standed here the binocular and petrographic microscopes. Insoluble testores were prepared by dissolving 10 gram samples of the shales to the chales and into the same and to one part waters the shales were shed, in one part hydrochloric acid to one part waters.

rparation consisted of placing silicified tossiliferous many phosphoric acid and five parts water, for each picture of the parts water, for each picture of the more fragile fossils.

Yespans, Subjecting fossiliferous limestone and shale interpretation for the fossils of the source of the fossils to leasen from the caused some of the fossils to leasen from the contact of fossil and matrix.

escription of fresh and weathered rock surfaces of the National Research Council rock color chirt, that the color of was excluded and different colors green, veilow, and brown were used.

reale was used for grain size analysis.

MISSISSIPPIAN SYSTEM

Great Blue Limestone

Throughout central Utah the upper member of the Great Blue Limestone crops out as a resistant ledge or cliff that can be very easily recognized from the overlying subdued slope forming Manning Canyon Shale.

The formation in the Oquirrh Mountains is exposed on the west flank of the range from Soldier Canyon on the north and west, south through Ophir Pass into the west side of Manning Canyon. On the east flank of the range it is exposed in Iron Canyon then south to just north of the town of Cedar Fort, where it is covered by alluvium.

In Lake Mountain the Great Blue Limestone is a ridge former, especially on the east side. At the mouth of Provo Canyon the ridge forming nature is again evident. This feature of the formation makes the upper contact with the basal overlying Manning Canyon Shale a recognizable stratal horizon. The complete sections in this report were measured from this contact.

The top of the upper member in the Great Blue Limestone in all areas that the present writer observed consists of medium and dark blue-gray limestone that weathers light gray, is very fine grained, and is thin to medium bedded in beds from one-half to twenty-four inches thick which weathers platy. This lithology is consistent throughout the area of study with only slight variations. In West Canyon and Soldier Canyon the weathered surface has a pink coloration. On Lake Mountain the upper member of the formation becomes cherty at the contact. In Manning Canyon, Chonetes and Archimedes are found along with other fossil hash very near the contact.

The Great Blue Limestone is Upper Mississippian (Girty In Gilluly, 1933, p. 30) in age.

Mountain '(top unexposed), 1910 feet, Average thickness for mention is 1500 to 1500 feet, The Saldier Canyon section is an nearly complete, is the best exposed, and represents the format best. A correlation chart (Plate A) shows the relationship of this and lithology for the formation in central Utah.

MISSISSIPPIAN PENNSYLVANIAN SYSTEMS

Manning Canyon Shale

Gilluly (1933, p. 31) named the formation as a distinct and separate lithology. Previously Spurr (1895, p. 376) called it the "upper shale" of the Great Blue Limestone.

Gilluly named the formation from exposures in Manning Canyon, but further states, "The section measured in Soldier Canyon is believed to represent the formation better than any of the others." In the present report the writer refers to the Manning Canyon section as the "type locality" and the Soldier Canyon section as the "type section". Most of the following discussion in the report centers around observations taken in Soldier Canyon where detailed measurements and descriptions were made.

Throughout the area of study the Manning Canyon Shale, at most localities, is a valley-former. The soft and impervious nature of the shales within the formation gives rise to fertile soil covered slopes and valleys. Springs occur in the formation in West Canyon, near Ophir Pass and in Manning Canyon. Vegetation and alluvial cover makes fresh exposures scarce.

The formation consists predominantly of shale, interbedded limestone, and quartzite, with few beds of siltstone and mudstone. Arkose is present locally. The Lake Mountain locality is the only section where shales do not predominate; there quartzite, subgraywacke, and arkose are abundant and make up over one-third of the formation.

The age of the formation is uppermost Mississippian and lowermost Pennsylvanian. Previous geologic work through fossil relationships has indicated its age most likely to be Upper Chester and Lower Springer.

Thicknesses of five measured sections are as follows: Soldier Canyon, 1559 feet; Ophir Pass, 1327 feet; Manning Canyon (base faulted), 1064 feet; West Canyon (top unexposed), 1683 feet; and Lake Mountain (top unexposed), 1910 feet. Average thickness for the formation is 1500 to 1600 feet. The Soldier Canyon section is most nearly complete, is the best exposed, and represents the formation best. A correlation chart (Plate 4) shows the relationship of thickness and lithology for the formation in central Utah.

Soldier Canyon (Type section). The formation crops out on the north side of Soldier Canyon (Plate 1) which is on the north nose of the plunging Ophir Anticline approximately three to three and one-half miles east and south of Stockton, Utah. Soldier Creek has eroded across the nose of the anticline and exposed the section. It crops out at the mouth of Soldier Canyon then trends east to one-quarter to one-half mile west of the mouth of South Fork where it passes under the valley floor and is exposed intermittently to the south in South Fork Canyon on to Ophir Pass.

Of all the localities visited by the writer the exposure in Soldier Canyon presents the formation more nearly undisturbed and most completely exposed. The section was measured on the north side of the canyon about one-half mile above its mouth, Sec. 33, T. 5 S., R. 4 W., Tooele County, Utah. The writer measured the section in four segments because of local soil cover (Plate 1). The formation was divided into 32 units and sub-divided into over 120 sub-units. Three additional units in the basal part of the Oquirrh Formation were closely observed and described to familiarize the writer with the upper contact of the Manning Canyon Shale.

The Manning Canyon-Great Blue contact is taken where the platy dense dark blue-gray limestones of the upper Great Blue Limestone change to the medium gray calcareous shales in the base of the Manning Canyon Shale. The Manning Canyon-Oquirrh contact was taken at the top of the uppermost dark gray shale and the base of a sandy medium brown-gray, brown weathering limestone and calcareous brown quartzite unit.

The basal five to ten feet of the Oquirrh formation contains abundant fossils, especially the brachiopod Spirifer. Immediately above this basal Oquirrh unit the shaly limestones weather so as to appear as shale on some spurs in Soldier Canyon. Care should be taken by stratigraphers not to confuse this overlying 92 foot zone with the uppermost shale in the Manning Canyon Shale. Within this shaly limestone about fifty to seventy-five feet above the base occur large Derbyia, abundant Echinoconchus, Allorisma?, various other productids, and abundant crinoid stems which are characteristic of the Oquirrh Formation and are diagnostic enough to differentiate these two similarly appearing units.

The lithologic composition of the Manning Canyon Shale in Soldier Canyon consists of approximately the following:

Rock type beruses	Total thickness in feet	Percentage of formation
Silty Shale	560	35.9
Clay Shale	391	25.1
Total Shale	951	61.0

Plate 1 Vertical air photo (CYO-12K-107, 1952) showing location of Soldier Canyon measured section.

Shaly Limestone Pure Limestone Sandy Limestone Total Limes	308 220 10	538	19.7 14.2 0.6
Quartzite		42	2.7
Sandy Shale, Silts Mudstone, and		_28	1.8
Total	thickness	1559	100.0

- 11-

The shaly limestones are distributed throughout the entire formation while the quartzite occur in two zones, one about 600 to 800 feet above the formation base and the other about 1300 to 1350 feet above the base. The shales predominate in the basal 900 feet, but persist throughout. Also clay shales are more predominant in the basal 600 to 700 feet then silty shales become the predominant type in the top 900 to 1000 feet.

In the laboratory the shales were checked for the amount of calcium carbonate. Table 1 shows the relation of color, silt to clay, and the percentage of calcium carbonate.

The brown and red-brown shales contain the largest percentages of calcium carbonate, the light and medium gray and brown-gray shales contain lesser amounts, and the dark gray and black shales contain the least.

The red-gray, red-brown, and brown shales are very calcareous (40 to 80 per cent calcium carbonate). The brown-gray, light gray, and medium gray shales are calcareous (10 to 40 per cent). The dark gray and black shales are slightly to very slightly calcareous (less than 10 per cent). Only a few of the black shales are calcareous; however, most are very carbonaceous and may in some areas be considered petroleum source rocks.

The silty shales are slightly more calcareous than the clay shales. As the color darkens to black, due to the increase in carbonaceous content in many beds, the clay content increases and the calcareous content decreases, suggesting change in depositional environment and provenance.

Of special interest in the basal half of the formation the pure limestones, fossiliferous limestones, and shaly limestones all have a one to six foot brown limestone "cap". These "caps" are evident in

Explanation of Plate 2

- Figure 1 Manning Canyon-Oquirrh contact and upper part of the Manning Canyon Shale, Soldier Canyon, Oquirrh Mountains, Utah. Viewed toward the west.
- Figure 2 Medial lime stone of Manning Canyon Shale, Soldier Canyon, Oquirrh Mountains, Utah. Viewed toward the west.
- Figure 3 Great Blue-Manning Canyon contact and lower part of the Manning Canyon Shale, Soldier Canyon, Oquirrh Mountains, Utah, Viewed toward the west

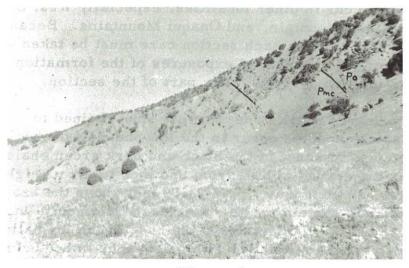


Figure 1



Figure 2

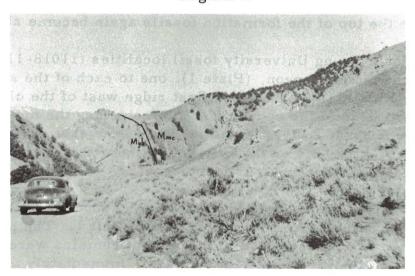


Figure 3

Soldier Canyon and occur in other localities, especially West Canyon, Manning Canyon, Lake Mountain, and Onaqui Mountains. Because of the numerous brown "caps" in each section care must be taken when correlation is attempted. In partial exposures of the formation the brown "caps" are indicative of the lower part of the section.

Siltstone, calcareous siltstone and very fine-grained to lithographic limestone nodules occur in the low one-half of the formation being more abundant in the black and dark green shales. In Soldier Canyon they range in size from one inch to six or eight inches in diameter. Within unit 7, 275 - 300 feet above the base of the formation, the nodules contain pyrite and iridescent limonite in the center. In unit 16, 640 - 700 feet above the base of the formation they are very dense calcareous siltstones with pelecypods imbedded in the outer edge.

Hydrothermal alteration is evident throughout the formation, but of minor importance. Many limestones contain secondary calcite and quartz stringers and vein fillings. Also, in many of the shales alteration takes place breaking down the shales with the addition of calcite and quartz. About 65 feet above the base of the formation this is very evident.

In the base of unit 7 about 264 feet above the base of the formation a four to six inch bed of gypsum, which appears to be a primary deposit, is present in Soldier Canyon.

The fossils found in the formation are dominantly marine invertebrates with a few plant remains, possibly terrigenous derived.

The lower 800 feet of the formation contains most of the fauna and flora. Near the top of the formation fossils again become abundant.

Four Brigham Young University fossil localities (11018-11021) are designated in Soldier Canyon, (Plate 1), one to each of the segments measured. The first, 11018, is on the first ridge west of the old mine tunnel on the north side of the canyon just after entering the mouth of the canyon. The second, 11019, is on the fourth spur east of the mine tunnel. The third, 11020, is on the fifth spur east of the mine tunnel, and the fourth, 11021, is on the fourth spur east of the mine tunnel.

Locality 11018 contains Yoldia, another pelecypod genus, and plant fragments in the top fifteen to twenty feet of unit 1. The pelecypods show no evidence of being reworked, but the plant remains were so fragmented that no identification could be made. Immediately above

Shale, dk. gv. to blk. wthrs, med, to lt. gy. calc., silty, contains interbedded med, gy, very fossiliferous lms, that wthrs, lt. org. -bwn, and lt. top. Unit contains Spirifer, Dictyoclastus, Cleiothyridina, Fenestella, Polypora, Archimedes 2 spp., Ptylopora and crinoids. lamestone, dk. gy., wthrs. lt. gy. throughout except med. bwn, in top 2 ft., very fine grained, thin bedded, platy and shalv, contains shale partings which become very numerous in basal 30 ft., top 6 ft. contains Acanthopecten, Spirifer, and ostracods (?). 7 7 7 7 7 Shale, blk. to dk. bwn. and med. gy., wthrs. blk and dk. bwn. -gy. calc. and silty in top one-half, clavey non-calc, in base. Very carbonaceous and graphitic, chippy, contains nodules with pyrite in center, fossils consist of Spirifer, Fenestella, Nucula, Amplexizaphrentis, gastropods, and crinoids (1 - 1 1/2 in. in diameter). Limestone, dk. gy., wthrs. lt. gy. throughout with med. org.-bwn. 3 ft. SHAL bedding, platy and shaly, abundant white and yell, calcite stringers in top 3 ft. Shale, lt. to med. red-bwn. and lt. bwn. wthrs lt. bwn.-gy. and red-bwn. silty and calc. in top 10 ft., friable, calc. and clayey in basal 13 ft., contains interbedded med. gv. fine grained, yell. -bwn. wthrg. lms. () Limestone, med. gy., wthrs. lt. pink-NNING gy., very fine grained, thin and laminated bedding, platy and shaly, ridge former, contains shale partings especially in basal 5 ft. Σ Shale, med. bwn.-gy, and med. bwn. becoming blk. in basal 30 ft., wthrs. _____ med, bwn, and med, gy., calc, and silty, papery and splintery in top 15 ft., becomes calc. and clayey in middle 40 ft. of unit, carbonaceous, non-calc, and slightly silty in basal 30 to 50 ft., contains interbedded med, bwn, and med, gy, fine grained med. to lt. bwn. wthrg. Ims. in beds 1-2 ft. thick that contain white calcite Limestone, med. bl. -gy., wthrs. med. and It, bwn. -gy., very fine grained, med, bedded, slabby, contains cherty beds in basal 6 ft., fossil hash in basal 3 ft. Shale, lt. pink-gy, in basal 5 to 10 ft., becoming med. gy. in middle and blk. in top, wthrs, med. and lt. gy., clayey and calc., very carbonaceous in top, contains two pelecypod genera and plant fragments in top 15 to 20 ft.

Great Blue Limestone

NO

CA

9

Richard W. Moyle 50 40 Shale, blk. med. gy.-bwn. med. 30 grn.-gy., wthrs, lt. yell.-bwn, lt. gr.-bwn. and lt. gy., calc. slightly V V V V silty, some fissile, contains 20 interbedded med, to It, gy., med. grained, bwn. wthrg qtzt., and med. 10 gy. -bwn., cryst. bwn. wthrg., lms. Quartzite, lt. gy., wthrs. med. grn., fine grained, massive, rough and Limestone, med. to dk. gy., wthrs. pock-like surface, crossbedded, salt med, bwn, -gy., very fine grained, and pepper appearance, white quartz thin and laminated bedding, platy and and well, calcite stringers. shaly, some beds are shaly and very carbonaceous, few calcite stringers, interbedded med. gy. shale. Shale, blk., med. bwn., med. bwn.-red. med. gy., wthrs, dk. gy., lt. pink bwn., lt. pink gy., lt. bwn.-gy., fissile, silty in top 12 ft .. clavey in rest of unit, contains a few interbedded dk. gy. and blk. fine grained, med. bwn. wthre, I ft. beds of lms., also a few beds of lt. to med. gy. fine grained, med. bwn. wthrg., qtzt. Some shales Igneous sill contain discontinuous beds. blobs. and nodules of blk. calc. siltstone, also in middle of unit pelecypods, gastropods, Cordaites ? and other Shale, dk. to med. gy., wthrs. med. gy. and med. to lt. bwn. calc. and silty, conts. 1-6 ft. beds of interbedded blk. to med. gy. fine grained SHALE lms. that wthrs. dk. to med. gy. ANYON Quartzite, lt. gray, wthrs. med. bwn., O med, grained, crossbedded, scintillating, rough surface, contains MANNING Limestone, blk. to dk. gy., wthrs. lt. a blk, clay shale in middle of unit. bl. -gy., very fine grained to litho., med. to thin bedded in top 27 ft., laminated bedding in basal 3 ft. slabby Shale, blk. wthrs. dk. gy., clayey. and flaggy, contains 8 ft. bed of blk. non-calc., contains pelecypods in carbonaceous clay shale. Shale, med. red-bwn., wthrs. lt. red-Limestone, med. gy. and med. to dk. gy., calc., clayey in top, med. gy. gy.-bwn. wthrs. lt. bwn. and gy.-bwn., fine grained, thin bedded, platy and calc, in basal 15 ft., contains med. gy, fine grained, med. bwn, wthrg, Ims. in top 5 ft. of unit, also a few 23 shaly, middle 11 ft. is med. gy. to blk. shale that is very fissile, calc. clayey, similar 1 ft. beds of lms, interbedded. contains Chonetes in top 2 ft. Shale, med, grn., lt. gv., grn, and yell.-gy., calc., clayey, contains med. grn., wthrs. dk. bwn. and med. org. -bwn, fine grained lms, in top Shale, med, ev. and med, red-bwn and in base, also med. gy.-bl. shaly 12 wthrs. med. gy. and med. pink-bwn., I = I - I - Ilms, in middle of unit. The lms, contain calc., friable, fissile, slightly silty. Spirifer, Nuculana, Dictyoclostus, 22 -1-1-1-15 contains 5 ft. zone of dk. gy. shaly Archimedes, Fenestella, horn corals, lms. in middle, fine grained laminated trilobites, crinoids, and gastropods. bedding, contains Chonetes. Shale, med. gy. to blk., wthrs. dk. gy., Limestone, dk. gy-bl. wthrs. lt. gy-bl., fine grained, thin to very carbonaceous, clayey, non-calc., 21 chippy. thick bedded, slabby to blocky, ---argillaceous in part, ledge Limestone, (medial lms.) med. gv.-bl., wthrs. med. to lt. gy., med. bwn. in 10 top 6 to 10 ft., fine grained, med. Shale, blk, in base, med, red-bwn, and bedded, calcite crusts on surface in bwn.-gy. in top 3/4 of unit, wthrs. 20 middle, contains Fenestella, Polypora, lt. bwn.-gy. and red-gy., calc., silty. Spirifer, Marginifera, Composita, Cleiothyridina, Kaskia, Griffithides, Limestone, med. to dk. gy., wthrs. Talpaspongia, Amplexizaphrentis, crinoids, obiculod brachiopods, and med. gy., fine cryst., laminated and thin bedded, platy and flaggy, 19 _____ rough pock-like surface, calcite ---stringers. 18

—Type Section of Manning Canyon Shale—Soldier Canyon, Oquirrh Mountains, Tooele County, Utah.

Vertical scale

Lamestone, blk, to dk, gv., wthrs, med. to lt. bwn.-gy., thin bedded, platy and chippy, contains shales, becomes shale in basal 32 ft., Spirifer, Echinoconchus, Derbya, Allorisma crinoids M OR Limestone, med. gy., wthrs. lt. gy. very fine grained, laminated bedding. platy, crinoids in top 6 ft., in basal I QUIRRI Limestone, med. by., wthrs. med. to lt. bwn.-gy. also bwn., fine grained thin bed. in beds 1/4 - 12 in. thick, platy and flaggy interbedded sandy lms., 33 shale, and calc. gtzt., contains Composita, Rhombopora, Fenestella, Spirifer, Cliothyridina, corals and Shale, blk. to dk. bwn.-gy., wthrs. med. bwn. -gy., silty and fine sandy, friable, calc. interbedded med. gy. and shaly lms., contains mud cracks, fucoids, 32 Spirifer, crinoids, and other fossil hash. Limestone, med. gy. to blk., wthrs. med. bwn.-gy., fine grained, thin and laminated bedding, platy and 宝宝 flaggy, shaly and sandy, Spirifer and other fossil hash, Composita (?) coquina in base. Limestone, blk. to med. gy. wthrs. lt. gy. very fine grained, thin bedded, 30 platy, contains brachiopod, trilobite, bryozoan, crinoid fragments, base is Η Limestone, dk. gy., wthrs. med. gy., fine grained, thin bedded, flaggy, 29 Limestone, blk. to med. gy., wthrs. lt. gy. and med. bwn., very fine ZZZ grained, thin and laminated bedding, 28 very shaly and carbonaceous, contains Spirifer and other fossil hash. Quartzite, med. gv., wthrs, lt. med. bwn. fine grained, thin bedded, argillaceous, calcareous, top 4 ft. is sandy lms.

NO.	FT ABOVE BASE	COLOR	WT AFTER ACID	PER CENT OF SOLUBLE MATTER	FIELD CARB. CLASSIF.	SILT & CLAY
33 33	1594 1573	DK-MD GY	7.72	22.8	· ·	
32	1558	MDGY	8.99	10.1	C ,	S
		DK-MD GY	8.72	12.8	С	S
32	1532	MD GY	7.63		C	5
32	1514	MD BWN-GY	8.02	23.7	VC	CL
32	1511	MD BWN-GY	7.79	19.8	C	CL
32	1501	LT-MD GY	7.79	22.1	C	SS
31	1488	DK GY		20.2	C	S
26	1209	LT BWN-GY	7.56	24.4	C	S
26	1192	MD BWN-GY	5.43	45.7	VC	CL
26	1182	LT GY	7.62	23.8	C	SS
26	1166	LT GY	6.25	37.5	C	55
25	1159	LT BWN-GY	4.43	55.7	VC	CL
25	1155	LT BWN-GY	5.50	45.0	VC	
25	1143	LT BWN-GY	4.94	50.6	V.C.	SS
25	1125	MD BWN-GY	5.04	49.6	VC	CL
25	1110		5.13	48.7	VC	CL
25	1090	DK-MD GY	6.44	35.6	C	SS
25	1074	MD BWN-GY	5.40	46.0		SS
25		DK GY	5.68	43.2	VC	S
	1054	LT-MD GY	6.00	40.0	VC	S
24	1014	BLK	10.01		VC	CL
23	1000	MD BWN-GY	5.04	NONE	NC	CL
18	868	MD BWN	8.16	49.6	VC	S
18	818	DK GY	9.73	18.4	C	SS
18	802	MDGY	10.03	2.7	VSC	CL
16	795	MD GY-BWN		NONE	NC	SS
16	784	B1,K	10.00	NONE	NC	SS
16	771	MD BWN	10.02	NONE	NC	SS
16	747	MD RED-BWN	9.58	4.2	VSC-	S
16	737	BLK	1.71	82.9	VC	
16	716		9.39	6.0	SC	SS
16	715	MD BWN-GY	9.04	9.6	SC	CL
16	714	MD RED-BWN	8.81	11.9		CL
		DK GY	9.93	0.7	C	SS
16	710	DK GY & BWN	10.01	NONE	VSC	CL
16	705	MD BWN	9.99		NC	CL
16	676	DK GY	10.00	0.1	VSC	S
16	640	DK GY	10.00	NONE	NC	CL
16	630	MDGY	10.10	NONE	NÇ	CL
15	625	MD BWN-GY	10.00	NONE	NC	CL
15	621	MD GY	9.96	NONE	VSC	SS
15	607	LT-MD GY		0.4	VSC	CL
14	577	DK GY	10.01	NONE	VSC	CL
13	572	MD BWN-GY	10.08	NONE	VSC	CL
13	555	MD RED-BWN	9.19	8.1	SC	SS
13	540	MD BWN	5.61	43.9	VC	SS
12	526	MD GY	8.10	19.0	C	5
12	524		8.60	14.0	C	
12	521	DK GY	10.03	NONE	SC	SS
12	515	LT-MD GY	10.00	NONE	SC	SS
11	507	MD GY	10.02	NONE	VSC	\$5
11	484	DK GY	9.99	0.1	NC	S S.
9	439	DK GY	10.00	NONE	NC	S
9	422	MDGY	9.30	7.0		CL
		MD BWN-GY	10.07	NONE	SC	CL
9	402	DK GY	10.02		VSC	SS
9	392	DK GY	10.00	NONE	NC	CL
7	331	MDGY	9.99	NONE	NC	SS
7	319	MDGY	10.20	0.1	VSC	SS
7	309	MDGY	9.70	NONE	NC	CI.
7	302	DK GY	10.01	3.0	VSC	CL
7	288	DK GY		NONE	NC	SS
7	274	MD GY	9.96	0.4	NC	SS
5	210	MD RED-BWN	9.67	3.3	VSC	Cl.
5	203	LT BWN-GY	7.59	24.1	C	SS
3	175	LT BWN	4.50	55.0	VC	SS
3	139		3.57	64.3	VC	
3	137	DK GY	9.97	0.3	VSC	CL
3		LT BWN	3.89	61.1		CL
	132	LT BWN-GY	3.86	61.4	VC VC	S
3	112	BLK	10.10	NONE	VC	SS
3	93	MDGY	10.30		NC	CL
3	83	DK GY	10.00	NONE	NC	SS
3	75	MD BWN-GY	9.10	NONE	NC	SS
3	67	BLK		9.0	SC	SS
1	50	BLK	9.86	1.4	VSC	CL
1	46	DK GY	8.30	17.0	5C	SS
1	42		6.70	32.0	C	
i	38	MD GY LT GY	7.80	22.0	č	\$S SS
			4.80	52.0	VC	C1.
Note:						
ordore mai	mples and 1:1 Hyd erial includes Ca alcareous, 40-80	rochloric acid were used. CO3, MgCO3 and other solubles. Der cent	non i	slightly calcareous, les		
- CONT. (CONT.)	eous, 10-40 per c	E-11 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	CL: Claye	ey, 100 per cent clay siz	a :	4
- Calcar	y calcareous, 5-1		5 Silty.	more than 25 per cent		

Table 1. Color, particle size, and soluble material of the shales in Soldier Canyon, Oquirrh Mountains, Utah.

this pelecypod zone, crinoid stems and a broken brachiopod hash comprise the basal three feet of limestone in Unit 2. Directly above in the base of unit 3 generically unrecognizable pelecypod fragments are present, but are absent five feet above. In the top of unit 4, which occurs 200 feet above the base of the formation, is a two inch bed of faecal-pellet limestone.

Locality 11019 is the most fossiliferous in Soldier Canyon. In the top twenty feet of unit 7 Spirifer, Composita, Amplexizaphrentis, Marginifera, Cleiothyridina, Dictyoclostus, Yoldia, Fenestella, and Nucula are well preserved in the shales and whole specimens weather out. In unit 7, 323 feet above the base of the formation, directly under a one foot bed of brown limestone, is a six to eight foot bed of large crinoid stems that are one to one-half inch in diameter.

In the top of unit 8, directly under the dark gray, brown weathering limestone is a six foot zone within the shaly limestone where large Acanthopecten (?) are abundant especially in the top six inches. A few Spirifer sp. and a few scattered ostracods (?) are mixed with the Acanthopecten (?).

The most fossiliferous shale and limestone units in the formation are 9 and 10. Within the shale bryozoans and brachiopods are abundant. Composita, Spirifer, Cleiothyridina, and Dictyoclostus are by far the most abundant brachiopods; Archimedes oquirrhensis, Archimedes owenanus, Fenestella are the most abundant bryozoans. Polypora debilis, Ptylopora, Marginifera, and Nuculana are present also in the shale. Mr. Lloyd Burkle (personal communication) stated that Archimedes owenanus is a mid-continent upper Osagean and lower Meramecian form. In Soldier Canyon this species occurs stratigraphically higher in upper Chesterian. The fenestellid bryozoans are abundant in the lower part of the shale in unit 9 and Archimedes becomes abundant in the top three to ten feet.

Unit 10 is a marker bed because of its fossil content. Local writers have designated this limestone as the medial limestone in the formation and have attempted correlation on this bed. In Soldier Canyon the bed varies from fifteen to twenty-three feet thick. The limestone contains Spirifer, Composita, Amplexizaphrentis, Orbiculoidea, Kaskia, Griffithides (?), Fenestella, Polypora, Talpaspongia calavata, abundant crinoid stems, and a few ostracods. Archimedes does not appear in the limestone, but in the shales and interbedded limestones above in unit 12 they occur and show their highest appearance in the formation at about 525 feet above the base of the formation. Fenestella and Polypora continue on up to the top of the formation, but in less abundance.

Talpaspongia occurs in less abundance than in the West Canyon area but does occur in the basal three to six feet of unit 10. This sponge appears to be a good guide fossil, but was found only in Soldier Canyon, Ophir Pass, and West Canyon by the writer. In thin section a few poorly preserved endothyrids (?) were observed in this medial limestone, but were unidentifiable.

Nuculana, Caneyella, Yoldia, and Nucula (?) become abundant in the shale zone above unit 10 in Units 12, 14, and 16. A zone of fragmented plants occurs in unit 16 about 708 feet above the base of the formation. According to Mr. Gilbert Greene (personal communication) the only recognizable fragments were those of Cordaites (?). The plants occur in a very silty, brown hard shale. Their fine details seem to be destroyed by weathering. Leaves, seeds, and pods were evident, but the abundant flora found in Lake Mountain was not observed in this section.

Locality 11020 contains Chonetes chesterensis. Only a few specimens were found in the shaly limestones at the top of unit 22 and the base of unit 23.

Locality 11021 is in the top 500 feet of the section and fewer fossils were found than in the lower two-thirds of the formation. Near the middle of unit 28 a dark gray-brown weathering limestone contains a few Spirifer and crinoid hash. The Spirifer are partially replaced by silica and were too fragile to be etched.

A very fine fossil hash is present in the basal two feet of unit 29. Brachiopod and trilobite fragments are the most abundant. Upon thin section study no microfossils were encountered. Unit 30 contains a few Spirifer, Composita, Dictyoclastus, Kaskia, crinoid and bryozoan fragments. A one-foot bed of brachiopod coquina, probably Composita, occurs a few feet above the fossiliferous limestone of unit 30. Throughout the top fifty feet of the formation Spirifer sp. and crinoids appear to be the most abundant fossils. They are intermixed with fossil hash.

Fossils found in Soldier Canyon are as follows:

Buxtonia semicircularis ? (Sutton and Wagner)

Cleiothyridina orbicularis ? (McChesney)

Cleiothyridina sp.

Chonetes chesterensis (Weller)

Composita subquadrata (Hall)

Composita sp.

Derbyia sp.

Dictyoclostus sp. da ssol no sampo disnogsagla Linoproductus ovatus (Hall) Marginifera lasallensis ? (Worthen) Marginifera sp. w see of covered lead page and Pinna sp. (?) Spirifer cameratus ? (Morton) Spirifer occidentalis (Girty) Spirifer opimus (Hall) Spirifer rockymontanus (Marcou) Spirifer sp. (Emos redi angest Description and or galances Acanthopecten sp. Aviculopecten sp. - mlade brase my and colle Caneyella richardsoni (Girty) Edmondia sp. (?) municipalist to the state and the Nucula croneisi (Schenck) Hity 1 626 contains Choncles alless Nucula sp. Nuculana bellistriata ? (Stevens) Nuculina sp. Yoldia glabia (Beede and Rogers) Yoldia sp. it leads 000 got but mest 150 i yill Stegocoelia turritella (Hall) Griffithides sp. (?) Kaskia chesterensis (Weller and Weller) Archimedes oquirrhensis (Condra and Elias) Archimedes owenanus (Hall) Fenestella sp. ni meserq er dand liesof enni Polypora debilis (Elias) Ptylopora sp. Talpaspongia clayata (King) Amplexizaphrentis sp. and to bed joet end Endothyra sp. Figure allies of old synds to a Ostracods acidem formation Soldier to the ost of the os Crinoid columnals Other gastropods Cordiates (?) fragments, seeds and pods

Ophir Pass: The Ophir Pass section is located about four miles to the south and east of the Soldier Canyon section on the divide between the head of the South Fork of Soldier Canyon and the head of Ophir Canyon, in Sec. 11, T. 6S., R. 4W., Tooele County, Utah. The base was taken at the top of the steeply east dipping Great Blue Limestone about 2000 to 3000 feet west of the log cabin located 200 to 300 feet down from the divide in Ophir Canyon. The section was then taped east to the upper contact of the Manning Canyon Shale.

On the pass the complete section is exposed with the lower and upper contact distinct. The lower part up to the fossiliferous limestone (unit 6) was covered with a thin soil so digging for samples was necessary. The upper part of the section is well exposed.

Lateral changes in the Ophir Pass section occur in three zones as follows: first, in a limestone near the base (unit 2) where the lithology changes from fine grained cherty limestone with a fine fossil hash in Soldier Canyon to a fine grained limestone containing intraformational sub-angular conglomerate with pebbles, one-fourth inch by one-half inch in diameter, in the base, along with abundant crinoids, and a few corals and brachiopods. The argillaceous content increases, but the color remains about the same. Second, the quartzite beds (units 8 and 10) increase in thickness but remain about the same in texture and lithology, except for the addition of more argillaceous material. The same quartzite zone in Soldier Canyon does not contain beds over six to eight feet thick, while in Ophir Pass the quartzite thickness is about seventy feet. The third zone of interest occurs 350 feet from the top of the formation where the quartzite and calcareous quartzite of Soldier Canyon (unit 27) becomes more calcareous in Ophir Pass. The color, grain size, and weathered surface remain the same, but the calcium carbonate percentage increases. This same unit (unit 14) in Ophir Pass might be termed a sandy limestone.

The measured thickness of the formation at Ophir Pass is 1347 feet. This thickness is about 200 feet thinner than at Soldier Canyon. The thinning probably is due to possible error in measurement of the dip in the beds, rather than lateral thinning in the formation.

Fossils are common in the limestone and calcarous shale beds. One Brigham Young University locality 11022 was designated in Ophir Pass. Within the Ophir Pass section Spirifer, Composita, and crinoids are most abundant, pelecypods and bryozoans common, and a few trilobites, corals, and sponges are present.

The sponges are of special interest. The zone of Talpaspongia clavata occurs in a three to four foot bed about 3 feet above the base of the fossiliferous limestone (unit 6, medial limestone). The sponges were replaced and weathered, but are characteristically the same shape and weather a light-medium orange-brown color.

The Wewokella zone immediately underlying the Talpaspongia zone in West Canyon was not evident or the specimens were covered in Ophir Pass. A higher zone of Wewokella was discovered in a five foot, black, very fine grained limestone about 200 feet stratigraphically

higher in the section. This is at 796 feet above the base of the formation. The limestone directly overlies the large quartzite bed in the section. Only one good complete specimen was found although the limestone bed was examined laterally for some distance.

Shales in Ophir Pass are covered and more poorly exposed than in Soldier Canyon, which probably accounts for the lack of the abundant pelecypod and bryozoan faunas, rather than lateral faunal changes within the section.

The <u>Chonetes</u> zone so evident in Soldier Canyon could not be recognized in this section. The <u>Archimedes</u> zones (Ophir Pass units 5 and 7) immediately above and below the fossiliferous limestone (unit 6 medial limestone) were not as evident; however, they are present.

The fauna collected (BYU loc. 11022) includes:

Avonia sp. (?) Cleiothyridina sp. Composita subquadrata (Hall) Composita sp. Dictyoclostus sp. or sales percentage incorpa Spirifer opimus (Hall) Spirifer sp. Nucula sp. Yoldia sp. Archimedes oquirrhensis (Condra and Elias) Fenestella sp. Polypora sp. Talpaspongia clavata (King) Wewokella n. sp. Vilisool Vils Amplexizaphrentis sp. Trilobite fragments Crinoid stems and fragments Other pelecypods

ages are of special interest, The sone of Talpaspongia

kin, a three to four foot bed about 3 feet above the base

Manning Canyon (Type Locality). The formation in Manning Canyon crops out discontinuously from the head of the canyon southeast to about one mile above the old townsite of Manning. A section aggregating 1064 feet was measured in three parts at the head and about one-half mile below the head of Manning Canyon in SW1/4 Sec. 4, T. 6 S., R. 3 W., Utah County, Utah. This is located about six and one-half miles southeast of Soldier Canyon. Two Brigham Young University fossil localities, 11023 and 11024, were designated in the canyon.

The basal 410 feet of the section was measured east from the steeply dipping Great Blue Limestone starting where the road leaves the old railroad bed. The basal part is exposed in the road cut. A possible fault exists in the base of the Manning Canyon Shale at this point as evidenced by the difference in strike between the Great Blue Limestone and the Manning Canyon Shale. The average strike in the Great Blue Limestone is about N. 20 W. while the limestones and quartzites in the basal part of the Manning Canyon Shale strike N. 50 W. The Great Blue Limestone over rides 200 to 300 feet of the base of the Manning Canyon Shale. Down the canyon from this measured basal unit black shales occur which could possibly be the basal black shales found in Soldier Canyon, but due to the displacement within the base a measured segment could not be attached to the section.

The middle 390 foot segment of the section was measured across the shallow valley directly east and north of the main road across Mercur Pass. This part of the section except for a few limestones and quartzites was covered, especially in the lower part, making dips and strikes difficult to obtain.

The top 255 feet of the formation was measured on the small east-west ridge about one-fourth to one-half mile north of the Mercur Pass road. The Oquirrh-Manning Canyon contact is similar to the one found in Soldier Canyon in that the pink-gray and purple-gray shales in the Manning Canyon Shale are overlain by basal sandy medium gray, brown weathering limestones of the Oquirrh Formation.

Lithology in Manning Canyon varies little from the sections to the north in Ophir Pass and Soldier Canyon except in the basal part. Quartzites are present about 120 to 150 feet above the base of the measured section, whereas in Soldier Canyon a silty shale occurs at about the same stratigraphic position. A lateral change in lithology occurs in Manning Canyon within the quartzite 570 feet above the base of the measured section. The quartzite is very argillaceous and fine grained which varies from the clean quartzites in Soldier Canyon and Ophir Pass. This quartzite contains Archimedes mcfarlani and Kaskia. The quartzite is calcareous enough that the fossils can be etched.

The remainder of the formation is similar to the type section. The limestone of unit 18 in Manning Canyon is more distinctly jointed and stylolites are present. These structures were not found in the equivalent units in West Canyon (15), Ophir Pass (12), and Soldier Canyon (24).

The top shale unit in Manning Canyon has a pink and purple coloration not evident in other localities. The basal black shales and

limestones of units 1 and 2 in Soldier Canyon and Ophir Pass and the overlying shaly limestones of units 4 and 6 are not exposed in Manning Canyon. This is possibly due to faulting in the base of the formation. The medial limestone, unit 10 in Soldier Canyon, might also be faulted and unexposed in Manning Canyon. However, the brown-gray, brown weathering fossiliferous limestone and shale of unit 7 in Manning Canyon contains the same fauna as unit 12 in Soldier Canyon which there overlies the medial limestone. In Manning Canyon unit 6 could possibly be the medial limestone although no sponges were found.

The fossils found in the basal part of the section in Manning Canyon (Loc. 11023) consist predominantly of Composita, Spirifer, and crinoids, with Archimedes oquirrhensis, Fenestella, corals and trilobite fragments present. They are found in the limestones and shales of units 6 and 7, 325 to 390 feet above the base of the formation.

In the upper 600 feet of the section, Loc. 11024, and argillaceous quartzite, unit 10, about 570 feet above the base of the formation contains Archimedes mcfarlani, Kaskia, sp. fenestellid bryozoans, and fragments of crinoids and brachiopods. This bed contains the only Archimedes mcfarlani found during the study.

The limestones of unit 12 are fossiliferous in the base containing abundant Spirifer, a few pelecypods and other broken fossil hash. About 70 feet from the top of the formation, in the base of unit 22, a zone of Spirifer, and crinoids with fossil hash, is present. Only a few pelecypod fragments were found throughout the shales. Also, no plants were found in the section.

The fossils found in Manning Canyon are:

Archimedes oquirrhensis (Condra and Elias)
Archimedes mcfarlani (Condra and Elias)
Polypora sp.
Kaskia sp.
Fenestella sp.
Spirifer sp.
Composita sp.
Cleiothyridina sp.
Dictyoclostus sp.
Corals
Crinoids

West Canyon: The formation is exposed on the north plunging nose of the Long Ridge Anticline. The section was measured south to north from the north slope of Iron Canyon to the north slope of West Canyon

about one to one and one-fourth miles up Iron Canyon from its mouth in Farmers Flat, Sec. 3, T. 5 S., R. 3 W., Utah County, Utah. The base of the formation is exposed in Iron Canyon and the overlying limestone and quartzite form the main ridge between Iron and West Canyons. The upper shales of the formation form the main valley of West Canyon.

A complete section of 1683 feet was measured in two parts. The lower 1283 feet was measured on the north slope of Iron Canyon. The upper 400 feet is largely covered and the interval was measured in West Canyon about one-fourth to one-half mile northwest from Farmers Flat.

The Great Blue-Manning Canyon contact is taken where the gray and blue-gray, light-gray weathering thin bedded, platy limestones are overlain by light pink-gray and black shales. The contact of the Manning Canyon Shale with the Oquirrh Formation was covered in the area, and an arbitrary boundary was selected. The writer walked the formation laterally to the head of West Canyon but no exposures of the upper part or contact were found.

The section is about eight miles east of the type section in Soldier Canyon and is appreciably thicker. The medial lime stone, Unit 6 in West Canyon, thickens to 98 feet and all the other units appear thickened. The lithology appears similar to the other three sections in the Oquirrh Mountains. The brown "capped" limestones are evident and thus correlation was effected.

The writer measured 1282 feet of section plus an estimated additional 400 feet at the top for a total thickness of 1682 feet. McFarland (1955, p. 12) reports a measured thickness of 2495 feet in Iron and West Canyons. Realizing the 800 feet difference the writer rechecked the thickness and feels the discrepancy may have been in the long shallow dip slope measurements of McFarland's units 62 to 66 (McFarland, 1955, p. 7). Although the formation thickens to the east it is considered that the true thickness is between 1600 and 1700 feet.

Fossil assemblages are similar to the Soldier Canyon faunas, however two sponge zones were found. According to Professor J. K. Rigby (personal communication) these sponges are the first found in the West in Mississippian rocks. Professor H. J. Bissell and the writer found a zone of Talpaspongia clavata in a 12 foot zone in the medial limestone about 682 to 694 feet above the base of the formation.

The writer found a zone of the sponges Wewokella contorta and Wewokella solida. These sponges weather out of the shales in unit 5 approximately 600 feet above the base of the formation. At present the Wewokella zone was found only at this one locality in central Utah.

Three Brigham Young University localities 11005, 11025, 11026 were designated in West Canyon. Locality 11005 is the Talpaspongia zone, 11025 is the Wewokella zone, and 11026 is designated as a general locality for fossils in West Canyon. The shale of unit 5 contains Wewokella, and also has thin beds of fossiliferous limestone containing abundant bryozoans and brachiopods. Archimedes oquirrhensis, Fenestella, Polypora, Dictyoclostus, Composita, Spirifer, Cleiothyridina, Kaskia, corals, and crinoid stems are present. The top of unit 4, 550 to 570 feet above the base of the formation contains abundant Spirifer, Composita, Polypora, Rhombopora?, Fenestella, Dictyoclostus, corals, trilobites and crinoids.

The quartzites and shales of units 10 and 11, 870 to 1000 feet above the base of the formation, contain plant fragments and a few pelecypod fragments. The plants were so fragmented and poorly preserved that generic identification was impossible.

Fossils present in West Canyon include:

Locality 11005

Talpaspongia clavata (King)

Locality 11025

Wewokella contorta (King) Wewokella solida (Girty)

Locality 11026

Cleiothyridina sp.

Composita subquadrata (Hall), seple arb wol

Composita sp.

Dictyoclostus sp.

Spirifer opimus (Hall)

Spirifer sp.

Kaskia sp.

Archimedes oquirrhensis (Condra and Elias)

Fenestella sp.

Polypora sp.

Rhombopora sp. (?)

Nucula sp.

Horn corals Gastropods Other pelecypods Crinoids

Lake Mountain: On the east flank of Lake Mountain a 1910 foot section of the Manning Canyon Shale was measured, east to west, on the west side of Seep Canyon 1000-2000 feet, north and west of the spring, Sec. 36, T. 6 S., R. 1 W., Utah County, Utah. The section is about 25 miles south-east of Soldier Canyon.

The Great Blue-Manning Canyon contact on Lake Mountain is at the top of a dark gray, light gray-weathering cherty limestone. where it is overlain by medium gray shales that become variegated above the base. The contact of the Manning Canyon Shale with the Oquirrh Formation is not readily located, but was taken beneath the dominant medium gray, brown weathering quartzite.

The lithology differs from other areas in that quartzites, subgraywackes and arkoses become more abundant than in any of the other sections. The lithologic summary is as follows:

Rock Type	Thickness in feet	Percentage
Shales	q8 817 979 1231 9	64.5
Quartzites, arkoses and	o to c straightfullen	
graywackes	600	31.4
Limestones (5) of the	179 1810 179 179 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	s fr 4.1
Total	1910	100.0

The section is thicker and the clastics are coarser grained. Limestones are less common and comprise only five per cent of the formation.

Large nodules 8 to 12 inches in diameter are present in the basal shale, unit 1, and consist of calcareous siltstone. Possible repetition may occur in the lower quartzite units on Lake Mountain especially in units 4, 6 and 7. Bullock (1949) suggests numerous faults and repetition of the beds occur on the east side of Lake Mountain. However, in the section measured no fault evidence or repetition was observed. Also, the section is 1910 to 2000 feet thick on the east side of Lake Mountain as compared to Bullock's (1949, p. 18) measurement of 1419 feet on the east and Calderwood's (1951, pp. 63-65) measurement of 1130 feet on the west side. The increase in the thickness of the section further suggests repetition. of to git at safe a sailte in the north slope of Provo Canvon,

The Brigham Young University fossil locality 11027 is designated in the section; however, the invertebrate fossils are rare. A plant fauna is abundant and well preserved. The few pelecypods found in the shaly siltstone of the upper part of unit 6, 1154 feet above the base of the formation, are poorly preserved.

Plants are in the light yellow-brown and yellow-gray shales of unit 13, 1588 to 1602 feet above the base. The same flora was found to the south about one to one and one-fourth miles in the Overcross clay pits located in the NE 1/4, SW 1/4, Sec. 1, T. 7S., R. 1 W., Utah County, Utah. Also to the north end of Lake Mountain in the Powell Pit located SE 1/4, NE 1/4, Sec. 3, T. 6S., R. 1 W., Utah County, Utah. The different genera were consistant throughout the three localities. The plants were not broken or fragmented as in other areas where they have been found. Alethopteris sp. and Neuropteris sp. appear to be most abundant genera.

Fossils collected from Lake Mountain are:

Flora:

Alethopteris sp.

Asterophyllites, 2 spp.

Cordaites sp. fragments

Calamites sp.

Eremopteris missouriensis

Eremopteris sp.

Neuropteris 3 or 4 spp.

Sphenopteris 2 spp.

Other seeds, pods and nuts (?)

Fauna:

Pelecypods
Gastropods
Crinoids
Crinoids

Provo Canyon. The Manning Canyon Shale was visited and observed at three localities in Provo Canyon. The first is above the Great Blue Limestone cliffs north of the Provo City Power Plant in the mouth of Provo Canyon, and the second is on the north side of the Provo River at Canyon Glen Camp Grounds, also south and across the highway. The third is east of the road at the head of Little Rock Canyon and also to the south at the head of Rock Canyon.

At the mouth of Provo Canyon the basal black shales in the formation are exposed, but no correlation of the upper part of these shales with the black shales at Canyon Glen Camp Grounds is positive because of numerous faults in the north slope of Provo Canyon.

At Canyon Glen Camp Grounds the section consists of black shales and fossiliferous limestones. They appear similar to shales and fossiliferous limestones in the lower part of the section, possibly near the black shales and medial limestone of Soldier Canyon units 7 and 12, 250 to 550 feet above the base of the formation. The upper medium and light gray shales and lime stones of the formation are exposed in the south slope of Provo Canyon, across the highway at Canyon Glen Camp Grounds.

The third location is the same area in which Baker (1947) measured 2600 feet of the formation. The section is so poorly exposed that no measurements were attempted. At the head of Little Rock Canyon a unit of quartzite occurs in the base of the formation and contains plant trunks and stems. The trunks show round stem scars. The fossils are poorly preserved, but appear to be one of the lepidodendrons, possibly Lepidodendron sp.

Traverse Mountain. The Manning Canyon Shale is exposed north of the Lehi-Cedar Fort highway on the south slope of Traverse Mountain. Only one location was visited where a partial section of the formation is exposed in the clay pits. The formation on Traverse Mountain is faulted and covered so that only in the clay pits is it exposed. Shale, quartzite, and limestone were observed which possibly compare to the middle part of the section. This is evidenced by the plant flora. The plants occur in a very light gray clay shale. South in Lake Mountain the shales occur in a more brown and yellow shale containing silt. The plants in the Traverse Mountain occurring in the clay shales suggest possible transportation, but no fragmentation of leaves and stems is evident.

The fossils found in the Traverse Mountain are:

Cordaites 2 spp. (?) Demande of blues notices a vidiased shall be set in the set in the

Five Mile Pass. The Manning Canyon Shale forms the low subdued slope that forms Five Mile Pass. The section appears faulted. Two localities were visited. One locality occurs on the southern tip of the

in this area. Pelecypods were observed to be abundant in the upp

Oquirrh Mountain about one to two miles north of the highway in Five Mile Pass, approximately on the Utah-Tooele County line. Variegated shales of red, yellow and gray shades containing pelecypod, gastropod and crinoid fragments were observed in the clay pits at this locality. Also, in this area plant fragments occur in the green-gray fine grained quartzite. From the occurrence of plants in quartzites the exposed partial section probably is in the middle to upper part of the formation.

The other locality visited in the Five Mile Pass area is located about eight miles south of the pass and about five miles east of Topliff, Utah. Only the uppermost 50 to 100 feet of the formation is exposed with a fault contact at the base. However good sedimentary contact with the overlying basal Oquirrh Formation is exposed in this area. The limestones and shales are sandy and appear very similar to the upper two units in the formation at Soldier Canyon (units 31 and 32). Fossil hash containing brachiopod fragments and crinoids occurs in the limestones and pelecypod fragments occur in the shales at this locality.

Stansbury Mountain. One short, faulted, partial section was visited in the northernmost end on the east flank of the range. The shale exposed in the clay pits is variegated green and brown-gray with some black. A few Spirifer sp., productid brachiopods, corals, crinoids and bryozoans were collected from the shale and the fossiliferous brown limestone that forms the west wall of the pit. The fossils were too poorly preserved for identification. The partial section observed probably occurs in the lower part of the formation.

Onaqui Mountains. A section in Johnson Pass in the northern end of the Onaqui Mountains was ivsited. The section appeared too faulted for measurement, but with further walking out of key beds possibly a section could be obtained. The section is predominately shale with quartzite and shaly limestone common in the lower one-half of the formation. Pure white quartzite occurs midway in the section with green and brown-green shales. Medium gray, brown-weathering limestones and brown and gray shales characterize the top of the formation.

Brigham Young University fossil locality 11028 was designated in this area. Pelecypods were observed to be abundant in the upper shale units with few occurring in the shaly limestones in the lower one-half of the formation. No fossil plants were found.

file. One locality occurs on the southern the of the

Fossils collected in the shales and limestones in the Onaqui Mountains include:

Cleiothyridina sp.

Dictyoclostus sp.

Composita sp.

Spirifer sp.

Amplexizaphrentis sp.

Nucula sp.

Fenestella sp.

Crinoids

Gastropods

Pelecypod fragments

formpulse. The second zone consists of the overlying shall have to say the sinus unperlones, and fossitile rous shales. This zone is at the sinustant and contains the mestines are east the top of the zone. The third lithologic form occur the silty shales and interbedded quartities in Soldier Canyon count to 900 feet above the base of the formation. This third are contained the vest above the dark of the formation. This third are contained to the rest. The mare and thicker quartities sed say it is correlated to the rest. The fourth zone consists of carbonaceous non fossiliferous pare and shall mestones containing interbedded calcareous shales. The first about 900 to 1050 feet above the base of the formation. The first about to litter alors the formation is predominantly gray and brow gray shale with a few thin beds of lines tone. Near the top it beds gray shale with a few thin beds of lines tone. Near the top it beds gray shale with a few thin beds of lines tone. Near the top it beds the structs and make calcareous. Midway in this top zone a calcareous opinite can be correlated in Soldier Canyon, Manning Canyon and Ophir Pass. This quartities not as evident in Lake Mountain, Womeyon, One qui Mountains not in Provo Canyon.

Individual lithologic and fossil zones also helped in the correlation of the sections throughout central Utah.

The fossiliferous medial limestone unit is evident in Soldier, Canyon, Ophir Pass, Manning Canyon, West Canyon (Plate 4) and helpful in determining the lower part of the section. The fauna is too different in this limestons than in the overlying and underlying units, however the abundance of fossils and the presence of Talps clavata in the basal part of this medial limestone makes it a good for correlation. The shale with interbedded limestone, overlying underlying this medial limestone also help to correlated this lime with comparable facies. Abundant bryozoans, especially Archime and Fenestella, the brachiopds Composita and Spirifer, and crind and Fenestella, the brachiopds Composita and Spirifer, and crind stems help to correlate these shales which in turn help designated medial limestone in West Canyon, Ophir Pass and Manning Canyon medial limestone in West Canyon, Ophir Pass and Manning Canyon

Fossils collected in the whales and Emestones in the Coartil

CONCLUSIONS

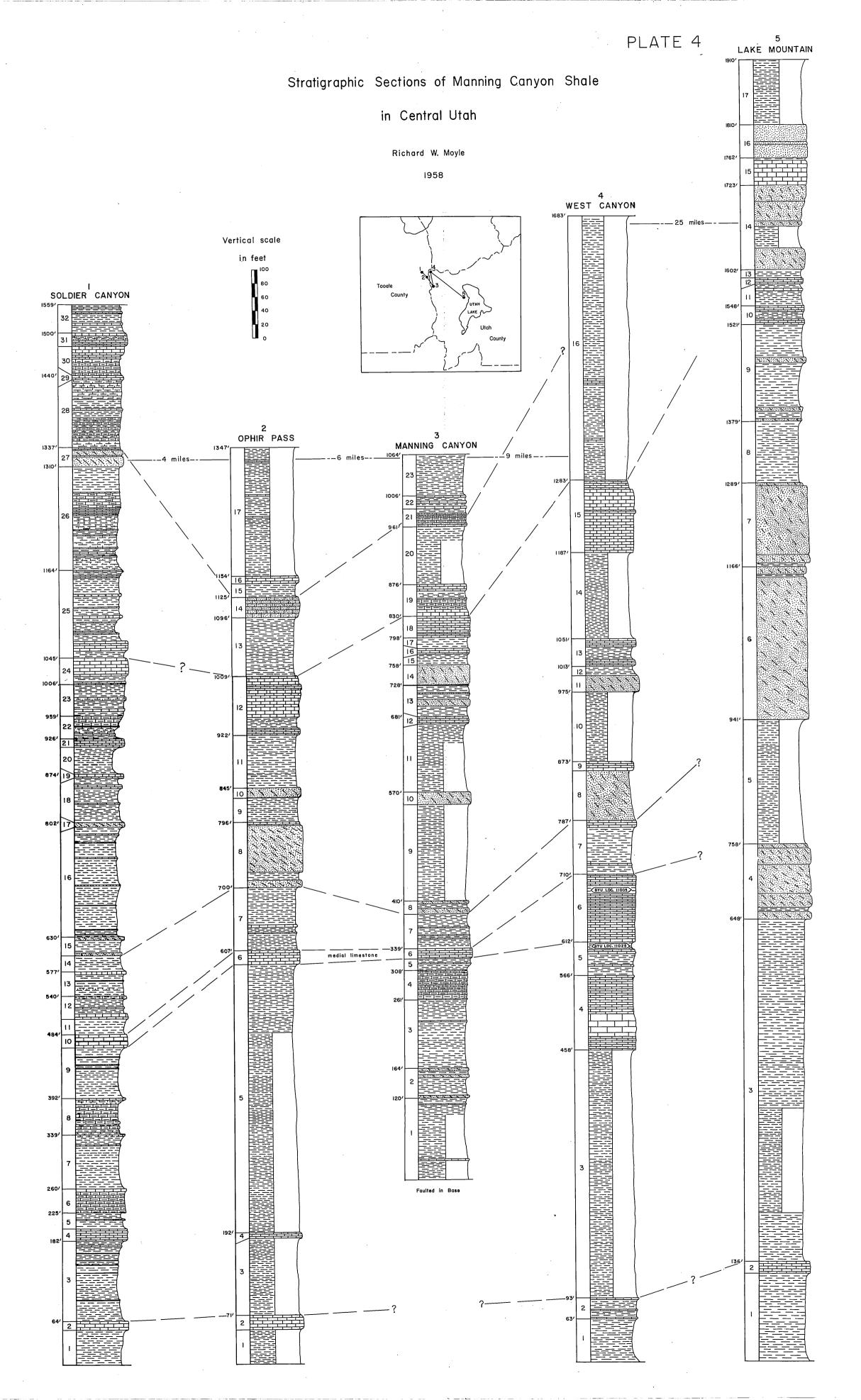
Correlation and Age

Correlation throughout central Utah within the Manning Canyon Shale is evident by means of key lithologic units. Four sections in the Oquirrh Mountains, one in the Onaqui Mountains and one in Provo Canyon appear to correlate well by means of these beds and fossil zones.

The first zone is in the black shales in the basal 200 feet of the formation. The second zone consists of the overlying shall limestone, fossiliferous limestones, and fossiliferous shales. This zone is about 200 to 550 feet above the base of the formation and contains the medial limestone near the top of the zone. The third lithologic zone occurs in the silty shales and interbedded quartzites in Soldier Canyon about 550 to 900 feet above the base of the formation. This third zone contains more and thicker quartzite beds as it is correlated to the east. The fourth zone consists of carbonaceous non fossiliferous pure and shaly limestones containing interbedded calcareous shales. The zone is about 900 to 1050 feet above the base of the formation. The fifth and top lithologic zone in the formation is predominantly gray and browngray shale with a few thin beds of limestone. Near the top it becomes sandy and more calcareous. Midway in this top zone a calcareous quartzite can be correlated in Soldier Canyon, Manning Canyon and Ophir Pass. This quartzite is not as evident in Lake Mountain, West Canyon, Onaqui Mountains nor in Provo Canyon.

Individual lithologic and fossil zones also helped in the correlation of the sections throughout central Utah.

The fossiliferous medial limestone unit is evident in Soldier Canyon, Ophir Pass, Manning Canyon, West Canyon (Plate 4) and is helpful in determining the lower part of the section. The fauna is not too different in this limestone than in the overlying and underlying units, however the abundance of fossils and the presence of Talpaspongia clavata in the basal part of this medial limestone makes it a good unit for correlation. The shale with interbedded limestone, overlying and underlying this medial limestone also help to correlated this limestone with comparable facies. Abundant bryozoans, especially Archimedes and Fenestella, the brachiopds Composita and Spirifer, and crinoid stems help to correlate these shales which in turn help designate the medial limestone in West Canyon, Ophir Pass and Manning Canyon.



Plants are good fossils for correlation in the middle and upper part of the formation within the sections measured in central Utah. The plants on Traverse Mountain may be an exception because they seem to appear stratigraphically higher in the section.

The brown "capped" limestones in the lower part of the formation appear to be helpful to correlate. However, because of their abundance in each section, care should be taken when using them individually to correlate.

The limestone of unit 24 in Soldier Canyon, along with its equivalent units in Ophir Pass (12), West Canyon (15), and Manning Canyon (18), can be used to correlate the upper part of the formation. A limestone with similar appearance in Provo Canyon might also be correlated with this unit.

Correlation with the major transgressions and regressions of the Manning Canyon sea is evident (Plate 5), especially in the sections in the Oquirrh Mountains. The Lake Mountain and Onaqui Mountain sections correlated partially using the lithologic cyclic nature of the formation.

Due to temporal transgression, the Maning Canyon Shale has been mapped by most geologic workers as Mississippian-Pennsylvanian. A few geologists working throughout central Utah have suggested differentiating Mississippian and Pennsylvanian rocks of the Manning Canyon Shale at the top of the predominate fossiliferous dark gray-blue limestone, termed the "medial limestone," that occurs about 450 feet above the base of the formation. Others suggest the time boundary occurs 200 to 330 feet higher in the formation at the zone of the arkosic and graywacke quartzites.

Elias (1956, p. 71) suggests that a similar controversy occurs relative to the Caney Shale throughout Oklahoma and that it might be solved somewhat by a practical lithologic differentiation. He worked with the Caney Shale in the northern and southern Arbuckle Mountains and states the differentiation between the two ages in the shale occurs between, ". . . the dark gray or black, more or less bituminous shale, with large to very large limestone and small phosphatic concretions below, to gray, non-bituminous hale with medium-size ferruginous, brownish to reddish concretions above."

It is impracticable to use such a lithologic change alone in the shales to differentiate time within the Manning Canyon Shale in central Utah. However a local differentiation based on the shale color change and the bituminous content of the shales and limestones might possibly be observed in the type section in Soldier Canyon (Plate 3). The time change liekly occurs in a zone between 1050 to 1150 feet above the base of the formation in the basal part of unit 25 or just above the dark gray and dark brown limestone of unit 24. Below unit 24 in the section the shales are predominantly black, dark gray and dark brown, very carbonaceous and contain zones of abundant calcareous siltstone nodules and concretions. Above unit 24 in the section the shales are medium and light gray, medium and light brown-gray, and are slightly carbonaceous.

The quartzites and sandy limestones above unit 24 conform closer to the overlying sandstones, quartzites, and limestones of the Oquirrh Formation than to the underlying quartzites, siltstones and limestone of the typical Manning Canyon Shale.

The paleontologic change between the faunas of the Mississippian and Pennsylvanian age throughout central Utah is not abrupt or conspicuous. The fauna change seems to occur without a profound break and may be transitional in all of the areas studied in this report. The writer feels that if a break is apparent in the fauna, it occurs in the Soldier Canyon section above unit 12, 550 feet above the base of the formation. The abundant fauna typified by Archimedes, Composita, Amplexizaphrentis, and Talpaspongia found in the formation below unit 12 disappears above the unit. Chonetes chesterensis occurs above this suggested break, but disappears about 1000 feet above the base of the formation in the top of unit 23. Dictyoclostus, Spirifer, Cleiothyridina, Derbya, and other brachiopods seem to carry thoughout the entire formation.

A distinct difference in lithology occurs directly above unit 12 in Soldier Canyon. The sediments show near shore to terrestrial conditions. The abundance of plants within the formation also occur in this overlying zone. The plants and the quartzites of this zone (units 14 to 18) in Soldier Canyon suggest a time break at this point, marked by a change in sedimentary conditions. The changes likely occurred in provenance and environment of sedimentation, but may have been more gradual than abrupt. The writer realizes that the time boundary is suggested near the close of a regression and not at the customary beginning of a transgression. The plants resemble Pennsylvanian forms thus suggesting the time boundary be arbitrarily drawn below the plant zone in unit 16; however, the se plants might well be Upper Mississippian in age.

The Mississippian-Pennsylvanian time boundary may be arbitrarily drawn in the Ophir Pass, Manning Canyon, and West Canyon sections at about the same stratigraphic position as in Soldier Canyon. The time boundary in Ophir Pass occurs above the

medial limestone unit 6 in the silty shales and quartzites of units 7 and 8 (Plate 4). Plants were not found in these units in this locality. In Manning Canyon the boundary might be placed in the shales and quartzites of units 7 to 10. The boundary in West Canyon may occur above the medial limestone of unit 6 in the shales and quartzites in units 8 to 10. Plant fragments occur in unit 11 within the West Canyon section (Plate 4).

Environment of Sedimentation

The environment of the shales as suggested by fossils, varying amount of silt and clay, varying amount of organic material and iron, and the calcareous content suggests both near-shore and shallow epineritic to shallow infraneritic conditions of deposition.

Sediments were possibly deposited on an unstable shelf directly adjacent to the deeper zones of the geosyncline. Calcareous content suggests a tidal flat environment for the brown and brown-gray shales. The brown color is probably due to oxidation of iron during deposition suggesting oxidating shallow near-shore conditions. Some of the silty and sandy shales are calcareous. The warmth of the water in the shallow near-shore conditions may be the cause of calcium carbona te in these sediments.

Abundance of organic material suggests deposition farther from shore or under slightly reducing conditions for the black shales. However, Hager (1928, p. 936) notes that "Black muds are deposited in estuaries, lagoons, and enclosed parts of the sea, where the water is relatively quiet and wave action is slight, or also in depression on the continental shelf where there is little current." The same type of origin might explain some of the black shales within the formation.

In summary of the different shales in the Manning Canyon Shale, two different vertical sequences exist in the section. The vertical lithologic change can probably be taken for lateral shift near-shore to deeper marine conditions. The first sequence consists of a change from near-shore to deeper marine, sandy shales, silty shales, clay shales, calcareous clay shales. The second sequence consists of calcareous sandy shales, calcareous silty shales, clay shales, and calcareous clay shales.

The environment of the limestones appears to have been shallow marine to open marine. The fossil fauna suggests epineritic clear water to shallow infrancritic clear water. The dense, hard, very fine grained to finely crystalline limestones devoid of fossils suggest

and at a time anatamillath

the outer edge of the shelf to a deeper marine lithotope. Fossiliferous limestones are present in the lower one-half and in the top of the formation, while the lithographic limestones are in the middle and upper part of the formation in central Utah.

Quartzites near the middle of the section at Soldier Canyon, Ophir Pass and West Canyon suggest that these sediments were washed into the geosyncline at a faster rate than the basin was sinking or that stabilization was occurring. This is evident from the reworked character of the sand grains, from cross-bedded and from the well-sorted subangular and subrounded grains. The reworked character of the sand grains is evident also on Lake Mountain.

The quartzite beds on Lake Mountain appear to retain the same thickness laterally. No evidence of channeling or lensing was evident, so possibly the sands were deposited as a blanket sediment in shallow sea conditions.

Current action during deposition of the shales and limestones was very slight. Fossils occur neither broken nor fragmented. Of special interest testifying to quiet bottom conditions is the Archimedes and Fenestella zone in Soldier Canyon about 400 to 525 feet above the base of the formation. Within this zone very fragile spiral fronds are well preserved in place and are unbroken.

Samples of quartzite from the basal Oquirrh Formation in central Utah, especially in Soldier Canyon, show an increased argillaceous content and less rounding of the grains. This suggests a faster subsiding basin at the close of Manning Canyon and the beginning of Oquirrh time.

Some of the sediments, especially sands, laid down in the formation were possibly of a terrestrial environment because of the presence of land plant fragments. Also, according to Bissell (personal communication) the formation locally contains a few thin coal beds.

The cyclic nature of the formation and the terrestrial fossils (possibly 2 to 3 zones) suggest that the strand line of the Manning Canyon sea migrated back and forth several times, and that transgression-regression occurred.

obably be taken for lateral shift near shorello

retalline limestones devold of fessils suggest

Cyclic Sedimentation. From the detailed study of the type section of the formation in Soldier Canyon cyclic sedimentation is evident.

Lithologic Oscillation Chart of Manning Canyon Shale

The cyclic sedimentary nature of the Manning Canyon Shale was believed by previous writers to be due to local diastrophic control, which consists of a combination of gradual sinking of the geosyncline and a gradual rising of the surrounding land masses.

Throughout central Utah the cyclic nature of the formation is best evidenced in plotting the different lithologies of the measured sections, then outlining the cyclic transgressions and regressions of the shallow Manning Canyon sea (Plate 5). Two transgressions and three regressions are evidenced in both lithologic and fossil changes.

The first transgression of the sea is evidence in the Soldier Canyon section (Plate 5, Sec. 1) by limestones and interbedded shale of units 4 to 13, 180 to 577 feet above the base, with the second one evident by the limestones and interbedded shales of units 21 to 26, 914 to 1310 feet above the base of the formation.

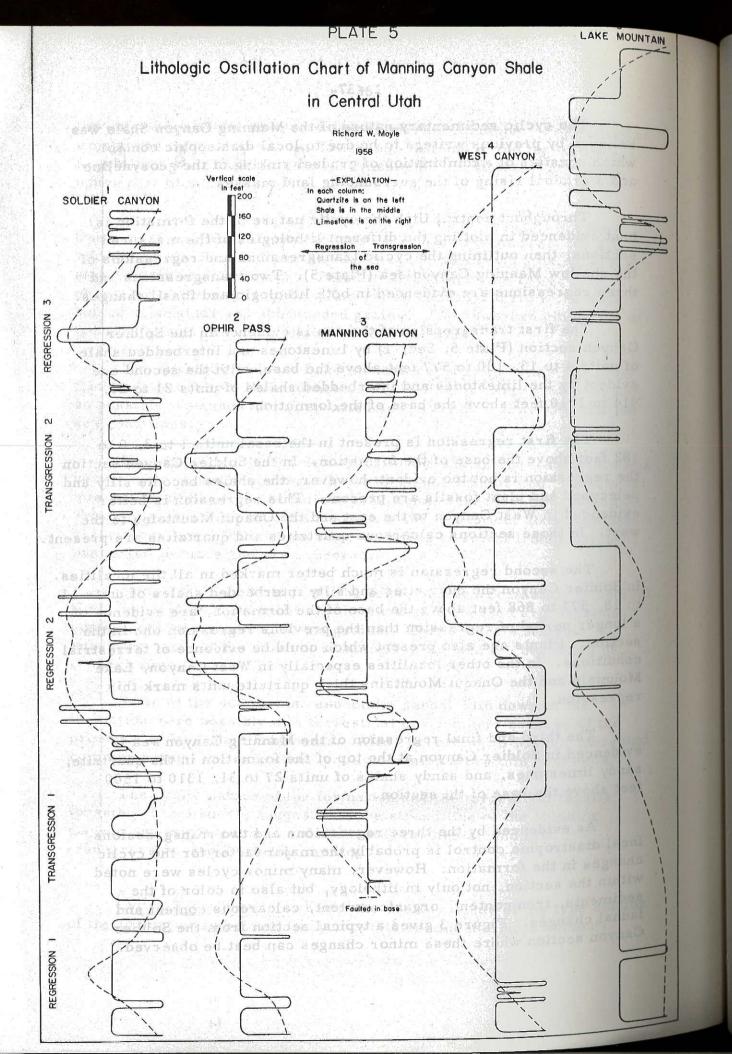
The first regression is present in the basal units 1 to 3, 0 to 182 feet above the base of the formation. In the Soldier Canyon section the regression is not too evident; however, the shales become silty and pelecypod and plant fossils are present. This regression is better evidenced in West Canyon to the east and the Onaqui Mountains to the west. In those sections calcareous quartzites and quartzites are present.

The second regression is much better marked in all the localities. In Soldier Canyon the quartzites and silty interbedded shales of units 14 to 18, 577 to 868 feet above the base of the formation, are evidence of a longer period of regression than the previous regression one in the section. Plants are also present which could be evidence of terrestrial conditions. In the other localities especially in West Canyon, Lake Mountain and the Onaqui Mountains thick quartzite units mark this regression.

The third and final regression of the Manning Canyon sea is evidenced in Soldier Canyon at the top of the formation in the quartzite, sandy limestones, and sandy shales of units 27 to 31, 1310 to 1500 feet above the base of the section.

As evidenced by the three regressions and two transgressions local diastrophic control is probably the major factor for the cyclic changes in the formation. However, many minor cycles were noted within the section, not only in lithology, but also in color of the sediments, iron content, organic content, calcareous content and faunal changes. Figure 3 gives a typical section from the Soldier Canyon section where these minor changes can best be observed.

I was been the sediments boild derive in



Also the minor cyclic changes in the formation are evident in the fossils (Figure 5). It is unreasonable to explain each of these minor changes as a separate diastrophic movement within the depositional basin. Climate may be a determining factor. Glaciation during the Upper Mississippian to Middle Permian possibly had some influence and should not completely be overlooked.

Wheeler (1957, p. 1985) states, "The persistent belief that late Paleozoic glaciation was confined to the Permian has influenced some to favor the diastrophic interpretation." Yet, in noting Permian glaciation studies of different men, they have dated tillites as early as the lower part of middle Carboniferous time. Also, as Wheeler (1957, p. 1998) states if ". . . others are correct in their belief that eustatic sea level fluctuation is the primary factor in sedimentation control, glaciation can be reasonably regarded as the most likely cause of fluctuation." He further states that glaciation constitutes an important factor in cyclothemic genesis.

The cyclic nature of the Upper Mississippian sediments in central Utah is evidently not as observable as in the Illinois Basin and other locations in the eastern United States, but during this present study, the sediments in the Manning Canyon Shale show a definie cyclic nature. With more detailed work in central Utah with the Upper Mississippian sediments a custatic change in the sea due to glaciation might become more evident.

Source Area. The varying thickness of the formation observed eastward from the Gold Hills through the Onaqui Mountains, Stansbury Mountains, Oquirrh Mountains, Lake Mountain, and Provo Canyon demonstrates thickening towards the east. Thickness varies from about 500 feet in the Gold Hills to 2600 feet in Provo Canyon. Also the size and amount of clastic sediments increases towards the east.

Black and brown carbonaceous shales in the base of the formation indicate that at least some of the sediments were derived from a source area containing a thick layer of mature soil. This area probably was to the east of the Mississippian Cordilleran miogeosyncline or a positive craton in or near it. The Colorado Mountains uplift was in its initial stage and sediments shed from its west slope into this geosyncline which were transported by ocean currents over wide areas.

Lake Mountain area, however, because of the thick beds of sandstone and quartzite suggests two possible areas from which sediments could be derived. First, the sediments could have been derived from the east as suggested by eastward thickening of the section and eastward increasing clastic size. Second, the sediments could derive from



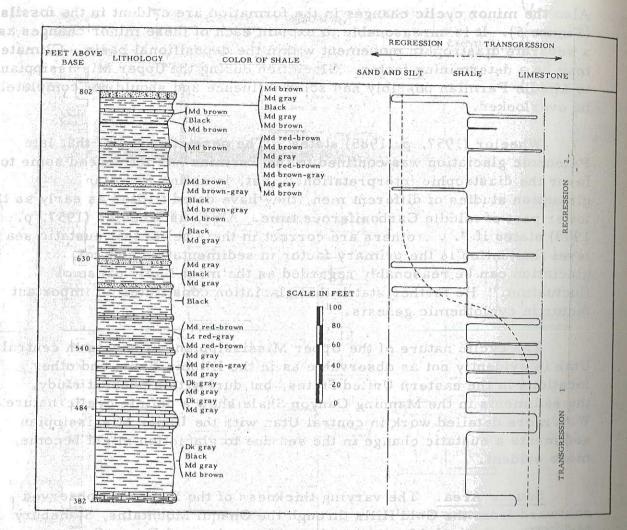


Figure 3. Cyclic patterns in Manning Canyon Shale, Oquirrh Mountains, Utah.

funt of clastic sedunents increases towards the east

pre-existing sediments of lower and middle Paleozoic sediments exposed in a small emergent high along part of western Utah as evident from a preliminary isopach map of Mississippian rocks prepared by Bissell (personal communication). Less than 500 feet of sediments represent the total Mississippian rocks in the area of this emergent high. In the San Francisco Mining district Upper Mississippian and Lower Pennsylvanian sediments are missing. This suggests either an area of non-deposition or an emergent area subject to erosion. Of the two suggestions there is no evidence of the latter, but both suggest an emergent local high from which sediments could have been derived.

Also, a source area to the north and west of central Utah should be considered a strong possibility. In this area the Antler orogeny, a late Mississippian and early Pennsylvanian disturbance, caused an uplift that could evidently have supplied sediments which were transported to the east and south into central Utah. Dott (1955, pp. 2222-23) reports the Tonka Formation (Mississippian-Pennsylvanian) of eastern Nevada consists of coarse conglomerates and quartzites which pass vertice-laterally into shales.

The coarse clastics intertonguing eastward with the black shales of the Upper Mississippian and the Lower Pennsylvanian limestones in central Utah prove this western source for at least some of the sediments. However, in Lake Mountain the abundance of coarser clastics that are sub-rounded and sub-angular suggest a much closer source area. This favors the second supposition above, namely that sediments likely were derived from the pre-existing Paleozoic sediments exposed in a small emergent high along part of western Utah. (Figure 4).

Another possible source area is to the south in southern Utah and northern Arizona. Nolan (1943, p. 155) states that "Over much of southern Nevada and the adjoining parts of southern California Upper Mississippian beds appear to be missing." Noble (1922) suggests that similar conditions prevail in the Grand Canyon region.

Figure 4 shows the thickness of the Manning Canyon Shale in central Utah and the thickness of correlative formations in the contiguous area. The thickness in turn suggests the four source areas of sediments in the Upper Mississippian and Lower Pennsylvanian time.

Paleoecology

The paleoecology has not been discussed for each unit and sub-unit of the formation. In this report six major zones, where one lithology predominates, and a few other minor zones of special interest are discussed.

The basal 50 feet of predominantly black shale in the formation contains abundant pelecypods and plant fragments which suggest shallow lagoonal conditions close to shore where the plant fragments could have washed in without undue maceration, or were aquatic types which lived in that environment. The presence of pelecypods suggests

Explanation of Figure 4

References to numbered localities

- 1 Agatston 1957 St. molgnon sancos se electron carvalland
- 2 Unterman and Unterman 1944 and the state well so they are
- 3 Sadlick 1955
- 4 Unterman and Unterman 1944
- 5 Williams 1948-As and Trawell and ball religious was M rag
- 6 Peace 1956 and it not solve are seen sitt svorg astu-
- 7 Peace 1956 nebrada and ristanoM axed at a reviewoil , a months
- 8 30 Olson 1956 saggra religing due but bebouwerdes are said
- 9 Moyle 1958 (this report) 10 Soos and a loval as fit as
- 10 Baker 1947 Ballsty bear of the beart beart beart with the real graduations
- 11 Bissell 1950 of sound lemergent high sound posed in a small emergent high sound posed in
- 12 Moyle 1958 (this report)
- 13 Bissell 1950
- 14 Moyle 1958 (section visited during present study)
- 15 Rigby and Bissell (personal communication)
- 16 Dott 1955 gradius to strad mere of be sit bas source.
- 17 Bissell (personal communication)
- 18 to Nolano1935 base O ent on lineary enoughnos resemble to the
- 19 Youngquist 1951
- 20 Ogden 1951 D zatasaM ed 'n saandalds edt ewode A
- 21 Nolan 1943 of temporal evister out to seem! Life als but
- 22 Bissell (personal communication)
- 23 Weller 1948al bha usigglesseel A ragglest status losses
- Westgate and Knopf 1932
- 25 Bissell (personal communication)
- Nolan 1943 and Weller 1948

rology has not been deschisched for each unit and traction. In this report six major somes, where bits matest art a few offer made conce of sperial interest as a conce of sperial interest as a conce of sperial interest and content fragments which will suggest the plant tragments in without undues parentian, or were equatic types a concent undues passence of pelecypods suggest and content undues passence of pelecypods suggest and content undues prosence of pelecypods suggests.

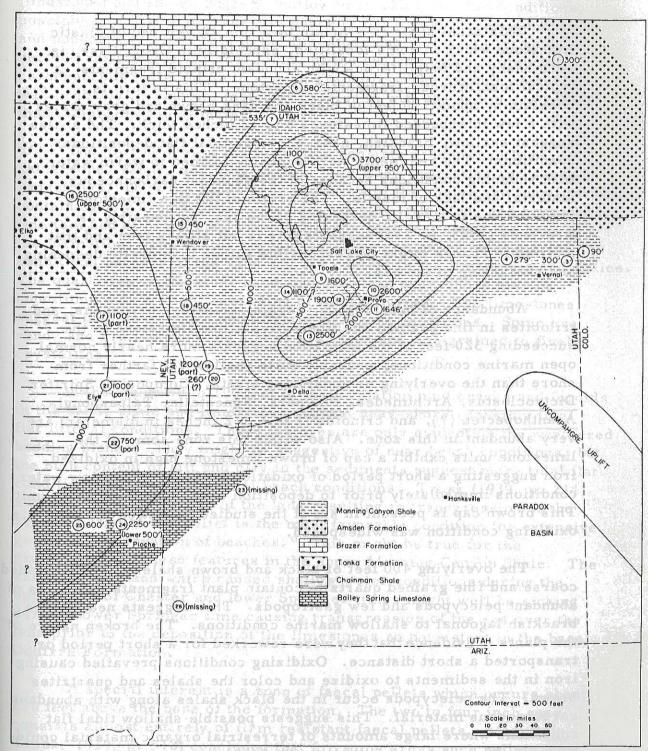


Figure 4. Isopachous map of Manning Canyon Shale and correlative formations in central Utah and eastern Nevada.

deline lime stone and calcareous shale contains thin beds o

the water was probably brackish, slightly aerated, yet somewhat turbid. The fauna shows no evidence of being washed into the area.

This basal shale is overlain by a thin 12 foot bed of clastic limestone which in some localities, especially in Ophir Pass, is very hashy and fossiliferous with abundant crinoid fragments—a typical fossiliferous-fragmental limestone. This in turn is overlain by about 110 feet of black and dark brown and gray shales with a few interbedded limestones 6 to 12 inches thick. The shales suggest shallow, brackish, near—shore or tidal flat depositional conditions. The sediments were probably formed in warm water with reducing conditions because of the abundant carbonaceous material and lack of fossils. The origin of the thin limestones might be due to a cyclic deepening of the sea, but possibly as Bruchner (1953, p. 236) suggests this is not always true and that a climatic change might be responsible rather than cyclic deepening and shallowing of the sea.

Abundance of brachiopods, corals, crinoids, bryozoans, and trilobites in fine grained and finely crystalline limestone in the succeeding 320 feet of the formation suggests more nearly normal open marine conditions, probably quite shallow yet farther from shore than the overlying or underlying strata. Composita, Spirifer, Dictyoclostus, Archimedes, Polypora, Fenestella, Talpaspongia, Acanthopecten (?), and crinoids (1 to 1 1/2 inches in diameter) are very abundant in this zone. Also within this zone some of the limestone units exhibit a cap of brown limestone high in oxidized iron suggesting a short period of oxidation probably in shallow water conditions immediately prior to deposition of the overlying shale unit. This brown cap is present in six of the studied localities showing this oxidizing condition was widespread.

The overlying 400 feet of black and brown shales and interbedded coarse and fine grained quartzite contain plant fragments and zones of abundant pelecypods and few gastropods. This suggests near-shore brackish lagoonal to shallow marine conditions. The broken nature of the plants is evidence that they were reworked for a short period or transported a short distance. Oxidizing conditions prevailed causing iron in the sediments to oxidize and color the shales and quartzites brown. The pelecypods occur in the black shales along with abundant carbonaceous material. This suggests possible shallow tidal flat conditions where large amounts of terrestrial organic material could be washed in and concentrated.

The next overlying 200 feet of interbedded fine grained to finely crystalline limestone and calcareous shale contains thin beds of

Chonetes, and Spirifer. The limestones which contain fossils suggest shallow normal marine conditions, whereas the non-fossiliferous limestones and shales suggest shallow near-shore marine conditions, possibly in a tidal flat. The abundance of black carbonaceous material and non-oxidized iron both in the limestone and shale support non-aerated reducing conditions. Some of the shales contain so much carbonaceous material as to appear graphitic. The organic content was probably terrestrial material carried in with the sediments rather than indigenous fossil material.

The top 500 feet of the formation consists predominantly of calcareous, sandy, and silty shales, with a few quartzites and sandy limestones. Only the limestones and a few of the shales are fossiliferous, containing few brachiopods, crinoids, corals, trilobites, and bryozoans. One thin bed of brachiopod coquina occurs 1486 feet above the base of the formation. The shales suggest shallow near—shore quiet marine conditions because of the sand and silt content in the limestones and shales. The quartzites on the other hand suggest clean near—shore marine conditions with a slight reworking of the sediments. The limestones suggest both near—shore and slightly off—shore conditions. Some are sandy and argillaceous and others are clean and fine grained to finely crystalline.

In the upper 50 to 100 feet of the formation mud cracks and fucoids in the sandy shale and limestones suggest near-shore, possibly tidal flat littoral conditions. These features are only evident on the weathered bedding planes. They might be suggestive of near beach conditions, but the high content of silt and clay in the sediments suggest more tidal flat conditions rather than normal beach conditions. Trask (1939, p. 214) states, 'Gradual sinking of the land resulting in transgression of the sea over land of low relief is the most favorable condition for extensive geologic preservation of beaches." This might be true for the preservation of these features in the upper Manning Canyon Shale. The eastward highland which caused shallow water conditions during the upper part of Chester and lower Springer times might well be sinking during Lower Springer time causing transgression of the sea over the land prior to the deposition of the limestones so persistent in the basal Brachiopode, tornals, bryozogna, trilobles and proposed atomat emparitic marine conditions. However, in some sediment

Of special interest is a zone of faecal pellets which occurs about 200 feet above the base of the formation. The two to four inch zone consists almost entirely of firm resistant faecal pellets. Moore (In Trask, 1939, p. 516) concluded that for some areas such pellets owe their origin to detritus eaters perhaps one of the pelecypods or holothurians.

Thinwever, are associated directly with Spirifer and crinoid of the pelectricity and the pelectricity and brachiopods show no evidence of being transported.

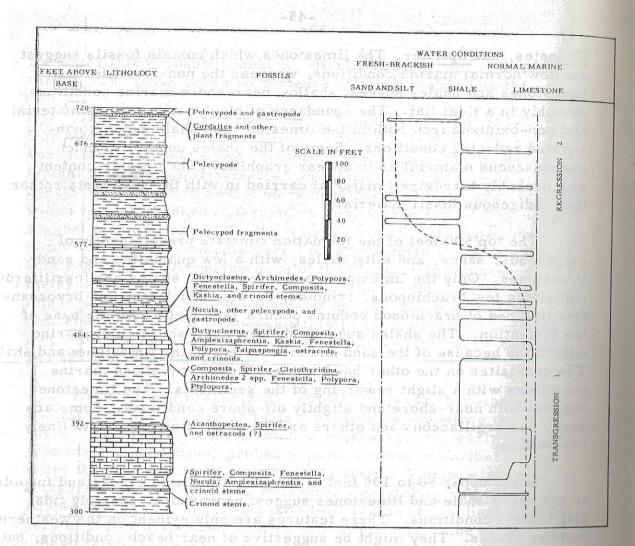


Figure 5. Fossil-sediment facies patterns in Manning Canyon Shale, Soldier Canyon, Oquirrh Mountains, Utah.

low relief is the most favorable condition for extensive

Fossils are associated throughout the formation with certain sedimentary facies. Figure 5 gives a typical fossil sediment facies pattern in the Soldier Canyon section and also gives the possible salinity of the water in which the fossils lived.

Brachiopods, corals, bryozoans, trilobites and crinoids characterize normal epineritic marine conditions. However, in some sediments within the formation, some bryozoans, brachiopods, and crinoids, are directly associated with silty shales and seem to be indigenous of that environment. No evidence of current reworking or fragmenting the fossils is present.

Pelecypods and gastropods suggest brackish water conditions. In Soldier Canyon they are commonly found in the dark gray and brown shales. Some, however, are associated directly with Spirifer and crinoids which suggests more normal marine environment for some of the pelecypods because the crinoids and brachiopods show no evidence of being transported.

The plants are found in silty shales, fine grained quartzites, and in one locality in pure clay shale. The shales and quartzites vary from ash gray, light yellow-brown for the shales to brown-green, green, and green-gray for the quartzites. The plants suggest very near-shore to terrestrial conditions. In Lake Mountain and Traverse Mountain the fossils have been found in clay and silt shales and are well preserved. In West Canyon, Five Mile Pass, Soldier Canyon, and Provo Canyon they are found in the quartzites and very silty shale and also show evidence of being reworked by currents.

Bive Line and chil where it is everlain by a medium gray shale we is had pink may in the base and becomes black in the top. The had 260 feet of the shale up to the top of unit 6 was measured and describe on the morth should treming spar about 200 to 300 feet, we stood the firming the firming the roll in the morth should side of the road after entering the canyons.

Served as a key bed to identify the sume und that spars to line each a lower had a suidued spar where the next set feet of the section when sured and described.

Because of the allinial cover on the fourth spur the next 2434 eet were meanted and described on the next spur east of the little four east of the mine hinsel.

Due to faults and folds in the upper part of the section on the first apair the top 514 feet of the Manning Canyon Shale and should reset of the assal Oquirth formation were measured and described the aext apart what or the fourth sput east of the mine tunnel and the aext approximent quartisite bed (unit 17) was the key bed which effected of setting measurement of the section to the fifth spur approximant timestone in the top of unit 24 was the key bed along which offset of the section back to the fourth spur was accomplish

Formation was taken where the shales and the shaly limestones change upward to sandy limestones and calcareous quartzites (based unit 33). These medium gray, brown weathering sandy limestor and calcareous quartzites should not be mistaken as the 27 foot sind calcareous quartzite of unit 20 because above unit 27 the shales are lighter to the typical Manning Canyon Shales whereas the shaly astones and limestones as well as the fossil fauna above unit 3 shales are shale basal Oquirin formation.

ed score the greek assessed Measured Sections and to year or years and the score with the conditions later to the conditions.

Soldier Canyon. A composite section was measured in four parts (Plate 1) on the north side of Soldier Canyon about one-half mile above the mouth, Sec. 33, T. 5 S., R. 4 W., Tooele County, Utah.

The base of the measured section was taken at the top of the Great Blue Limestone cliff where it is overlain by a medium gray shale which is light pink-gray in the base and becomes black in the top. The basal 260 feet of the shale up to the top of unit 6 was measured and described on the north-south trending spur about 200 to 300 feet west of the first mine tunnel on the north side of the road after entering the canyon.

The dark gray-brown weathering limestone in the top of unit 6 served as a key bed to identify the same unit four spurs to the east on a lower more subdued spur where the next 542 feet of the section were measured and described.

Because of the alluvial cover on the fourth spur the next 243 feet were measured and described on the next spur east or the fifth spur east of the mine tunnel.

Due to faults and folds in the upper part of the section on the fifth spur, the top 514 feet of the Manning Canyon Shale and about 140 feet of the basal Oquirrh Formation were measured and described on the next spur west or the fourth spur east of the mine tunnel.

The prominent quartzite bed (unit 17) was the key bed which effected offsetting measurement of the section to the fifth spur, and the dominant limestone in the top of unit 24 was the key bed along which offset of the section back to the fourth spur was accomplished.

The contact of the Manning Canyon Shale and the Oquirrh Formation was taken where the shales and the shaly limestones change upward to sandy limestones and calcareous quartzites (base of unit 33). These medium gray, brown weathering sandy limestones and calcareous quartzites should not be mistaken as the 27 foot calcareous quartzite of unit 27 because above unit 27 the shales are similar to the typical Manning Canyon Shales whereas the shaly limestones and limestones as well as the fossil fauna above unit 33 resemble the basal Oquirrh Formation.

The writer felt that inasmuch as a general description of the Soldier Canyon section is given on Plate 3 it would be repetitious to include in the appendix the detailed emended type section of the Manning Canyon Shale in Soldier Canyon. The detailed section is placed on file with the Geology Department of the Brigham Young

noitari 495

University at Provo, Utah, and is available on request.

Manning Canyon. A composite section was measured in three parts at the head and about one-half mile below the head of Manning Canyon, SW 1/4, Sec. 4, T. 6S., R. 3W., Utah County, Utah.

The basal 410 feet of the measured section was taken at the top of the Great Blue Limestone on the first large bend in the road about one-half mile down Manning Canyon from the head, where the Great Blue Limestone is steeply east-dipping along the old railroad bed. The middle 390 feet was measured across the valley directly east and north of the main road across Mercur Pass. The upper 255 feet was measured on the small east-west ridge about one-fourth to one-half mile north of the Mercur Pass road.

The quartzite beds in unit 8 were used as markers to connect the basal and middle segments of the section. A dark gray-blue limestone of unit 18 was a marker bed to connect the middle and the upper parts of the section.

The Oquirrh-Manning Canyon contact was taken at the top of the pink-gray and purple-gray shales in the top of the Manning Canyon Shale, and the overlying sandy medium gray, brown weathering limestone containing fossil hash, cross-bedding, and crinoid stems(?). This upper contact is very similar to the one found in Soldier Canyon.

Limestanel Dark, gray, weathernow

Pennsylvanian

Oquirrh Formation s find abod at hebbed doubt bearing

Unit No. Description Thickness Above in feet Base

Limestone: Dark gray, weathers light gray and light gray with pink tinge in basal 12 ft., lithographic to finely crystalline, thin bedded in beds 1 to 10 in. thick in the top 14 ft. and 1/2 to 2 in. thick in the base, weathers platy with laminae on

Unit.	No.	Descri	ption	Thickness in feet	Feet Above Base
off lo no.	contains	ed surface in to bryozoans (? its, marker be	or crinoid	a writer for	iT
e i maira	ipe ifiction detailed so	Al benname bu e Canyon I be	distrib defat de la Soldre:	n th©appen Slanyen Sha	solani.
	gray, w local mo	ne: Dark gray eathers light b ottling of light	rown-gray w red-brown a	rith, com ya. nd	e kepali Rasyini Ta
	light ora	ange brown, fin	ne grained, t	thin	land 4
e idtni ba grunnstë b	contains	n beds 4 to 12 fossil hash .	deathman	16.68.64 onl	M J
iksti	Total	thickness mea	sured of Oqu	irrh Fm.	ind of
	on was take	nedeured secti	feet of the r	30 al	
		ne first large			
Mississip	pian-Penns	sylvanian	Manuing Can	mile dowd nestatie i's s	
Manning	Canyon Sha	ed across the	was niesew roselecze	that 308 slb	born od!
23	Shale (Ca	alcareous): Da	taga Hama'a	di no berba.	
			, weathers i	neulum	
	purple-g	ray and pink∞b	rown, conta	ins	
anid-v	light yell	.ow-brown and	medium ora	nge	
e and the	Contain	the base, siltens a few dark	y, calcareou gray limesto	nes	
	that weat	her medium ye	ellow-brown	and	
lo got es	light ora	nge-brown, ve	ry fine grain	ed, no o	
and garan	Limestor	ous, beds 1/2	to 6 in. thic	k, biral value	
	the basal	es become mo	re abundant	in	
insurbig	per cent	20 ft. making of the sedimen	up about 50	- 0	1064
ALCO A STATE	ound in Sol	i eno sol of in	ingara, Kapara	., 58	1064
22	Limeston	e: Dark gray,	weathers		
	medium a	and light brown	, very fine		
	grained,	thick bedded in	bed 3 ft.	10.42 (0.4	
	thick, con	ntains a few ve	rtical white	northerno	
Feet Above	calcite st	ringers 1/4 in	thick	: th 3	1006
Base	Shale: Da	ark gray, weat	hers to	ore since it is	
	calcareou	prown-gray soi s, slightly sil	I slope,	. 6	1003
	Limeston	e: Dark gray i	n basal 3 ft.		
ta egin	becoming	medium gray	in the top,		

weathers platy with lamine on

-51-

		FAC *	Feet
Un	it No.		Above
nord, 8		in feet	Base
		weathers medium and light brown-	
		gray, also medium brown, very	
		fine grained to finely crystalline,	
		thin bedded in beds 2 to 12 in. thick,	
		shaly in middle 4 ft., contains	
		brachiopods and crinoids in basal and the	
		3 ft. with a finely broken fossil hash	
		of mesters (Shalf) Medium brogot and ni	997
	21	Shale: Medium gray, weathers to	
		light brown-gray soil slope, sandy,	
		calcareous distributed beating 6	987
		1/4 to ! in. thick, weathers very platy 26	1500
		Limestone: (Sandy) Medium gray-	
		brown, weathers medium yellow-brown,	
		fine grained, thin bedded in beds 1/2	
		to 2 in. thick, weathers flaggy 4	981
		in bedd 2 to 10 in. thick, weathers	
		Shale: Medium gray, weathers to	
		medium brown-gray soil slope,	
		silty, calcareous	977
83		10 ft. from base in a 10 in, bed 21	
		Limestone: (Sandy) Medium gray,	
		weathers dark and light brown mottled,	
		fine grained, thin bedded in beds 6 to	
		12 in. thick, contains small amount of	
		mica, weathers blocky and slabby with	
		pock-like surface 5	966
0.8		2 cot	,
	20	Shale: (Covered) Dark and medium	
		gray, weathers to medium gray fine	
		grained sandy soil slope, calcareous,	
		contains fine quartz sand which makes	
		up about 20 per cent of the shale 85	961
		Ilaggy, contains a tew very thin	/
801	19	Limestone: Medium red-gray in basal	
		3 ft. becoming medium gray in middle	
		4 ft. and medium brown in top 3 ft.,	1
		weathers light and medium orange and	
		pink brown, very fine grained, thin	
		bedded in beds 1/2 to 2 in. thick,	
791		weathers platy 10	876
		Shale: Dark gray and black also share.	
		- 10 P. L. P. L. L. P. L.	

Est Limestolle: Dark gray, weathers shad shad median blue gray, we waters and the median blue gray, we water grained.

			it was in the	Feet
Unit No.		Description	Thickness	Above
Contrate and the contrate	urwi vii	ativitibns matieur	in feet	Base
		eous) Black and		
	dark gray, we	athers to medium	gray	
		oil slope containing	ga	
	few chips, ver	y calcareous and	n in Visite	
	carbonaceous		10	866
			Aller of the	
66		haly) Medium bro		
		3 ft. becoming me		
	그리는 것이 하는데 하는데 이 그렇게 되어서요 선택하였다.	eathers medium y		<u> </u>
		t orange-brown, f		
		nated bedding in be		
	1/4 to 1 in. th	ick, weathers ver	y platy 26	856
	vetal	(Sandy) Medawn	Linesion	
18	Limestone: D	ark gray, weather	s w nword	
	medium and li	ght blue-gray, ver	y fine	
v.p	grained to lith	ographic, thin bed	ded	
	in beds 2 to 10	in. thick, weathe	rs	
	blocky and sla	bby due to good pa	rallel	
	jointing, lamir	nae on weathered s	surface,	
	그렇지 않는 아이들이 얼마나 되었다. 그 아이들이 얼마나 되었다는 그 없는 것은 바다 모든 것으로 했다.	taining styolites a		
	10 ft. from ba	se in a 10 in. bed	21	830
		es (Sandy) Medium		
		ark gray and black	맛있다. 그 도마일 및 그 및 것으로 다른 VX	
		gray, fine graine	AND AND ADDRESS OF THE PARTY OF	
		beds 1 to 6 in. thi		
		rough and irregul		
		ins few styolites i	n	0.00
	top 2 in		3	809
	medium	overed Dark and		
		ark gray, weather		
		blue-gray, very fi		
		ographic, thin bed		
		4 in. thick, weath	ersods qui	
		ns a few very thin		806
	calcite stringe	rs rg-bar mulbaM	aholesmi <mark>8</mark>	800
17	Shale Dark a	Caray medican terras	t. beco	
11		reen and dark gre		
		s medium green-g	ray	
		reen-brown chip	http://www.	
		lope, clayey, slig		798
178	carcareous .	n jak yele	10	
16	Limestone: D	ark gray, weather	s	
10		gray, weather gray, very fine gr		
	incarain blue-	gray, very rine gr	arrica,	

	Ac.		
Unit No	Noderical and		Feet
SVOOT AND	Description	Thickness	Above
easd re		in feet	Base
	thin bedded in beds 1 to 12 in.		
	thick, weathers flaggy, contains	310-11	
	numerous horizontal and vertical		
	white calcite stringers 1/8 to	LE LAND	
	1 1/2 in. thick, styolites displace	at Æf	
670	calcite stringers.	8	782
15	Shale: Dark green and green-black,		
	weathers medium green and medium	1544.20 : 11 i	
	green-gray, silty, non-calcareous,	₹525	
	contains abundant chips in the soil		
		ibenii	55.
	nedium red onay that		774
14	() 110 mtmato: Mandiana 1	téez	
	weathers light medium brown-gray	Jaio	
	with a green tinge, very fine graine	d.	
	thin to thick bedded in beds 2 in. to	AR COLLEGE	
	8 ft. thick, average thickness 2 to 1	2	
	in., weathers slabby with scintillati	ng.	
	surface, medium red-brown iron sp		
	present in a few zones on the weather	loiches	
	surface	red	2013
	surface bening and reversional	. \ 30	758
13	Shale: Dark green and green-black,	VELY	
15	Shale: Dark green and green-black,	agns	
	weathers to medium and light green	Arcj.	
570	brown chip covered soil slope, silty	5 073.8 cl	
	locally very carbonaceous, contains		
	interbedded quartzite in beds 1 to 10	า รมสน	
	ft. thick, dark green becoming medi	.um-	
	brown-green on the weathered surfa		
	fine grained		728
	markey soil slope centaining	word	multi-safili
12	Limestone: Dark-medium brown-gr	av	
	in basal 8 ft. becoming dark gray in		
	top, weathers medium brown, fine	iJibJ	
	grained, thin bedded in beds 2 to 12	73.52	- 75
	in. thick, weathers flaggy and blocky	Teus	
	Contains white calcite stringers 1/4	ed	
	contains white calcite stringers 1/4	Shall	318 2
784	to 1/2 in. thick, very fossiliferous	npou	
	and hashy, brachiopods and pelecypo		
	predominate		681
11	Shale: Dark gray and black, weather	rs	
	to light brown-gray and medium gray	· MILL	
	here flaggy, scintillating surface	J.59W	

	-54-	Feet
Jnit N	o. Description Thickness	Above
	chip covered soil slope, silty,	Base
	non-calcareous, contains few interbedded black-brown	
	weathering limestone beds 6 to 12 in. thick with abundant	
	fenestellid bryozoans . 1986 . 100	670
10	Quartzite: (Argillaceous) Light gray, weathers light brown-gray in the top 4 ft. becoming dark gray, medium gray weathering in the	
	underlying 6 ft. The next underlying 2 ft. is medium red-gray that	
	weathers dark red-gray and light pink-gray, the basal 6 ft. is light-	
	medium gray, weathers light and medium yellow-brown contains argillaceous spots and blobs which cause the weathered surface to be	
	pock-like, blobs are medium yellow and medium brown about 1/8 in. by	
	1/2 in. Unit is very fine grained, very dense and hard, weathers to subrounded blocks, contains	
	Archimedes, Kaskia, crinoids and brachiopod fragments	570
9	Shale: (Covered) Medium green, dark green, medium green-brown, medium brown-gray and black,	
	weathers to medium brown and light brown-gray soil slope containing	
	chips, silty and some clayey, non- calcareous, darker shales very carbonaceous. Base contains brown	
	shales, middle contains green shales, and top is black. Middle green	
	shales contain 4 to 6 in. black siltstone nodules that weather light brown 42	552
8	Quartzite: Dark gray, weathers medium green-brown and medium green-gray, fine grained, top 2 ft. massive, basal 2 ft. thin bedded in beds 1 to 2 in. thick,	
	/ II Thin hedged in hode I to / in thick	

		-53		Feet
Uni		44 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 -	Thickness	
		De, religion	in feet	Base
E 7 EL		Shale: Medium gray, weathers to		
		light brown-gray very chippy slope,		
		silty, non-calcareous		406
		and the continuous state of the		
		Quartzite: Dark gray, weathers		
		medium green-brown and light green		
08.		gray, fine grained, thin to medium		
		bedded in beds 4 to 24 in. thick,		
		weathers slabby, scintillating surfac		
		contains few white quartz veins		400
		and begin ift massive, upperson		
	7	Shale: Medium brown-gray and medi	lum	
		gray, weathers to light brown-gray s		
		slope containing fine shale chips,		
		calcareous and slightly silty, contain	is of	
		interbedded limestones which are dan	rk	
		gray and medium brown-gray, which	eri.I	
		weather light and medium orange-bro	own,	
3.0		fine grained, beds 6 to 12 in. thick,	very	
		fossiliferous containing Composita,	3	
		Spirifer, Archimedes, Fenestella, co	orals,	
		crinoids, and trilobites	55.52	391
28		Lili, thick apidt .th L		o de t
	6	Limestone: Dark blue-gray, weather		
		light brown-gray, very fine grained,		
		basal 2 ft. thin bedded in beds 1/2 to		
		2 in. thick, overlying 12 ft. laminate		
		bedding 1/16 to 1/2 in. thick, weathe	rsw	
		platy and flaggy, contains a few white	.	
		calcite stringers 1/16 to 1/2 in. thick	k. 14	339
		t brown althoughphic, sidal boddeds ig		337
àS	5	Shale: Medium red-gray, weathers t		
		light pink-gray fine grained soil slope	a .	
		calcareous, slightly silty	Lad27 F	325
		weather a to wery old his guarands aw	TE TO	323
		Limestone: Light gray-brown, weath	ers	
36		mottled light brown and light gray, ve		1.7511
		fine grained, laminated bedding in be-	da	124
		1/4 to 1 in. thick, weathers platy	or II	210
		the control of the key of the control of the contro		318
23		Shale: (Calcareous) Light gray, ben	Form Profile	
U 101		weathers about the same, very	1b1g	
		calcareous, slightly silty	1-420	215
				317
		there medium brown gray with fine a in the soil slope, clayey, pink		
			of at	
			ALL AND	

42.57				Feet
Unit No	Sealer Messelle	Description	Thickness	Above
44 10 144			in feet	Base
	weathers light abundant carbo	medium brown- brown-gray, con naceous materia	tains l,	
		ains a few interb		
	dark brown wea	of medium gray- athering fine grain light brown in ba	ined	
	becoming black silty, weathers	in top 3 to 4 ft. to fine grained	, very	
	slope	jpy while the und	70	234
2	weathers medic	lium brown-gray ım yellow-brown edded in bed 1/2	, fine	
The Park	in. thick, weat		· · · · · 3	164
		007.11	. 1325	201
	light brown=gra silty, weathers	brown-gray, we y and medium br to chippy soil sl lecypod fragmen	own, lope,	
2 20		2 in. beds of dan		
	siltstone.	Some (Borevet	y• k•• ₉ • 5	161
85 8	weathers mediu grained, thin be	ium brown-gray m yellow-brown edded in beds 1/2	, fine 2 to 6	
6	white calcite s	ners platy and fla	Firm an interest and an address about	156
	gray, weathers	gray and medium to light brown-g	ray soil	Great
	slope, slightly	silty	27	151
	weathers light h	ium brown-gray, prown-gray, fine ng in beds 1/2 to	grained.	.144
	thick, weathers	platy	4	124
foldler Cany 4 W., Too	gray and light b friable, clayey,	gray, weathers rown, very fissi weathers to fine	light and based a	da life de la composición della composición dell
de and the	grained soft soil	slope	Blue Lithie w.	120
	alcareous	a age the serie	novemio 4 65	

Feet

rt A la diamen	Design of the second of the se	nit No.	reet
Unit No.	Description	Thickness	Above
	The second transfer of Mr. St. St.	in feet	Base
	Limestone: (Argillaceous) Medium gray, weathers medium yellow-brow fine grained, contains light brown argillaceous partings, thin bedded i beds 2 to 4 in., weathers to rounded slabs	wn, n	111
	1991 N. 1991 N	5.500	
	Shale: (Covered) Dark and medium gray, weathers medium gray and lig	sa ht	
	yellow-gray, silty, the top 25 ft. is	mole i	
	very chippy while the underlying 15		
	20 ft. weathers in splinter-like chip contains few interbedded medium an	d	
	light gray, medium brown weathering	ıg,	
42	fine grained quartzite beds 6 to 18 i.	n.	
	thick	. 80	110
	Limestone: Medium brown, weathe medium orange-brown, very fine	rs rs	
	grained, massive, few white calcite	g silty	
	stringers	5 tr >> 2	30
	y iew i to 2 in, bed o day weigh	a ver	30
	Shale: (Covered) Black and medium	ailta i	
	gray in weathered slope, slightly si		
	non-calcareous	28	28
	iers inclium yellow-lidd a line	weat	

Fault (?) contact vessel and a red sew , south .mt:

Great Blue Limestone (top)

Limestone: Medium blue-gray, weathers light gray, very fine grained, thin bedded in beds 1 to 12 in. thick, weathers flaggy and platy, cliff former.

Ophir Pass. The Ophir Pass section is located on the east-west trending divide between the head of the South Fork of Soldier Canyon and the head of Ophir Canyon in Sec. 11, T. 6S., R. 4 W., Tooele County, Utah. The Manning Canyon Shale forms the divide. The Great Blue Limestone forms the ridge west of the divide and the Oquirrh Formation forms the ridge east of the divide.

The Great Blue-Manning Canyon contact is at the top of the steeply east dipping medium gray, light gray weathering limestone slope about 2000 to 3000 feet west of the log cabin which is located 200 to 300 feet down the canyon from the head of Ophir Canyon.

The Manning Canyon-Oquirrh contact was taken where the medium gray shales and interbedded limestones in the upper part of the Manning Canyon Shale are overlain by the sandy medium gray, brown weathering limestones of the basal Oquirrh Formation.

Conformable contact at 1 and at beabed and apparent

Mississippian and Pennsylvanian

Manning Canyon Shale

	0	Black, weathers to medium grays.	Skale	Feet
Un	it I	No. Description Description	Thickness	Above
	2023	Make the state of the second section is sufficient to the second section.	in feet	Base
	17	Shale (Covered): Medium gray, weathers to light brown-gray soil s calcareous and silty, contains interbedded medium gray, medium	lope,	
932		brown-gray weathering, fine graine	dilagu	
		limestone in beds 6 to 24 in. thick		1347
922	16	weathers light and medium brown, grained, sandy, thin bedded in beds	fine 4 to	
		12 in. thick, weathers slabby, cont		1.
		white calcite stringers, also calcite coating on weathered surface	o mediur	1154
	15		athers m ltv. 19	1144
845	14			
		grained, thin bedded in beds 2 to 12	in. Shala	616
		thick, weathers slabby, very sandy	on bas	
		weathered surface	29	1125
	13	Shale (Covered): Medium gray, we to medium gray soil slope, silty and	athers	.42
831		calcareous ! . shed elizitate pena	87	1096

Feet bove Base
bove Base
3ase 0 009
009
.009
.009
.009
005
997
948
932
922
845
019
4 100
831

1501						Feet
Unit No.	9 S GLENIA LEI 1	Des	scription	Tł	nickness	Above
2 92-1-4		re - ve norsa kuf	the mark (d)		in feet	Base
8			weathers			
			ay, fine gra			
			2 to 6 in.	thick,	M amad	
	weathers s		ontains		w. to i way z	0.726
	Wewokella	n. sp.	alc . Algur	cyfnid a	5	796
	Quartzite:	Dark gr	reen-gray i	n the bas	sal	
			edium gray			
			50 ft. and r			
	gray in the	top 10 f	t., weather	rs mediu	m	
			ery fine gra		1901	
	dense and	hard, th	in and medi	um bedd	.ed	
	in beds li	n. to 3 f	t. thick, we	eathers		
	flaggy and	slabby,	contains cr	oss-bed	ding	
ar a	locally .		ikag Pala	의 22 의 명시나 나는 14 - 1873	69	791
	Shale: Day	uls amoon	Area Tari			
			-gray, and medium gr		r) searcil	
			soil slope,			
	silty	nword in	18 ALMOLD W	mangma b	light an	722
	sirty	89 S ((e)	ed at babb	Walter !	benist.	122
	Quartzite:	Medium	gray and r	nedium	a van Sa	
			ers medium		Til treot	
			ne grained,		19117 वृद	
	bedded in h	oeds 1 to	8 in. thick	nerings a	Amplex	
	weathers f		inga, wéd. Paga, wéd.	litt Bil	10 fragmet	710
7	Shale: Me	dium gra	ıy and medi	um brow	⁄n =	
	gray, weat	hers to 1	ight and me	edium	Shale	
			ained soil s		peconic	
			ecially in to		n ediai	700
	0	cr der at	J.H. AWOIG	mediam		
	Limestone:	Mediu	n brown-gr	ay,	isw	
mi é an more	weathers li	ight brov	vn, fine gra	ained to	THILDSAI	
	finely crys	talline,	thin bedded	in beds	91.601.60	
	1/2 to 1 in	. thick,	weathers fl	aggy,	o this in	
	contains a	few Spir	ifer, bryoz	oan and	100010	
	crinoid fra	gments	LKINY AF	atherea	aelmoon	645
	light gray :	and Lylif	pink gray,	very in	13	
			wn gray, v		Basmil	. 4
			soil slope,	eather	R.F. B.Y. W	1 1 10
	calcareous	and slig	ntly silty	ind grain	brown	642

				Feet
Unit No.	De	escription	Thickness	Above
	Easton III.	MODE DESCRIPTION	in feet	Base
9814	Limestone (Argill brown-gray, weat gray, fine grained beds 1/2 to 2 in.	hers light brown- l, thin bedded in		3
	argillaceous, wea contains <u>Dictyoclo</u> <u>Fenestella</u> , <u>Archi</u>	thers flaggy, stus, Spirifer, medes, trilobite		0
	fragments, and cr Shale: Black and weathers to mediu gray soil slope co chips, clayey, cal	dark green-gray, im greenish brow ntaining a few		634
	fossiliferous in ba containing Archim	sal 5 to 10 feet	in bode i re Naggy and Jesselv	
	Composita, Spirif	er, and crinoids	26	633
6	Limestone: Dark and medium brown	to medium gray	Shalo Dur black vea	
	light and medium grained, thin bedd	brown-gray, fine led in beds 1/2 to	annem race	
	12 in. thick, weat fossiliferous, con Spirifer, Fenestel Amplexizaphrentis Composita, trilob	tains, <u>Dictyoclos</u> la, <u>Polypora,</u> s, Talpaspongia,		
0.14	fragments	ite and Crinoid	21	607
5	Shale: Medium gr becoming black an	d dark green-gra	ıy	
1007	in the middle of the gray to medium by ft., weathers to do medium green-brocalcareous and cla	own in the top 15 ark gray and own chippy soil s ayey in base beco	lope,	
	clayey and non-cal middle of the unit 100 to 150 ft., con	and silty in the to		
Cld	nodules about 1 to			586
4	Limestone: Dark gray, weathers lig	and medium blue		
14-3	brown, fine graine		8 11 2 9 T G D T T	

		Feet
Unit No. Description	Chickness	Above
[2] 2 전 1 전 1 전 2 전 1 전 1 전 1 전 2 전 2 전 2 전	in feet	Base
cross-bedded and irregular bedded		to be the
in beds 1/2 to 3 in. thick, weathers		
flaggy, sandy in top, contains white		of Evidian
calcite stringers, hashy in basal		
one half	6	192
	. 0	192
3 Shale (Covered): Medium gray,	Office of	
weathers to light gray and light		
	ad, des	
brown-gray, fine grained soil slope	. 115	186
2 Limestone: Medium blue-gray in	erafera.	. The Later
The state of the s		
base becoming medium and light		
brown-gray in top, weathers medium	Latings of	d dans a les
and light gray with some brown,		
medium and coarse grained in basal		1 1 2 5 1 2 0
5 ft., fine grained throughout the res	t	d appropri
of the unit, medium bedded in beds 1		12
to 1 1/2 ft. thick, weathers slabby		D Burners
and blocky, contains intraformational	i e	
limestone pebble conglomerate in the		
basal 5 ft., arenaceous in the top 5		
to 10 ft., contains brachiopod, coral,		
pelecypod, and crinoid fragments .	21	ا 7 1
The contract of the contract o	និងថ្មីទីវិសម្រប	
l Shale: Medium to dark gray with	B-mward	
medium brown partings, weathers lig	April 1	
pink-gray in base and light gray in the		
upper 45 ft., slightly calcareous and	ribean of	
slightly silty	5.0	5.0
nei Medium brown and medium	. 50	50
r.y, weathers light Egymand	y-nword	
Conformable contact		
d tedding 1,4 to 1 in thirty		
Mississippian		
Wississippian		
C 71	inteam: 1	
Great Blue Limestone (top)	aro itrax	
Limestone: Medium gray, weathers	L'Ac at	1015
light gray and light pink-gray, very fi	ne	42
grained, thin bedded in beds 1/2 to 2	in.	
thick, weathers flaggy and platy,	CIROLAL	
contains black chert and fossil hash.	dark gra	
TO bin lints, weathers there	in beds	
a file while calcite a tangers		
Ducks 1239 1	I ALL SAL	

West Canyon: The formation was measured in two segments. The lower 1283 feet was on the north slope of Iron Canyon about one to one and one fourth miles up the canyon from the mouth. The upper 400 foot interval was measured in West Canyon about one-fourth to one-half mile northwest of Farmers Flat. Both of these localities occur on the north plunging nose of the Long Ridge Anticline, Sec. 3, T. 5 S., R. 3 W., Utah County, Utah.

The Great Blue-Manning Canyon contact is exposed in Iron Canyon on the north side of the stream bed where the medium gray, light gray weathering platy limestones are overlain by black shales. The Manning Canyon-Oquirrh contact is arbitrary because the upper 400 feet of the formation is covered.

	nichain an	adiron	Torrest.	ab ye	ord.
Conformable contact	(covered)	d estiba	il. Lw		bas.

	and fight go de with some bedying the	
Mississinni	ian-Pennsylvanian ongge in beet in the first description of description of the contract of the	
Wilsbibbippi	bit. The grained throughout the control of the control of the	
Manning Ca	nyon Shale	
	Description of the control of the co	Feet
Unit No.	Description Thickness	Above
	in feet	Base
16	Shale (Covered): Medium gray, weathers to light gray and light	
7 98 90	brown-gray soil slope, contains	
	interbedded medium gray limestone	
	that weathers medium and light gray	
	to medium and light brown 400	1683
	Appreciate of the state of the	
15	Limestone: Medium brown and medium	
	brown-gray, weathers light brown and	
2 8 8 9	light orange-brown, fine grained,	
	laminated bedding 1/4 to 1 in. thick,	
	weathers platy	1283
	Limestone: Medium and light gray,	
	weathers very light gray, fine grained,	
	laminated and thin bedded in beds 1/4	
	to 2 in. thick, weathers platy 31	1270
	ted thin bedund in bedy 142 to line	
	Limestone: Dark gray to black, weathers	
	dark gray, very fine grained, thin bedded in beds 2 to 6 in. thick, weathers flaggy,	
	contains a few white calcite stringers	
	1/2 in. thick	1239

10104h	Sample Continue Language of Trickman	Feet	
Unit No.	- correction Thickness	Above	
	in feet	Base	
	Limestone: Dark gray, weathers light gray, fine grained to lithographic, thin bedded in beds 1 to 12 in. thick, weathers flaggy and slabby, contains a		
	few white calcite stringers	1236	
14	Shale: (Covered): Black and dark green, weathers to medium brown-green and medium green-gray soil slope, silty, non-calcareous, contains a few beds of dark green, fine grained quartzite that		
	weathers medium brown-green 127	1178	
13	Limestone: Dark gray, weathers light		
	gray, very fine grained to lithographic, laminated and thin bedded in beds 1/2 to 3 in. thick, very argillaceous in basal 3 ft., weathers platy, contains a few white calcite stringers 1/2 to 1 in.		
	Shale: Black to dark green, weathers	1051	
828	dark and medium green-gray chippy soil slope, silty, contains a few siltstone	1041	
	artigoria matt. The public of ere in the	1041	
	Limestone: Dark-medium gray, weathers dark gray, fine grained, laminated bedding in beds 1/8 to 1/4 in. thick, weathers platy, contains a few horizontal and vertical white	*17.	
787	calcite stringers 10	1023	
	Shale: Dark brown and medium green- gray, weathers to medium and light green-brown soil slope, silty and		
	slightly calcareous	1013	
	Quartzite: Dark and medium gray-green, weathers medium brown-green and		
	green-gray, thin bedded in beds 1 to 12 in thick, weathers slabby, contains plant fragments 24	000	
		999	

11 27		Feet
Init No.	Description Thickness	Above
svede-	in feet	Base
10	Shald (Cavarad in base). Black and	
10	Shale (Covered in base): Black and	
= 0.2	dark green, weathers to medium brown -	
	green and medium green-gray soil	
	slope with fine chips, slightly silty 102	975
5.86 1	Limestone: Black, weathers dark	
	gray with medium orange-brown	
	stains on surface, very fine grained	
	to finely crystalline, thin bedded in	
	beds 1/2 to 12 in. thick, weathers	
	flaggy	873
878	The state of the s	
0	Quartzite: Light gray, weathers light	
	gray, fine grained to finely crystalline,	
	scintillating, thin bedded in beds 6 to 24	
	in. thick, weathers slabby 22	860
	Quartzite: Light gray, medium green-	
	gray, medium gray and dark gray,	
	weathers medium brown-gray, medium	
1991	brown-green and medium brown,	
	medium and fine grained, medium	
	bedded in beds 1 to 2 ft. thick becoming	
	thin bedded in the top 15 ft., cross-	
	bedded, weathers slabby and blocky . 51	838
1401	bedded, weathers stabby and blocky . 51	030
7	Limestone (Sandy and argillaceous):	
	Medium gray and medium blue-gray,	
	weathers to light blue-gray and light to	
	medium brown, very fine grained to	
	finely crystalline, thin bedded in beds	
	1 to 6 in. thick, weathers flaggy and	
1023	rounded due to argillaceous content. 9	787
	Tounded due to arginaceous content.	101
	Shale: Dark brown and dark gray	
	becoming black in basal 15 ft., weathers	
	to medium brown-gray soil slope, slightly	
£101	silty, calcareous in upper one half,	
	contains a few thin beds of medium brown-	
Y.	gray quartzite that weathers medium brown,	
	pelecypod fragments occur in the basal	
	one half	778

1991		-01-		
Unit No.		Description	Thickness in feet	Feet Above Base
6	becomes me 4 in., weath fine grainted 2 to 6 in. the Dictyoclostu	Dark medium gray dium brown in top 2 there light gray, very d, thin bedded in beds ick, contains as, brachiopod spines, ssil hash	love police	710
	Limestone: weathers lig yellow tinge in beds 1 to	Medium to light brow the brown with pink an fine grained, thin be 4 in. thick, weathers	n, d edded	
		laty, contains crinoid and brachiopod fragi		708
	weathers liggray, fine gand medium thick, contained thick	Dark and medium gracht gray and light brown rained to lithographic bedded in beds 1 to 2 ins Dictyoclostus, Cona, Spirifer, Fenestell alpaspongia, horn constant above to about 4 ft. above to ational limestone pebles, the pebbles are an /2 in. in diameter.	thin thin in. mposita, a, cals, bed	702
e an den bes Reden seen Bus tie Wes Artins tie St	Limestone: light gray, f in beds 6 to	Medium gray, weather ine grained, thin bedo 12 in. thick, weather ains Dictyoclostus, Sp	Shair Cove to Tight bell alope cons	
ondo::::	Shale: Black weathers to	and ccrinoids	5 Shale: Blace to medium: V	677
e l'echio.	nodules, silt interpedded fossiliferous dark and me	ins a few black siltstory and calcareous, medium brown fine grass limestones that weat dium brown occur in hick. The shales continuous	m beniarg rained track her historia beds	est Vova Base
		very and wright friend		

Jnit No.	Description Thickness	Feet ss Above
	in feet	
assa.	tal ni	
	Wewokella 2 spp. and the limestones	
	contain Archimedes, Fenestella, Polypora, Rhombopora (?), Kaskia,	
	Spirifer, Dictyoclostus, Composita,	
	horn corals, crinoids, a few	r The
	pelecypods 46	612
n rev	N. C.	012
4	Limestone: Medium brown, weathers	
	medium orange-brown, fine grained,	
	thin bedded in beds 2 to 6 in, thick,	
	weathers flaggy, vertical white calcite	
	bringers 1/2 to 1 in. thick, contains	
	Dicty octosius and other lossif	
	fragments	300
	Limestone: Dark gray becomes medium red-gray in top 10 ft., weathers medium gray becomes light pink-gray in top 50 ft., fine grained to lithographic, thin bedded in basal 20 ft. becomes massive in middle 32 ft., and thin bedded in top 53 ft., weathers flaggy and slabby, contains white calcite stringers in the middle of the sub-unit and becomes fossiliferous in the top 10 ft., brachiopod fragments and crinoids are abundant in the fossiliferous part 106	
3	Shale (Covered): Dark gray in base becomes light gray in top, weathers to light gray and light red-gray soil slope, contains black siltstone nodules	450
773	and pelecypod fragments 365	458
2	Shale: Black and dark brown, weathers to medium brown-gray soil slope, calcareous and slightly silty, contains interbedded medium brown, fine grained, medium orange-brown weathering limestone in beds 6 to 24	
	in. thick	93
	The bar in the control of million but	

orange all the states contain

		-07-69-		
Feet		* 3 X **		Feet
Unit No.	Thickness	Description (Thickness	Above
Buse	tosa ki		in feet	Base
1	Shale: Medi	ium red-gray in basa	1.6	4 00 00 00 00 00 00 00 00 00 00 00 00 00
1913	ft. becomes medium bro	black in top, weathe wn-gray soil slope, careous, contains bla	rs to silty,	
	calcareous	siltstone nodules in b	asal	
	one-half	ys and medium brook rained becomes mad		63
Conformabl	le contact	ase and very fire at a bedded in beds 2 to i	nida jast si	
Mississippi		- bedder especially in		
intar.				
Great Blue	Limestone (1	top)		
	Figure atach	Medium brown, weal		
		Dark gray, weather		
		ht brown gray, fine g		
1785		in beds 1 to 12 in. th		
	weathers fla	aggy and platy, ledge	former.	
	Scownin Apr	Medium gray, becom	Sandstange	
	angdica. #T	-gray in top, weathe	in edittri rad	
		1910 foot section on		
about 1000	to 2000 feet i	east to west, on the north and west of the Utah.	spring, Sec.	
	was the TAG	kes up about 10 to 60	Ent tenable	
gray, light	Great Blue-M gray weather	lanning Canyon conta ring, cherty limestor ing Canyon-Oquirrh	ct is taken wh ne are overlai	n by medium

ark um gray shales. The Manning Canyon-Oquirrh contact is not readily located because of soil cover. The contact was arbitrarily taken beneath the dominant medium gray, medium brown weathering quartzite of the Oquirrh Formation.

Conformable contact (covered) and institution was a

Mississippian Pennsylvanian of A abod of habbad muibem

Manning Canyon Shale also patrid was a spring control of the contr

Unit No. Description Thicknes in feet

17 Shale (Covered): Medium gray, Feet Thickness Above Base

weathers to light gray and light

Feet				
Unit No.		That was the	Thirkbooks	Feet
	Josi Wi	Description	Thickness	
7 S 40 C 4			in feet	Base
	•		3.7	
		oil slope, silty,		
		eous in base		1910
• •	스크 레이트 이 기가 있다면 스크리 유모를 보다	edale lies van - rom	4.7	
16		gillaceous): Medium		
- exception		athers light gray,		
C.M. T. SERVIN		and medium brown-		
		ined becomes mediu		
		e and very fine grain		
¥.		dded in beds 2 to 12	in, skroveka	s mrije krol
		pedded especially in		
	시대를 잃었다. 그들은 하나 사람이 얼굴하다 가는 만든 장면 있어?	, weathers slabby an		gia a
	blocky	h i i i h i i i i i i i i i i i i i i i	25	1810
	The second secon	(qoi)	è l'imestorie	ul
	Quartzite: Me	edium brown, weathe	rs	
		ine to medium graine		
	thin bedded in	beds 4 to 12 in. thic	kjais vata	
		weathers slabby		1785
		tlangy and platy, led		
	Sandstone: Me	edium gray, becomes	5	
		ray in top, weathers		
ate (190 de	medium to dar	k gray and dark red	Mountain:	
Seep Ca	gray, medium	grained, thin bedded	was measu	
	in beds 2 to 12	in. thick, weathers	Ø to 2000 See	
	slabby, arkosi	ic in middle 10 ft.,	Itah County	
	feldspar make	s up about 30 to 40 p	er	
		kə.açvasə şa.m.sM-		1782
dn by me	tone are everl	nering, cherty limber	tray weat	
value t	Shale: Varieg	ated medium gray an	et . The Mb.	
lly taken	medium brown	, weathers light gra	you to senting	
		n-gray, calcareous,		
			TimpO 2	1762
	sepe, 770 ser,			
	Limestone: D	ark gray to dark blue	36	
	gray, weather	s light brown-gray a	ale manacha	
	light gray, ver	ry fine grained, thin	and	
	[전기 시항 : [10] [2] [12] [2] [2] [2] [2] [2] [2] [2] [2] [2] [d in beds 2 to 24 in.		
	thick, weather	s slabby with rough		
		ins a few white calci	te de la compa	
Jee"i	stringers	on stanger order	37	1760
svodA 6	THICKNE	Description		
0814	Quartzite: Lig	ght gray, weathers li	ght	
		light brown-gray, ve		
		nedium bedded in be		
		Mark 11일 등 시민 100 100 100 100 100 100 100 100 100 10		

210		- / L=			
Fee				Feet	
Jnit	No.	Description	Thickness	Above	
Bass	in test		in feet		
			III leet	Base	
	1 to 3 ft 3 3	weathers slabby and		X 1 2 1 2	
			I the state of		
	rounded, so	cintillating in base	23	1723	
		the first of the f			
	Quartzite:	Medium gray and dark	gray,		
	weathers m	edium and light brown-	gray		
	with a green	n tinge, very fine grain	ed,		
	thin bedded	in beds 1/2 to 6 in. thi	ck,		
	weathers fla	aggy, cross-bedded .	29	1700	
	1 1 1,50	light pink-gray soil slo	bas ymg	1100	
	Quartzite:	Medium gray and light	draw (6:		
	mottled, we	eathers dark brown and	bad Al S		
*		y, fine grained, thin			
881		eds 2 to 6 in. thick, wea			
	maggy, cros	ss-bedded	7	1671	
	1 10 1 1	of Jan 1 in the second	21219		
		red): Medium gray and			
	medium gre	en-gray, weathers to m	nedium		
	and light ye.	llow-brown and medium	, benisty		
FC.1	brown-gray	, silty, contains interb	edded		
	medium gra	y and light gray fine gr	ained.		
	medium bro	wn weathering quartzite	e heds		
	6 to 12 in. t	hick ben weller-men	Caurl32ca	1664	
	I Ohrstandgl	redium, red-gray and I	n aword	1004	
		Medium and light gray,			
	weathers me	adium brown - 1°	and light		
15.5		edium brown, medium g			
	hada 2 t- 12	ine grained, thin bedde	d in		
	beds 2 to 12	in. thick, weathers sla	abby		
		, contains medium brow			
	argillaceous	specks in top one-half			
		s brown argillaceous	nn bede 4		
entre v	laminae band	ding in basal one-half		1632	
152	0	cite stringers	white cal		
13	Shale: Medi	um green-gray, light g	rav.		
	and light yel	low-brown, weathers to	n light	6	
	and medium	brown soil slope, well	yellow-b	-1	
	foliated, cla	yey, slightly calcareou	AY LAYET A		
	locally sligh	tly silty, contains a fev	gray, we		
152	interhedded	calcareous siltstone be	and medi		
	1/2 to 2 in	thick, Alethopteris,	us		
	Neurontoria	Condition 1	the state of		
	plant for	Cordaites, and other		1289	
	piant fragme		weathers		
	yellow-brow		. 14	1602	
red r	919	to 12 in, thick, weather			
147	V		laggy.		

T		Feet
Unit No.	Description Thickness	Above
vouet es	in feet	Base
12	Limestone: Medium gray, weathers	
1.7.7 1.7.9.1	light and medium gray, very fine	1000
	grained, thin bedded in beds 1/2 to	
	2 in. thick, weathers slabby and	
	blocky, contains white calcite contains white	
	stringers	1588
	thin bedded a beds 1/2 to 6 in thick.	
01/31/1	Shale: Medium gray, weathers to light	
	gray and light pink-gray soil slope,	
	2 ft. bed of dark gray, fine grained,	
	medium gray weathering limestone that	
	has laminated bedding in beds 1/2 in.	
	thick	1580
	And the first of t	
10	Limestone: Medium and dark blue and limestone	
	gray, weathers light gray, very fine	
	grained, thin bedded in beds 6 to 12 in.	
	thick, weathers flaggy and platy	1548
	medium gray and near gray fine grained,	
	Shale: Variegated medium yellow, muibem	
1-101	medium brown-yellow, medium green-	
	brown, medium red-gray, and light	
	gray, weathers to light yellow-brown and light brown-gray soil slope,	
	slightly calcareous	1539
	·	
	Limestone: Dark gray, weathers light	
	gray, very fine grained, thin bedded	
	in beds 4 to 6 in. thick, weathers	
-1632	platy and flaggy, contains very thin	
	white calcite stringers 6	1527
	Shale: Medium green gray fight dray	
9	Shale: Variegated medium and light	
	yellow-brown, medium and light red-	
	gray, very light gray, and light yellow-	
	gray, weathers to light yellow-brown,	1521
	and medium brown soil slope 48	1341
	Quartzite: Medium gray-brown,	
	weathers light brown and medium	
	brown-gray, fine grained, thin bedded	
1602	in beds 6 to 12 in. thick, weathers	
	flaggy 9	1473
	10.00 (1995)	

Feet		- 13-		
Unit No	Thickness in fect	Description	Thickness in feet	Feet Above Base
	61. 1 5 ^[X]	aci, wealingth to medica		
		irk green, medium green		
		nd green-black, weathers		
		n green-brown soil slope		
		fine grained dark green		
		and siltstone beds 1/2 to		
		k, pelecypod fragments		
	30 ft	pecially in the basal 20 to	e pelite	- 12
0011	JO 11		65	1464
	Ouartzite	: Medium gray and medi		
		becomes medium green-		
		., weathers light gray as		
PELL .		gray, fine grained to fin		
	crystallin	e, medium bedded in bed	ery	
	1 to 2 ft.	thick, scintillating, cont	Sinc	
		rown argillaceous partin		1200
	SHAT AND	20 m arginaccous partin	gs · wood	1399
	Shale: Me	edium green and medium	tiet bas	
		ay, weathers about the sa		
	silty, cont	tains pelecypods in basal	Barris	
25014	l to 2 ft.		Maria de la Compaño de la Comp	1393
	ink requ	B William Control		13/3
	Quartzite:	Medium gray, weather	ollaise s	
	light and r	medium brown, very fine	My daery	
1.10	grained, t	thin bedded in beds 4 to 1	2 1 1031	12.5
941		weathers flaggy	4	1383
			1	, ,
aforc 8 and	Shale: Da	rk green to black, weath	ers	*
		n green-brown soil slope		
<u>Charley</u>	wilty, con	tains a few pelecypod	7 -	
	fragments	come very coarse grant it of sub-unit, thin to t	90	1379
		The fill metal to the file and fill to be and	out brother and	
7	Quartzite:	Light and medium gray	1 nr aped	
	weathers 1	light brown and light grav	Y.,	
	fine to me	dium grained, thin and n	nedium	
758	bedded in	beds 2 to 36 in. thick, w	eathers	
	slabby and	l blocky, becomes more	B	
	argillaceo	us in top 5 ft. and in bas	al	
	13 ft., als	so in the top and the base	the	
333	unit is dar	ker brown	118	1289
		Light to medium gray,	Ouartzite	
	bres	light and medium brown		

	-85-78		Feet
nit No.	Description	Thickness in feet	Above Base
, S	feet and the second	111 1000	Dasc
	Shale: Black, weathers to medium gray soil slope containing fine chip silty, and non-calcareous		1171
6	Quartzite: Medium gray-brown, we light gray-brown, fine grained, medbedded in beds 2 to 4 ft. thick, crobedded in base, weathers blocky, 1	edium oss-	
<u>.</u>	former	12	1166
	Shale: Black, weathers dark gray, silty, contains abundant pelecypod	Quartaitu redegray	
	fragments	hiq digi	1154
	Quartzite: Light gray, medium grand light brown-gray, weathers me	edium	
	brown, medium and dark red-brow		
	and light brown-gray, fine and med		
	grained, medium bedded to massiv weathers blocky and rounded,	re,	
	scintillating	211	1152
5	Shale (Covered): Black to medium		
	green-gray, weathers to medium a	ind	
	light green-brown soil slope contai	ning	- No.
	fine chips	183	941
4	Quartzite: White to light gray, we light gray, medium and dark red-hand light to dark brown, fine to me grained becomes very coarse grain in basal 3 ft. of sub-unit, thin to the badded in hade 1 to 4 in this badded i	orown, edium ned nick	
	bedded in beds 1 to 4 in. thick in the base to 4 to 6 ft. thick in top, weat slabby and blocky, scintillating, leformer in basal 20 ft., cross-bedd	hers dge	
	throughout	92	758
	Shale: Black, weathers to medium gray chippy soil slope, slightly sil		
(2)	non-calcareous	6	666
	Quartzite: Light to medium gray, weathers light and medium brown a	and	
	The state of the s		

Unit No.	Dε	escription	Thickness in feet	Feet Above Base
	brown-gray, fine medium bedded in thick, weathers s	beds 6 to 24 in.		660
3 05	Shale: Light brow medium red-gray and black in top 43 medium gray and soil slope, clayey	on in basal 20 ft., in overlying 52 ft 30 ft., weathers to medium brown-grand, slightly	······································	erun e
jen Laine	Limestone: Dark medium gray, fine bedded in beds 1 to	gray, weathers grained, thin of thick,	itingeni Marie	14(4)***
1	weathers platy Shale: Variegated yellow-brown, light gray, light gray, red-gray, weather light red-gray fine calcareous, silty; and chert nodules and light yellow-gray.	light yellow, dar nt and medium ye and medium and l es to light gray an e grained soil slop contains siltstone that are light yell	k llow- ight d	146
	diameter			129

Conformable contact

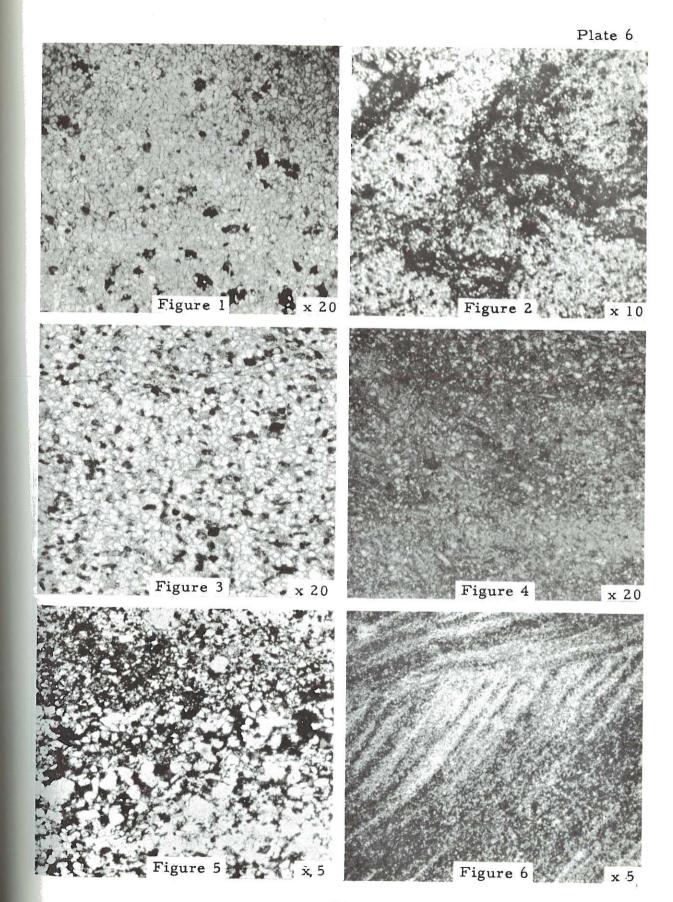
Mississippian

Great Blue Limestone (top)

Limestone: Medium and dark gray, weathers light gray, fine grained, thin bedded in beds 1 to 12 in. thick, weathers platy and flaggy, contains discontinuous black and gray chert beds.

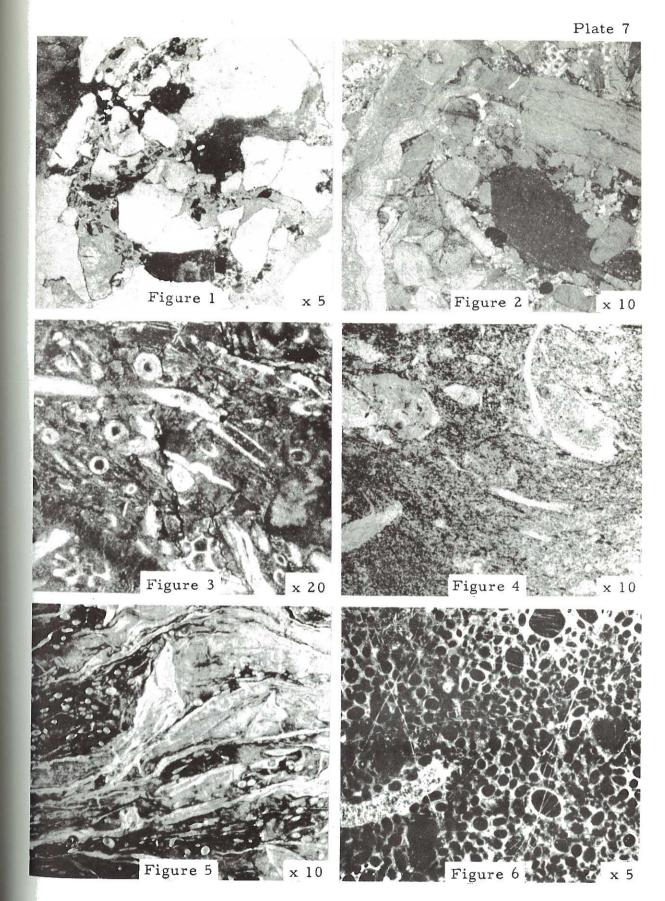
Explanation of Plate 6

- Figure 1 Fine grained quartzite from unit 17 in Soldier Canyon, Oquirrh Mountains, Utah; 800 feet above the base of the formation, magnified 20 times, plain light.
- Figure 2 Fucoidal sandy limestone from the top of unit 32 in Soldier Canyon, Oquirrh Mountains, Utah; 1540 feet above the base of the formation, magnified 10 times, plain light.
- Figure 3 Fine grained calcareous quartzite from the top of unit 27 in Soldier Canyon, Oquirrh Mountains, Utah; 1335 feet above the base of the formation, magnified 20 times, plain light.
- Figure 4 Fine grained sandy limestone from unit 29 in Soldier Canyon, Oquirrh Mountains, Utah; 1435 feet above the base of the formation, magnified 20 times, plain light.
- Figure 5 Medium grained arkosic sandstone from the base of unit 16 in Lake Mountain, Utah;
 1770 feet above the base of the formation, magnified 5 times, plain light.
- Figure 6 Cross-bedded sandy limestone from the top of the formation in Five Mile Pass, Utah; megascopically the cross-beds are evident only on the weathered surface, magnified 5 times, plain light.



Explanation of Plate 7

- Figure 1 Fragmental argillaceous limestone from the top of unit 7 in West Canyon, Oquirrh Mountains, Utah; 785 feet above the base of the formation, magnified 5 times, plain light.
- Figure 2 Crinoidal limestone from unit 2 in Ophir Pass, Oquirrh Mountains, Utah; 60 feet above the base of the formation, magnified 10 times, plain light.
- Figure 3 Hashy fossiliferous limestone from unit 32 in Soldier Canyon, Oquirrh Mountains, Utah; 1513 feet above the base of the formation, magnified 20 times, plain light.
- Figure 4 Fossiliferous limestone from unit 10, medial limestone, in Soldier Canyon, Oquirrh Mountains, Utah; 480 feet above the base of the formation, magnified 10 times, plain light.
- Figure 5 Brachiopod coquina from unit 31 in Soldier Canyon, Oquirrh Mountains, Utah; 1485 feet above the base of the formation, magnified 10 times, plain light.
- Figure 6 Faecal pellet limestone from top of unit 4 in Soldier Canyon, Oquirrh Mountains, Utah; 200 feet above the base of the formation, magnified 5 times, plain light.



web, Upper Missigsippies and Lower Fernsylvenian

SELECTED REFERENCES

- Agatston, R. S., 1957, Pennsylvanian of the Wind River Basin: Wyoming Geol. Assoc. 12th Ann. Field Conf. Guidebook, pp. 29-33.
- Arnold, C. A., 1947, An Introduction to Paleobotany: McGraw & Hill Co., New York and London.
- Axenfeld, Sheldon, 1952, Geology of the North Scraton area, Tooele County, Utah: Univ. of Indiana, (Film E. 84).
- Bacon, C. W., Jr., 1948, Geology of the Confusion Range, west-central Utah: Geology Soc. Am., Bull., Vol. 59, Part 2, pp.1027-1052.
- Baker, A. A., 1947, Stratigraphy of the Wasatch Mountains in the vicinity of Provo, Utah: U. S. Geol. Survey, Oil and Gas Inv. Preliminary Chart 30.
- Baker, A. A., Huddle, J. W., and Kinney, D. M., 1949, Paleozoic geology of north and west sides of the Uinta Basin, Utah: Am. Assoc. Petrol. Geol., Bull., Vol. 24, pp. 617-635.
- Beath, O. A., Gilbert, C. S., and Eppson, H. F., 1939, The use of indicator plants in locating seleniferous areas in Western United States, Part I: Amer. Jour. Bot., Vol. 26, p. 267.
- Bissell, H. J., 1936, Pennsylvanian Stratigraphy in the southern Wasatch Mountains, Utah: Iowas Acad. Sci. Proc., Vol. 43, pp. 239-243.
- Jeff 1950, Carboniferous and Permian stratigraphy of the Uinta Basin area: Guidebook to the Geol. of Utah, No. 5, pp. 71-96.
- Valley Quadrangle, Utah: Am. Assoc. Petrol. Geol., Bull., 36, pp. 575-634.

- Bissell, H. J., and Hansen, G. H., 1935, The Mississippian-Pennsylvanian contact in the central Wasatch Mountains, Utah. Utah Acad. Sci., Arts, and Letters, Vol. 12, p. 163.
- Bruchner, W. D., 1953, Cyclic calcareous sedimentation as an index of climatic variations in the past: Jour. of Sed. Petrol., Vol. 23, No. 4, pp. 235-237.
- Bullock, K. C., 1951, Geology of Lake Mountain, Utah: Utah Geol. and Min. Surv. Bull., No. 41, pp. 9-32.
- Calderwood, K. W., 1951, Geology of the Cedar Valley Hills area, Lake Mountain, Utah: The Compass of Sigma Gamma Epsilon, Vol. 29, No. 1, pp. 21-32.
- Cooper, C. L., 1947, Upper Kinkaid (Miss) Microfauna from Johnson County, Illinois: Jour. Paleo., Vol. 21, pp. 81-94.
- Cooper, G. A., 1937, Brachiopod Ecology and Paleoecology: Rept. Comm. Paleoecology, National Research Council, pp. 20-53.
- Croft, M. G., 1956, Geology of the Northern Onaqui Mountains, Tooele County, Utah: Brigham Young University Research Studies, Geology Series, Vol. 3, No. 1, p. 45.
- Dapples, E. C., Drumbein, W. O., and Sloss, L. L., 1948, Tectonic control of lithologic associations: Am. Assoc. of Petrol. Geol. Bull., Vol. 32, pp. 1924-47.
- Am. Assoc. Petrol. Geol. Bull., Vol. 33, pp. 1859-1891.
- Degens, E. T., Williams, E. G., and Keith, M. L., 1957, Environmental studies of carboniferous sediments, Part I: Geochemical criteria for differentiating marine from freshwater shales: Am. Assoc. Petrol. Geol. Bull., Vol. 41, No. 11, pp. 2427-2455.
- Dott, R. H. Jr., 1955, Pennsylvanian stratigraphy of Elko and Northern Diamond Ranges, northeastern Nevada: Am. Assoc. Petrol. Geol. Bull., Vol. 39, No. 11, pp. 2211-2305.
- , 1958, Cyclic patterns in mechanically deposited

 Pennsylvanian limestones of northeastern Nevada: Jour. of
 Sed. Petrology, Vol. 28, No. 1, pp. 3-14.

- Elias, M. K., 1937, Depth of deposition of the Big Blue (late Paleozoic) sediments in Kansas: Geol. Soc. Am. Bull., Vol. 48, pp. 403-432.
- , 1956, Upper Mississippian and Lower Pennsylvanian
 Formations of South-Central Oklahoma: Am. Assoc. of Petrol.
 Geol. of So. Okla. Reprint, p. 134.
- , 1957, Late Mississippian fauna from the Redoak Hollow Formation of southern Oklahoma, Part 1: Bryozoa, Jour. Paleo., Vol. 31, No. 2, pp. 370-427.
- Formation of southern Oklahoma, Part II: Brachiopoda, Jour. Paleo., Vol. 31, No. 3, pp. 487-527.
- Formation of southern Oklahoma, Part III: Pelecypoda, Jour. Paleo., Vol. 31, No. 4, pp. 737-784.
- Fenton, C. L., 1935, Viewpoints and objects of paleoecology: Jour. of Paleo., Vol. 9, No. 1, pp. 63-78.
- Gilluly, James, 1932, Geology and ore deposits of the Stockton and Fairfield Quadrangles, Utah: U. S. Geol. Surv. Prof. Paper 173, pp. 31-34.
- Grim, R. E., 1951, The depositional environment of red and green shales: Jour. Sed. Petrology, Vol. 21, pp. 226-232.
- Hager, D. S., 1928, Factors affecting the color of sedimentary rocks: Am. Assoc. Petrol. Geol. Bull., Vol. 12, pp. 901-938.
- Hebertson, K. M., 1950, Origin and composition of the Manning Canyon Formation in central Utah: Unpublished M. S. Thesis, Brigham Young University.
- _______, 1957, Some characteristics of the Manning Canyon
 Formation in central Utah: Intermountain Assoc. of Petrol.
 Geol. 8th Ann. Field Conf. Guidebook, pp. 78-81.
- Hernon, R. M., 1935, The Paradise Formation and its fauna: Jourof Paleo., Vol. 9, No. 8, pp. 653-696.
- Hyatt, E. P., 1956, Clays of Utah County, Utah: Utah Geol. Miner. Surv., Bull. 55, p. 83.

- Keller, W. D., 1953, Illite and montmorillonite in green sedimentary rocks: Jour. Sed. Petrology, Vol. 23, No. 1, pp. 3-9.
- Kelley, V. C., 1956, Thickness of strata: Jour. Sed. Petrology, Vol. 26, No. 4, pp. 289-300.
- Krumbein, W. C., and Sloss, L. L., 1951, Stratigraphy and Sedimentation: W. H. Freeman and Co., San Francisco, California.
- McFarland, C. R., 1955, Geology of the West Canyon area, northwestern Utah County, Utah: Brigham Young University Research Studies, Geology Series, Vol. 2, No. 3, pp. 21.
- McKee, E. D., and Weir, G. W., 1953, Terminology for stratification and cross-stratification in sedimentary rocks: Geol. Soc. Am. Bull., Vol. 64, pp. 381-390.
- Miller, A. K., and Youngquist, Walter, 1948, The cephalopod fauna of the Mississippian, Barnett Formation of central Texas: Jour. of Paleo., Vol. 22, pp. 649-671.
- Moore, R. C., Lalicker, C. B., and Fisher, A. G., 1952, Invertebrate Fossils: McGraw-Hill Book Co., New York.
- Noble, L. F., 1922, A section of the Paleozoic Formations of the Grand Canyon at the Bass Trail: U. S. Geol. Surv. Prof. Paper 131, pp. 23-73.
- Nolan, T. B., 1935, The Gold Hill Mining District, Utah: U. S. Geol. Surv. Prof. Paper 177.
- _____, 1943, The Basin and Range Province in Utah, Nevada, and California: U.S. Geol. Surv. Prof. Paper 197 D, pp. 141-196.
- Ogden, Lawrence, 1951, Mississippian and Pennsylvanian stratigraphy, Confusion Range, west-central Utah: Am. Assoc. Petrol. Geol. Bull., Vol. 35, No. 1, pp. 62-82.
- Olson, R. H., 1956, Geology of Promontory Range: Guidebook to Geol. of Utah, No. 11, pp. 41-75.
- Ornelas, R. H., 1953, Clay deposits of Utah County: Unpublished M. S. Thesis, Brigham Young University.
- Parks, J. M., Jr., 1951, Corals from the Brazer Formation of northern Utah: Jour. of Paleo., Vol. 25, No. 2, pp. 171-186.

- Peace, F. S., 1956, History of exploration for oil and gas in Box Elder County, Utah, and vicinity: Guidebook to the Geol. of Utah, No. 11, pp. 17-31.
- Pettijohn, F. J., 1948, Sedimentary Rocks: Harper and Brothers, New York.
- Robertson, J. A., 1941, A study of the occurrence of selenium and the possibility of selenium poisoning in Utah: Unpublished M. S. Thesis, Brigham Young University, pp. 63-72.
- Sadlick, Walter, 1955, The Mississippian and Pennsylvanian boundary in northeastern Utah: Unpublished M. S. Thesis, University of Utah, 1.77.
- ______, 1955, Carboniferous formations of northeastern Uinta Mountains: Wyo. Geol. Assoc. Guidebook, 10th Ann. Field Conf., pp. 49-59.
- , 1957, Regional relations of carboniferous rocks of northeastern Utah, Guidebook to the geology of the Uinta Basin: Intermountain Assoc. Petrol. Geol., 8th Ann. Field Conf., pp. 56-78.
- Schemel, M. P., 1950, Carboniferous plant spores from Daggett County, Utah: Jour. Paleo., Vol. 24, pp. 232-44.
- Shimer, H. W., and Shrock, R. R., 1949, Index Fossils of North America: John Wiley and Sons, Inc., New York.
- Siever, Raymond, 1953, Petrology and sedimentation of Upper Chester sandstones: Jour. Sed. Petrology, Vol. 23, No. 4, pp. 207-219.
- Spreng, A. C., 1953, Mississippian cyclic sedimentation, Sunwapta Pass area, Alberta, Canada: Am. Assoc. Petrol. Geol. Bull., Vol. 37, pp. 665-689.
- Spurr, J. E., 1895, Economic geology of the Mercur Mining District, Utah: U. S. Geol. Serv. 16th Ann. Report, Part 2, pp. 375-377.
- Stokes, W. L., 1951, Paleozoic stratigraphy of Great Basin (abstract): Am. Assoc. Petrol. Geol. Bull., Vol. 35, p. 1107.
- Sutton, A. H., 1938, Taxonomy of Mississippian Productidae: Jour. Paleo., Vol. 12, No. 6, pp. 537-569.

- Thompson, M. L., 1953, Primitive Fusulinella from southern Missouri: Jour. of Paleo., Vol. 27, pp. 321-27.
- Trask, P. D., 1939, Recent Marine Sediments: Am. Assoc. Petrol. Geol., George Banta Publishing Company, Menasha, Wisconsin.
- Twenhopel, W. H., 1932, Treatise on Sedimentation: The Williams and Wilkins Company, Baltimore, New York.
- Unterman, G. E., and Unterman, B. R., 1944, Geology of Dinosaur National Monument and vicinity, Utah and Colorado: Utah Geol. and Miner. Surv., Bull. 42, p. 31.
- Wanless, H. R., 1950, Late Paleozoic cycles of sedimentation in the United States: International Geological Congress, Report of the 18th Session, Great Britain 1948, Part IV.
- Weller, J. M., 1948, Correlation of the Mississippian formations of North America: Geol. Soc. Am. Bull., Vol. 59, pp. 91-196.
- Westgate, L. G., and Knopf, Adolph, 1932, Geology and ore deposits of the Pioche District, Nevada: U. S. Geol. Surv. Prof. Paper 171, pp. 21-22.
- Wheeler, H. E., and Murray, H. H., 1957, Base-level control patterns in cyclothemic sedimentation: Am. Assoc. Petrol. Geol. Bull., Vol. 41, No. 9, pp. 1985-2011.
- White, David, 1899, Fossil flora of the lower coal measures of Missouri: U. S. Geol. Surv. Mono. 37.
- Williams, J. S., 1948, Geology of the Paleozoic rocks, Logan Quadrangle: Geol. Soc. Am. Bull., Vol. 59, pp. 1121-1164.
- , 1948, Mississippian and Pennsylvanian boundary problems in the Rocky Mountain region: Jour. Geol., Vol. 56, No. 4, pp. 327-351.
- Young, J. C., 1955, Geology of the southern Lakeside Mountain, Utah: Utah Geol. & Miner. Surv., Bull. 56.
- Youngquist, Walter, 1949, The cephalopod fauna of the White Pine Shale of Nevada: Jour. Paleo., Vol. 23, No. 3, pp. 276-303.

- Youngquist, Walter, 1951, Notes on the Mississippian rocks of the Great Basin with special reference to the Confusion Range, Utah: Guidebook to the Geol. of Utah, No. 6, pp. 54-59.
 - Zeller, Doris E. N., 1953, Endothyroid foraminifera and ancestral fusulinids from the type Chester (Up. Miss.): Jour. Paleo., Vol. 27, pp. 183-199.
 - Zeller, E. J., 1950, Stratigraphic significance of Mississippian endothyroid foraminifera: Univ. Kansas Contr. Paleo., Protozoa, Art. 4, pp. 1-23.
 - , 1957, Mississippian endothyroid foraminifera from the Cordilleran Geosyncline: Jour. Paleo., Vol. 34, No. 4, pp. 679-704.